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# The Study of Crystal Structure on Grossular–Andradite Solid Solution

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**Abstract:** The effects of  $\text{Al}^{3+}$ – $\text{Fe}^{3+}$  substitution on 10 synthesized garnet samples along the grossular–andradite binary solid solution were investigated using both powder and single-crystal X-ray diffraction. Results showed that cell volume increased with andradite content. Small negative excess volume was observed in the Al-rich samples. By measuring the bond length, polyhedral volume, octahedral distortion parameter ( $\sigma$ ), and tetrahedral rotation angle ( $\alpha$ ), we determined that the distortion occurred on the dodecahedral site. The width of the diffraction peaks was obviously related to the composition of the solid solution. Full width at half maximum of diffraction peaks was used to calculate the microstrain, which may have a relationship with enthalpy of mixing.

**Keywords:** grossular–andradite; X-ray diffraction; bond length; excess volume; distortion

## 1. Introduction

Garnet ( $\text{X}_3\text{Y}_2\text{Si}_3\text{O}_{12}$ ) is a significant phase in the upper mantle and transition zone because it can be stable over a broad temperature and pressure range [1,2]. Natural garnets always exhibit extensive substitution and exist as complex solid solutions. Due to different ionic radii, substitution between cations changes the microstructure of garnets, which may further affect the relevant physical–chemical properties [3–7].

Previous studies on the mixing behavior of binary garnet solid solutions have mainly focused on the effects of divalent cation substitutions ( $\text{Mg}^{2+}$ – $\text{Fe}^{2+}$ – $\text{Ca}^{2+}$ ) [5,8–13]. It has been shown that the larger the size difference between the cations that mix on the X sites, the larger will be the deviation from the ideal mixing solid solution [13]. The excess volume of pyrope (Pyr,  $\text{Mg}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ )–almandine (Alm,  $\text{Fe}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ ), almandine–grossular (Gro,  $\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ ), and pyrope–grossular solid solutions can reach 3.0, 4.3, and 7.5  $\text{\AA}^3/\text{cell}$ , respectively. The nonlinear variation of thermal expansion, bulk modulus, and thermodynamic properties (e.g., H and S) for pyrope–grossular solid solution is considered to be related to the nonideal behavior in structure [5,14–16].

According to Boffa-Ballaran and Woodland (2006) [17] and McAloon and Hofmeister (1995) [18], unit-cell parameters have a roughly linear correlation with garnet composition in spite of the large radius difference between  $\text{Fe}^{3+}$  and  $\text{Al}^{3+}$  on the octahedral site. However, Heuss–Aßbichler and Fehr (1997) [19] observed a negative and a positive deviation from the ideal mixing cell volume at grossular-rich and andradite (And,  $\text{Ca}_3\text{Fe}_2\text{Si}_3\text{O}_{12}$ )-rich compositions, respectively. Therefore, more global observations of the mixing behavior of grossular–andradite solid solution should be carried out.

In this study, we investigated the grossular–andradite binary system to examine the effects of changing octahedral cations. We synthesized garnet solid solutions along the grossular–andradite

binary system at 3 GPa and 1100–1200 °C and investigated the effects of Al<sup>3+</sup>–Fe<sup>3+</sup> substitution on garnet structures using both powder X-ray diffraction (PXRD) and single-crystal XRD methods. The variation in cell volume and bond lengths of polyhedrons was measured. Excess volume and microstrain were calculated by unit-cell parameters and full width at half maximum (FWHM) of the diffraction peaks, respectively. The results show that excess volume may have a relationship with dodecahedral distortion.

## 2. Experimental Methods

The raw materials (analytical grade reagents SiO<sub>2</sub>, CaCO<sub>3</sub>, Fe<sub>2</sub>O<sub>3</sub>, and Al<sub>2</sub>O<sub>3</sub>) were dried at 450 °C in open air for 12 h. After that, these chemical compounds were mixed and pressed into pellets of grossular and andradite composition, respectively. These pellets were decarbonated at 1000 °C in high-*T* furnace. Then, the degassed pellets were crushed into fine powders and mixed at appropriate proportions, which were further used for high-*P* experiments to synthesize the Gro–And solid solutions (Table 1).

**Table 1.** Starting compositions used in this study.

Target <i>x</i> <sup>a</sup>	Composition	SiO <sub>2</sub> <sup>b</sup>	Al <sub>2</sub> O <sub>3</sub> <sup>b</sup>	Fe <sub>2</sub> O <sub>3</sub> <sup>b</sup>	CaO <sup>b</sup>
0.0	Gro <sup>c</sup> <sub>100</sub> And <sup>d</sup> <sub>0</sub>	40.02	22.64	0.00	37.35
0.1	Gro <sub>90</sub> And <sub>10</sub>	39.51	20.11	3.50	36.88
0.2	Gro <sub>80</sub> And <sub>20</sub>	39.02	17.66	6.91	36.42
0.3	Gro <sub>70</sub> And <sub>30</sub>	38.53	15.26	10.24	35.97
0.4	Gro <sub>60</sub> And <sub>40</sub>	38.06	12.92	13.49	35.53
0.5	Gro <sub>50</sub> And <sub>50</sub>	37.61	10.64	16.66	35.10
0.6	Gro <sub>40</sub> And <sub>60</sub>	37.16	8.41	19.75	34.68
0.7	Gro <sub>30</sub> And <sub>70</sub>	36.72	6.23	22.77	34.27
0.8	Gro <sub>20</sub> And <sub>80</sub>	36.29	4.11	25.72	33.88
0.9	Gro <sub>10</sub> And <sub>90</sub>	35.88	2.03	28.61	33.49
1.0	Gro <sub>0</sub> And <sub>100</sub>	35.47	0.00	31.42	33.11

<sup>a</sup>  $x = \text{Fe}^{3+}/(\text{Fe}^{3+} + \text{Al}^{3+})$  (atomic ratio); <sup>b</sup> Content in wt %; <sup>c</sup> Gro stands for grossular Ca<sub>3</sub>Al<sub>2</sub>Si<sub>3</sub>O<sub>12</sub>; <sup>d</sup> And stands for andradite Ca<sub>3</sub>Fe<sub>2</sub>Si<sub>3</sub>O<sub>12</sub>.

High-*P* syntheses were conducted at 3 GPa and 1100–1200 °C using the cubic press at the High-Pressure Laboratory of Peking University [20]. The new cell arrangement reported in He et al., (2014) [21] was used (BJC-11 [20]). In order to host our large Pt capsules (OD = 5 mm, ID = 4.75 mm, and L = 7 mm), which were arc-welded to contain the experimental precursors, the dimensions of the small parts inside the graphite heater were accordingly modified. In addition, all small parts made of *h*-BN were replaced by MgO, so the oxygen fugacity in this experimental assembly was significantly increased. As illustrated in Liu et al., (2012) [20], no pressure correction is necessary for this cell arrangement, and the accuracy of the experimental *P* measurements should be ±0.1 GPa. A 1 mm thick MgO disk was used to prevent the capsule being in direct contact with the thermocouple (a type C W5%Re–W26%Re thermocouple). The fluctuation of the nominal *T* readings of the thermocouple was always less than ±1 °C. Due to the large size of the Pt capsules, the real temperature exerted on the experimental sample varied from ~70 °C higher to ~50 °C lower than the nominal thermocouple reading [22].

The experimental products from LMD581, LMD584, and LMD589 were simply dug out of the Pt capsules without any attention paid to observe the presence/absence of gas. Small parts of the synthesized products were arbitrarily picked, mounted in epoxy, polished, and then examined with scanning electron microscopy (SEM). Subsequent electron microprobe analysis (EMPA) was employed for composition analysis. After completing all the above experiments, suitable crystals were carefully picked out for single-crystal XRD. The remaining samples were grinded into fine powders for PXRD analyses (Pt powder added as internal standard). In contrast, the experimental products from other synthesizing experiments were first sanded open with dry silicon carbide abrasive paper and checked

to determine whether there was a gas phase by signs of bubbling (Table 2). The capsules were then cut longitudinally by a low-speed diamond saw at positions about 1 mm away from the capsule edge. These small portions of the samples were mounted in epoxy, polished, and prepared for SEM, EMPA, and single-crystal XRD. The large portions were grinded into fine powders for powder XRD analyses.

**Table 2.** Experimental conditions and results (all experiments were performed at 3 GPa).

Run #	X <sup>a</sup>	T (°C)	t (h) <sup>b</sup>	Phase Assemblage	Notes
LMD611	0.0	1200	24	Gt + melt	Gt size ~20–100 μm; vol % of melt ~5% <sup>c</sup>
LMD634	0.1	1200	28	Gt + melt + gas <sup>d</sup>	Gt size ~20–500 μm; vol % of melt ~20%; almost all Fe lost to the Pt capsule
LMD584	0.2	1200	24	Gt + melt	Gt size ~30–50 μm; vol % of melt ~5%
LMD618	0.3	1200	24	Gt + melt + gas	Gt size ~20–200 μm; vol % of melt ~5%
LMD589	0.4	1200	24	Gt + melt	Gt size ~30–80 μm; vol % of melt ~5%
LMD622	0.5	1200	28	Gt + melt + gas	Gt size ~20–50 μm; vol % of melt ~5%
LMD631	0.6	1200	28	Gt + melt + gas	Gt size ~40–60 μm; vol % of melt ~5%
LMD613	0.7	1200	24	Gt + melt + gas	Gt size ~50 μm; vol % of melt ~5%
LMD581	0.8	1200	24	Gt + melt	Gt size ~50–100 μm; vol % of melt < 5%
LMD632	0.9	1100	28	Gt + melt + gas	Gt size ~20–100 μm; vol % of melt ~5%
LMD625	1.0	1100	28	Gt + melt	Gt size ~100–500 μm; vol % of melt ~30%

<sup>a</sup> Target  $x$ , with  $x = \text{Fe}^{3+}/(\text{Fe}^{3+} + \text{Al}^{3+})$ ; <sup>b</sup> Experimental duration in hours; <sup>c</sup> Melt volume contents in notes were estimated from electron backscattered images; <sup>d</sup> Gas was identified by the observation of bubbling from the Pt capsule when the capsule was sanded open with dry silicon carbide abrasive paper because water and oxygen are the main components of the gas. We did not check the presence of gas for samples LMD581, LMD584, and LMD589.

The appearance of the samples was checked using the scanning electron microscope (FEI Quanta 650 FEG) at the School of Earth and Space Sciences, Peking University. The working conditions were 10 kV (accelerating voltage) and 10 μs (dwell time), with other working parameters shown in Figure 1.

The compositional data were collected using the JEOL JXA-8230 electron probe microanalyzer at the School of Earth and Space Sciences, Peking University. The analytical conditions were 15 kV (accelerating voltage) and 10 nA (beam current). Beam diameter was 1–2 μm for garnet crystals and 5–20 μm for melt; in particular, a 200 μm diameter was used for large areas of melt in samples LMD586 and LMD625. Calibration was based on optimization to some standards provided by the SPI Corporation (USA), with diopside for Ca, jadeite for Si and Al, and hematite for Fe. Data correction was performed with the PRZ method. Ferric iron content was calculated by charge balance and setting the sum of oxygen ions to 12. To obtain an average composition, more than 10 analyses were performed on various grains from different locations in each sample.

Powder XRD patterns were collected at ambient temperature and pressure using an X'Pert Pro MPD diffractometer with an X'Celerator detector at the School of Earth and Space Sciences, Peking University. The measuring conditions of the X-ray powder diffraction were Cu Kα radiation, 40kV and 40 mA, 10–90° (2θ) scanning range, 0.017° step size, and 20 s per step. We analyzed the patterns using PeakFit v4.12 software. Diffraction peaks were fitted using Gauss + Lorentz peak shape function to obtain the positions and FWHM of the diffraction peaks. Peak positions were further used to refine the unit-cell parameters of each sample by UnitCell [23].

The single crystals were immersed in silicone oil, and suitable samples were picked out under an optical microscope for single-crystal X-ray diffraction analysis. Intensity data were collected on a Bruker Smart ApexIII Quazar microfocussed diffractometer using Mo Kα radiation ( $\lambda = 0.71073 \text{ \AA}$ ). The raw data were processed and corrected for the absorption effects using SAINT+ v8.38A and SADAB embedded in Bruker Apex3. The initial structures were set up based on garnet structure (space group Ia-3d) and refined by a full-matrix least-squares method using the SHELXT v2018.3 software included in the SHELXTL package. The heaviest atoms were first located unambiguously in the Fourier maps, and the O atoms were then found in the subsequent difference maps. All atoms were refined with anisotropic displacement parameters. The refinement results are listed in Table 3.

**Table 3.** Atom fraction coordinate <sup>a</sup> and refinement results.

Comp. <sup>b</sup>	Oxygen Atom Coordinate			Refinement Results		
	<i>x</i>	<i>y</i>	<i>z</i>	<i>R</i> <sub>1</sub>	<i>wR</i> <sub>2</sub>	<i>Goof</i>
0.00 (0) <sup>c</sup>	0.0384 (1)	0.5453 (1)	0.3487 (1)	0.0267	0.0769	1.012
0.15 (2)	0.0387 (2)	0.5456 (2)	0.3480 (1)	0.0269	0.0780	1.001
0.26 (1)	0.0383 (5)	0.5465 (6)	0.3475 (5)	0.0682	0.1517	1.077
0.33 (2)	0.0379 (4)	0.5464 (4)	0.3471 (4)	0.0521	0.0937	1.059
0.45 (2)	0.0390 (3)	0.5467 (3)	0.3466 (3)	0.0377	0.0938	1.015
0.54 (2)	0.0389 (3)	0.5471 (3)	0.3464 (3)	0.0377	0.0657	1.014
0.70 (2)	0.0394 (4)	0.5475 (3)	0.3456 (3)	0.0423	0.1048	1.010
0.75 (1)	0.0392 (1)	0.5479 (1)	0.3455 (1)	0.0189	0.0541	1.045
0.90 (1)	0.0393 (2)	0.5485 (2)	0.3449 (2)	0.0274	0.0746	1.091
0.99 (0)	0.0395 (3)	0.5488 (2)	0.3446 (2)	0.0371	0.0940	1.006

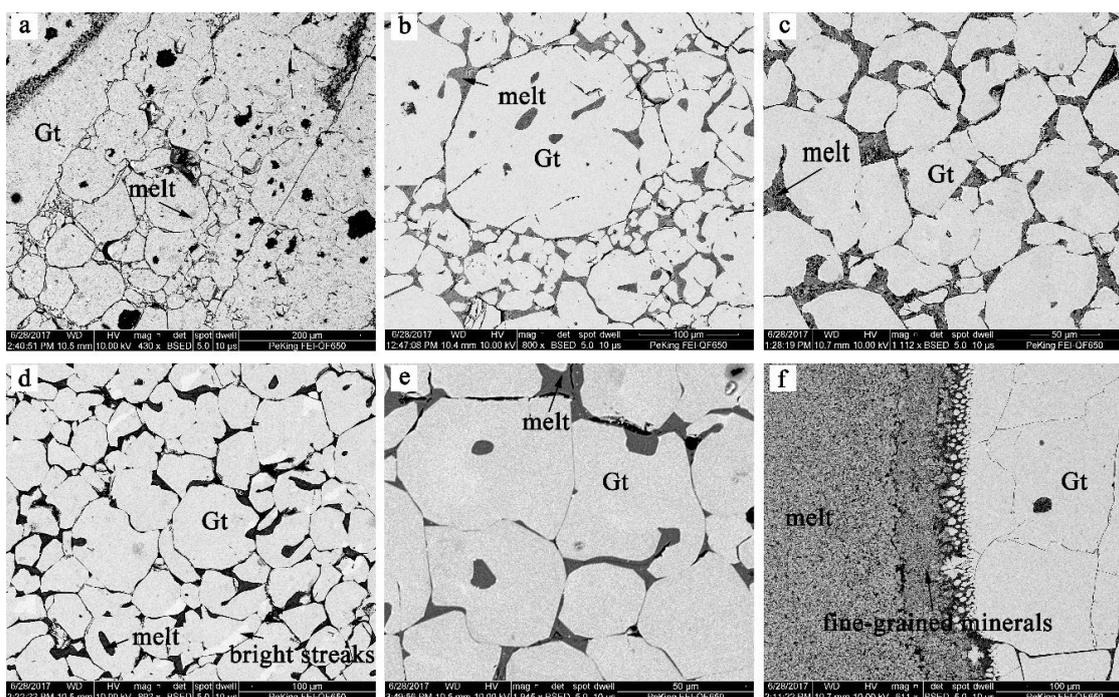
<sup>a</sup> Coordinate of Ca was fixed as (1/8 0 1/4), Si was fixed as (3/8 0 1/4), Al/Fe was as fixed (0 0 0); <sup>b</sup> Comp. stands for composition ( $\text{Fe}^{3+}/(\text{Fe}^{3+} + \text{Al}^{3+})$ ); <sup>c</sup> Number in the parentheses represents standard deviation in the rightmost digit.

### 3. Results

#### 3.1. Phase Compositions

As shown by the backscattered electron (BSE) images in Figure 1, two phases—garnet and melt—were observed. In most cases, the size of the garnet was between 20 and 100  $\mu\text{m}$ ; however, a larger size of about 100–500  $\mu\text{m}$  was observed in sample LMD625, possibly due to a higher degree of melting (Figure 1f). The glass phase constituted around 1–10% of the experimental product and existed in the gaps between garnet crystals and along the cracks of garnet grains. A few bright-colored minerals between the garnet grains in the BSE images might have been a Ca-rich quench product formed from the melt, which was always less than 1%. However, in sample LMD625 with pure andradite starting composition, we observed about 30% melt by volume in the center of the capsule, which might have been caused by the lower melting point of andradite compared to grossular (Figure 1f). Fine mineral grains with poor crystallinity could be found between the melt and garnet crystals, formed during the quenching process.

As Fe can be lost to the Pt capsules [24], most samples had a higher grossular component compared to the target value according to EMPA analyses. Due to the lower andradite content, almost no Fe was observed in the garnet of sample LMD634, although our target  $\text{Fe}^{3+}/(\text{Al}^{3+} + \text{Fe}^{3+})$  ratio was 0.1 (Table 4). Therefore, sample LMD634 was not considered in the following studies.



**Figure 1.** Electron backscattered images showing the textures of the synthesized samples: (a) LMD611, (b) LMD618, (c) LMD622, (d) LMD631, (e) LMD613, and (f) LMD625. Garnet (Gt) and melts are shown in each figure. A little bit of unreacted starting material (less than 1%) was contained in large grains of garnet in sample LMD611.

Garnet solid solutions were compositionally homogeneous according to EMPA analyses, with the standard deviation of  $\text{Fe}^{3+}/(\text{Al}^{3+} + \text{Fe}^{3+})$  ratio ranging from 0 to 0.02. In sample LMD631, some bright streaks were observed on the garnet grains in BSE images scanned by SEM (Figure 1d). However, the light and dark areas had rather similar compositions, with  $\text{Fe}^{3+}/(\text{Al}^{3+} + \text{Fe}^{3+})$  ratios of 0.54(2) and 0.53(2), respectively. In fact, the brightness difference could not be identified in the backscattered electron images using the EMPA. We did not distinguish the light and dark areas thereafter. The degree of Fe reduction is a key issue for the study of Gro–And solid solutions. The content of ferrous iron was estimated by charge balance. As demonstrated by Xu et al., (2017) [25], charge balance method has a high accuracy for garnets. Our calculations showed that, with the exception of low-Fe sample LMD584, the ratio of ferric to ferrous iron was always >95%.

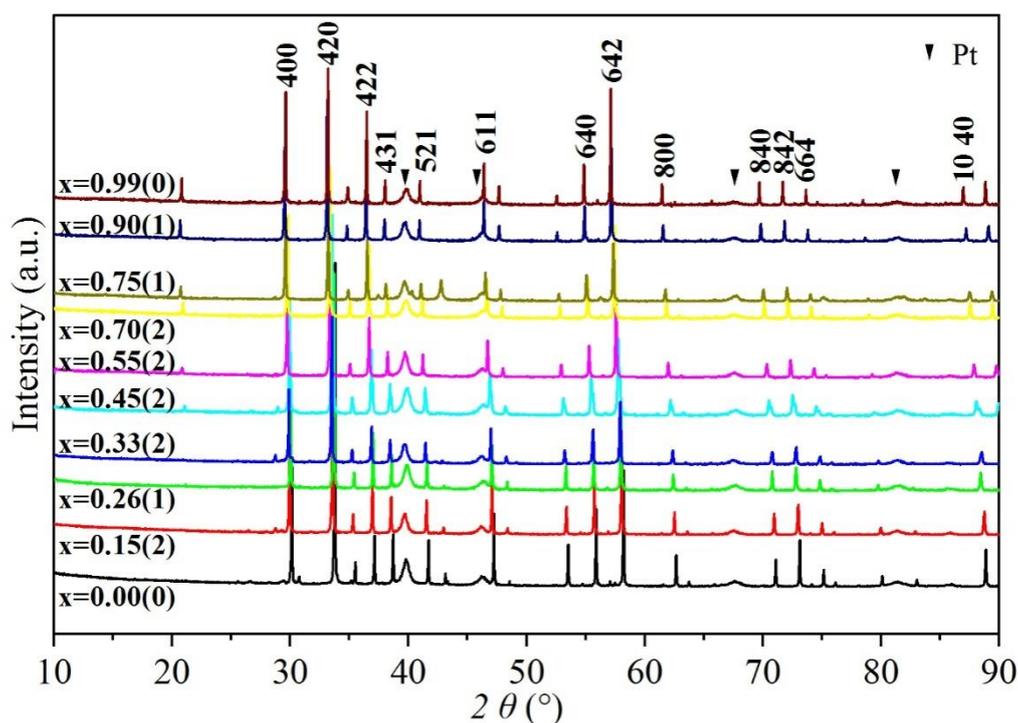
Table 4. Compositions of garnet.

Run #	LMD611	LMD634	LMD584	LMD618	LMD589	LMD622	LMD631	LMD613	LMD581	LMD632	LMD625
Target <i>x</i>	0	0.1	0.2	0.3	0.4	0.5	0.6	0.7	0.8	0.9	1.0
EMPA #	11	18	19	23	12	20	28	24	10	28	12
Wt %											
SiO <sub>2</sub>	39.92(15) <sup>a</sup>	40.27(21)	39.27(16)	38.49(17)	38.23(28)	37.35(23)	37.06(29)	35.86(17)	36.05(36)	35.54(29)	34.73(23)
Al <sub>2</sub> O <sub>3</sub>	22.49(20)	22.23(22)	18.88(51)	16.49(28)	14.77(48)	12.11(52)	9.69(52)	6.48(36)	5.12(34)	2.01(20)	0.15(2)
Fe <sub>2</sub> O <sub>3</sub> <sup>b</sup>	0.00(0)	0.40(34)	5.30(55)	8.87(44)	11.27(81)	15.67(67)	18.14(71)	22.89(58)	24.38(58)	28.68(50)	31.85(31)
FeO <sup>b</sup>	0.00(0)	0.32(27)	0.65(40)	0.46(36)	0.90(45)	0.24(25)	0.70(36)	0.05(9)	1.02(44)	0.66(47)	0.00(0)
CaO	37.76(22)	37.30(36)	36.07(35)	35.53(26)	34.92(33)	34.67(26)	33.98(23)	33.63(25)	32.80(39)	32.61(28)	32.62(14)
Total	100.23(37)	100.61(32)	100.26(41)	99.93(45)	100.08(54)	100.11(50)	99.64(42)	99.01(47)	99.47(79)	99.59(55)	99.44(51)
Cations on the basis of 12 oxygen											
Si	2.99(1)	3.01(1)	3.00(2)	2.99(1)	2.99(1)	2.96(1)	2.99(2)	2.96(2)	2.99(2)	2.99(2)	2.96(1)
Al	1.98(2)	1.96(1)	1.70(4)	1.51(2)	1.36(4)	1.13(4)	0.92(5)	0.63(3)	0.50(3)	0.20(2)	0.01(0)
Fe <sup>3+</sup>	0.00(0)	0.03(2)	0.31(3)	0.52(3)	0.66(5)	0.94(5)	1.10(5)	1.45(4)	1.52(4)	1.82(3)	2.07(2)
Fe <sup>2+</sup>	0.00(0)	0.02(2)	0.04(3)	0.03(2)	0.06(3)	0.02(2)	0.05(2)	0.00(1)	0.07(3)	0.05(3)	0.00(0)
Ca	3.03(1)	2.99(2)	2.95(2)	2.95(1)	2.93(3)	2.95(2)	2.94(2)	2.98(2)	2.91(2)	2.94(2)	2.98(1)
Total	8.00(0)	8.00(1)	8.00(1)	8.00(1)	8.00(0)	8.00(1)	8.00(0)	8.02(2)	8.00(0)	8.00(0)	8.02(1)
Fe <sup>3+</sup> /ΣFe	–	–	0.88(7)	0.95(4)	0.92(4)	0.98(2)	0.96(2)	1.00(0)	0.96(2)	0.98(2)	1.00(0)
Observed <i>x</i>	0.00(0)	0.01(1)	0.15(2)	0.26(1)	0.33(2)	0.45(2)	0.54(2)	0.70(2)	0.75(1)	0.90(1)	0.99(0)
X <sup>2+</sup>	3.03(1)	3.01(1)	2.99(1)	2.98(1)	2.99(1)	2.96(1)	2.98(2)	2.98(1)	2.98(2)	2.99(2)	2.98(1)
Y <sup>3+</sup>	1.98(2)	1.98(2)	2.00(3)	2.03(3)	2.02(2)	2.07(3)	2.02(3)	2.08(4)	2.02(3)	2.02(3)	2.08(2)
Z <sup>4+</sup>	2.99(1)	3.01(1)	3.00(2)	2.99(1)	2.99(1)	2.96(1)	2.99(2)	2.96(2)	2.99(2)	2.99(2)	2.96(1)
Si	2.99(1)	3.01(1)	3.00(2)	2.99(1)	2.99(1)	2.96(1)	2.99(2)	2.96(2)	2.99(2)	2.99(2)	2.96(1)
Al	1.98(2)	1.96(1)	1.70(4)	1.51(2)	1.36(4)	1.13(4)	0.92(5)	0.63(3)	0.50(3)	0.20(2)	0.01(0)
Fe <sup>3+</sup>	0.00(0)	0.03(2)	0.31(3)	0.52(3)	0.66(5)	0.94(5)	1.10(5)	1.45(4)	1.52(4)	1.82(3)	2.07(2)
Observed <i>x</i>	0.00(0)	0.01(1)	0.15(2)	0.26(1)	0.33(2)	0.45(2)	0.54(2)	0.70(2)	0.75(1)	0.90(1)	0.99(0)
X <sup>2+</sup>	3.03(1)	3.01(1)	2.99(1)	2.98(1)	2.99(1)	2.96(1)	2.98(2)	2.98(1)	2.98(2)	2.99(2)	2.98(1)
Y <sup>3+</sup>	1.98(2)	1.98(2)	2.00(3)	2.03(3)	2.02(2)	2.07(3)	2.02(3)	2.08(4)	2.02(3)	2.02(3)	2.08(2)
Z <sup>4+</sup>	2.99(1)	3.01(1)	3.00(2)	2.99(1)	2.99(1)	2.96(1)	2.99(2)	2.96(2)	2.99(2)	2.99(2)	2.96(1)

<sup>a</sup> Number in the parentheses represents standard deviation in the rightmost digit. <sup>b</sup> Fe<sub>2</sub>O<sub>3</sub> and FeO are calculated using the charge balance method.

### 3.2. Unit-Cell Parameters

Powder XRD patterns of Gro–And garnet solid solutions were obtained at ambient condition.  $K\alpha_2$  peaks were eliminated during data processing by Jade 5.0 software. Due to narrow and sharp shapes, peaks of garnets could easily be identified and are marked with number of plane in Figure 2. The Pt powder was used as an internal standard to correct diffraction peaks in the XRD experiment. Four Pt peaks could be observed with Bragg angle of  $10^\circ$ – $90^\circ$  (marked by solid inverted triangles in Figure 2). Because of broad diffraction peaks, Pt correction was not used in this study to avoid fitting error. In addition, there were some peaks with low Bragg angles, which might be Ca-, Si-rich minerals (e.g., wollastonite) crystallized during melt quenching.



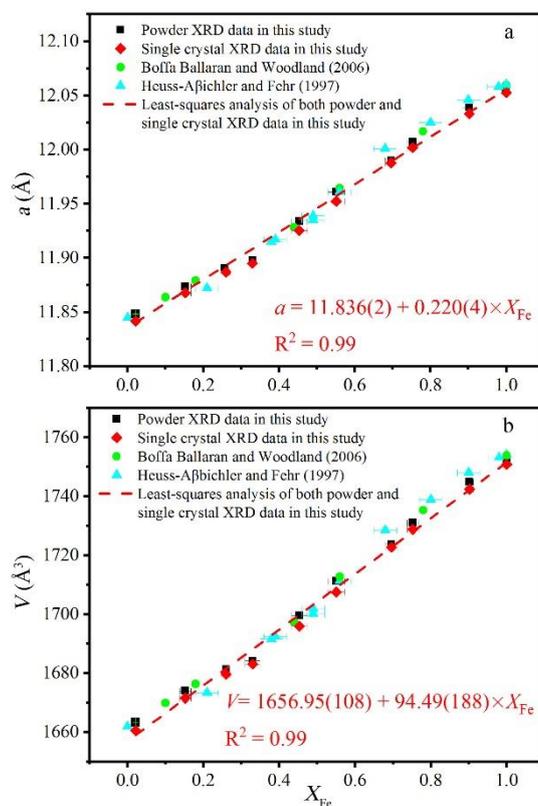
**Figure 2.** X-ray diffraction patterns of grossular–andradite solid solutions at ambient P–T condition. All the major peaks attributed to garnets are marked with hkl symbol (Miller indices). The Bragg peaks of Pt powder are denoted by solid triangles (Pt standard was not used in this study, see text). The number represents the mole fraction of andradite component ( $x = \text{Fe}^{3+}/(\text{Fe}^{3+} + \text{Al}^{3+})$ ).

The unit-cell parameters and volumes of various samples using both powder and single-crystal XRD showed little difference (Table 5). The unit-cell parameters of grossular and andradite end-member garnets were 11.848(0) and 12.054(0) Å (PXRD) and 11.842(0) and 12.053(1) Å (single-crystal XRD), respectively, which are consistent with previous studies (Table 6). The least-squares method was used to fit both powder and single-crystal XRD data in this study:  $a = 11.836(2) + 0.220(4) \times X_{\text{Fe}}$  ( $R^2 = 0.99$ ). The fitting results of four sets of data were as follows:  $a = 11.840(1) + 0.218(2) \times X_{\text{Fe}}$  ( $R^2 = 0.99$ ) (Figure 3).

**Table 5.** The unit-cell size, excess volume, and microstrain of grossular–andradite garnet solid solutions (3 GPa, 1200 °C).

Comp. $X^a$	Powder XRD					Single-Crystal XRD			
	$a$ (Å)	$V$ (Å <sup>3</sup> )	Mol $V$ (cm <sup>3</sup> /mol)	Excess $V$ (cm <sup>3</sup> /mol)	Microstr ain(10 <sup>-2</sup> )	$a$ (Å)	$V$ (Å <sup>3</sup> )	Mol $V$ (cm <sup>3</sup> /mol)	Excess $V$ (cm <sup>3</sup> /mol)
0.00(0) <sup>b</sup>	11.8484(3)	1663.31(11)	125.16(1)	0	0	11.8418(3)	1660.55(11)	124.96(1)	0
0.15(2)	11.8736(3)	1673.97(11)	125.97(1)	−0.209(1)	1.06(28)	11.8677(4)	1671.47(11)	125.78(1)	−0.213(1)
0.26(1)	11.8909(3)	1681.30(11)	126.52(1)	−0.348(1)	1.39(25)	11.8869(9)	1681.30(11)	126.52(1)	−0.308(1)
0.33(2)	11.8977(3)	1684.17(11)	126.73(1)	−0.620(1)	3.33(56)	11.8950(4)	1683.04(11)	126.65(1)	−0.549(1)
0.45(2)	11.9338(3)	1699.55(11)	128.08(1)	−0.283(1)	3.02(29)	11.9253(8)	1695.93(11)	127.62(1)	−0.417(1)
0.54(2)	11.9610(3)	1711.19(12)	129.04(1)	−0.053(1)	2.83(34)	11.9524(15)	1702.52(12)	128.49(1)	−0.206(1)
0.70(2)	11.9899(3)	1723.64(12)	129.70(1)	−0.080(1)	1.57(24)	11.9878(14)	1722.73(12)	129.64(1)	−0.047(1)
0.75(1)	12.0070(3)	1731.02(12)	130.26(1)	0.100(1)	3.18(44)	12.0018(3)	1728.78(12)	130.09(1)	0.024(1)
0.90(1)	12.0390(3)	1744.89(12)	131.30(1)	0.153(1)	0.60(38)	12.0332(3)	1742.38(12)	131.11(1)	0.034(1)
0.99(0)	12.0542(3)	1751.50(12)	131.80(1)	0	0.31(32)	12.0525(7)	1750.78(12)	131.75(1)	0

<sup>a</sup> Comp. stands for composition,  $x = \text{Fe}^{3+}/(\text{Fe}^{3+} + \text{Al}^{3+})$ ; <sup>b</sup> Number in the parentheses represents standard deviation in the rightmost digit.

**Figure 3.** Effect of composition ( $x$ ) on (a) unit-cell parameters and (b) unit-cell volume of grossular–andradite solid solution. The equations shown are fitted by least-squares analysis.

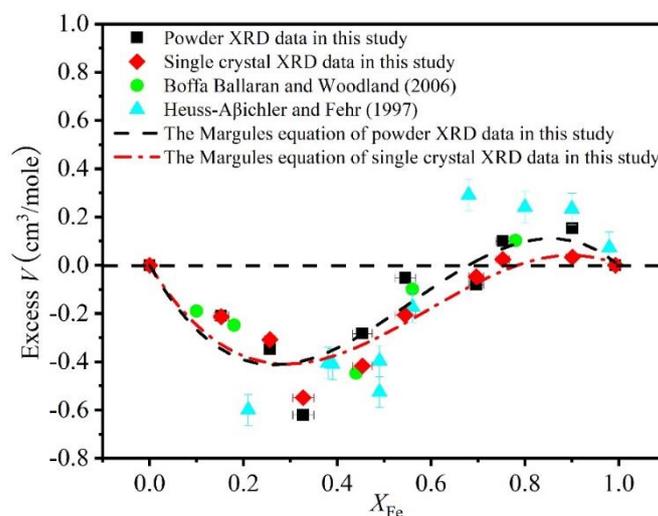
**Table 6.** Unit-cell parameter of quenched end-member grossular and andradite garnets from different studies.

Comp <sup>a</sup>	T (°C)	P (GPa)	t (h) <sup>b</sup>	a (Å)	Reference
Gro <sup>c</sup>	1200	3.0	24	11.8484(3)	This study, powder XRD
Gro	1200	3.0	28	11.8418(3)	This study, single-crystal XRD
Gro	1150	2.7	48	11.851(1)	Geiger et al., (1987) [26], powder XRD
Gro	1350–1400	4.0–4.2	48	11.8515(2)	Ganguly et al., (1993) [8], powder XRD
Gro	1400	6.0	0.5	11.850(2)	Du et al., (2016) [15], powder XRD
And <sup>d</sup>	1100	3.0	24	12.0542(3)	This study, powder XRD
And	1100	3.0	28	12.0525(7)	This study, single-crystal XRD
And	1200	2.0		12.063(1)	Armbruster and Geiger (1993) [27], single-crystal XRD
And	1100	1.2	91	12.0596(2)	Woodland and Ross (1994) [28], powder XRD

<sup>a</sup> Comp. stands for composition; <sup>b</sup> Experimental duration in hours; <sup>c</sup> Gro stands for grossular  $\text{Ca}_3\text{Al}_2\text{Si}_3\text{O}_{12}$ ; <sup>d</sup> And stands for andradite  $\text{Ca}_3\text{Fe}_2\text{Si}_3\text{O}_{12}$ .

### 3.3. Excess Volume

Excess volume is the difference in the unit-cell volume between the real mixture and the ideal solid solution, which can also be used to evaluate structural distortion (Table 5). Figure 4 shows the relationship between excess volume and composition along Gro–And solid solution. Negative excess volume value was observed in Al-rich samples in all three studies, whereas a slight positive excess volume value was observed in andradite-rich samples in this study and in Boffa Ballaran and Woodland (2006) [17]. The apparently more positive value in andraditic garnet in Heuss-Aßbichler and Fehr (1997) [19] might have been caused by inhomogeneous composition [17].



**Figure 4.** Excess volumes of mixing in grossular–andradite solid solution. A negative deviation was observed in Al-rich samples for all studies. The dashed line is fitted using single-crystal XRD data by two-parameter Margules equation.

A two-parameter Margules equation was used to model the excess volume values of garnets:

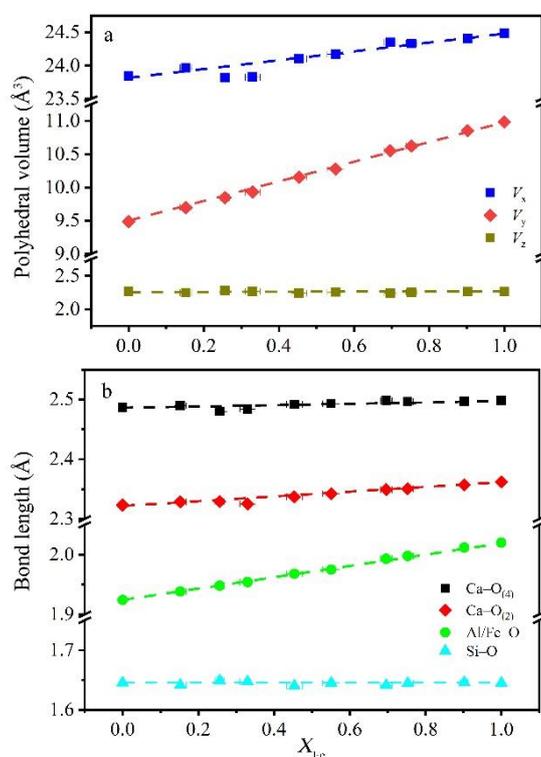
$$V(X)_{\text{excess}} = X_{\text{grossular}} \cdot X_{\text{andradite}} (X_{\text{grossular}} \cdot W_{\text{Vandradite}} + X_{\text{andradite}} \cdot W_{\text{Vgrossular}}) \quad (1)$$

where  $X$  is the mole fraction of different garnet solutions, and  $W$  is the excess volume parameter. According to fitting results,  $W_{\text{Vgrossular}} = -3.47 \pm 0.57 \text{ cm}^3/\text{mol}$  and  $W_{\text{Vandradite}} = 1.64 \pm 0.60 \text{ cm}^3/\text{mol}$  were calculated for powder XRD data, and  $W_{\text{Vgrossular}} = -3.15 \pm 0.42 \text{ cm}^3/\text{mol}$  and  $W_{\text{Vandradite}} = 0.87 \pm 0.44 \text{ cm}^3/\text{mol}$  were calculated for single-crystal XRD data. Data from the previous two studies were also modeled, giving  $W_{\text{Vgrossular}} = -2.50 \pm 0.53 \text{ cm}^3/\text{mol}$  and  $W_{\text{Vandradite}} = 1.01 \pm 0.67 \text{ cm}^3/\text{mol}$  for Boffa Ballaran and Woodland (2006) [17] and  $W_{\text{Vgrossular}} = -5.65 \pm 0.65 \text{ cm}^3/\text{mol}$  and  $W_{\text{Vandradite}} =$

$3.42 \pm 0.65 \text{ cm}^3/\text{mol}$  for Heuss-Aßbichler and Fehr (1997) [19]. The results showed that data in both studies did not fit well with Margules equation. This might have been a result of  $\text{Al}^{3+}$ – $\text{Fe}^{3+}$  cation misfit, microstrain, or different degree of  $\text{Al}^{3+}$ – $\text{Fe}^{3+}$  occupation disorder [15].

### 3.4. Bond Length and Polyhedral Volume

Bond lengths and polyhedral volumes were correlated with cation substitution (Figure 5). This is in agreement with previous studies showing cation substitutions on octahedron mainly affect the Y–O bond length, with the Z–O distance remaining constant [12,29,30]. From grossular to andradite, cation substitution occurred exclusively at the octahedral site, and the dodecahedron and tetrahedron sites were nearly occupied by Ca and Si. As shown in Figure 5a, volume change pattern turned from steep at the octahedron site through the dodecahedron site to almost flat at the tetrahedron site. Although cation substitution occurred in the octahedrons, significant negative deviation could be observed at the dodecahedron when the garnet had 30–40% And (sample with the 30–40% andradite component), which seemed to be attributed to excess volume.

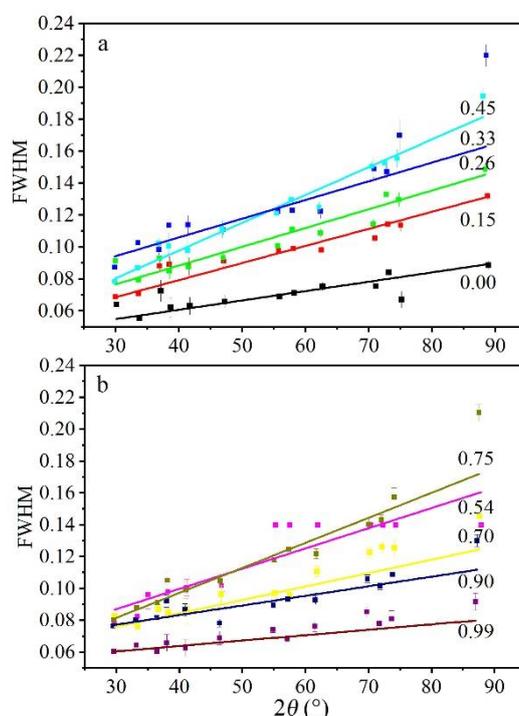


**Figure 5.** (a) Volume of polyhedra and (b) cation–oxygen distances at the X (dodecahedron), Y (octahedral), and Z (tetrahedral) sites as a function of composition.  $\text{Ca-O}_{(2)}$  and  $\text{Ca-O}_{(4)}$  are shorter and longer bond in dodecahedron, respectively. Negative deviations were observed in Al-rich samples depending on variations in dodecahedral volume and Ca–O bond length.

Because of the larger radius of  $\text{Fe}^{3+}$  compared to  $\text{Al}^{3+}$ , Al/Fe–O length linearly increased from 1.924 to 2.020 when andradite end-member content increased from 0% to 100%. Compared to octahedral bond length variation, the two inequivalent dodecahedral bond lengths—shorter  $\text{Ca-O}_{(2)}$  and longer  $\text{Ca-O}_{(4)}$ —slightly increased with andradite composition, with the former having a deeper slope. A negative deviation at 40% And was observed at both  $\text{Ca-O}_{(2)}$  and  $\text{Ca-O}_{(4)}$  bond lengths, corresponding to the volume change of dodecahedron. The Si–O bond lengths were constant within the analytical error (Figure 5).

### 3.5. XRD Peak Broadening

Peak broadening of powder XRD can be expressed by FWHM. Generally, garnets with intermediate composition showed broader FWHM than pure end-members (Figure 6), which is in accordance with previous studies [15,31]. However, the garnet with 75% andradite was an exception, exhibiting the largest FWHM.



**Figure 6.** The full width at half maximum (FWHM) values of X-ray diffraction peaks (used for unit-cell calculation) on garnets with different compositions. (a) grossular-rich samples, (b) andradite-rich samples. The number represents the mole fraction of andradite component. Garnet solid solutions with intermediate composition showed broader XRD peaks than end-members. For samples with a constant composition, XRD peak broadening increased with Bragg angle (lines with different colors only serve as labels and have no practical meaning).

For samples with a certain composition, in general, XRD peaks became broader with  $2\theta$ . The peak broadening reflects the effect of instrument and sample [15]. For samples with pure end-member composition, the FWHM was narrow and represented the instrumental effect (grossular showed little peak broadening and had a small difference from  $\text{LaB}_6$  in Du et al., (2016) [15]). As a result, if we adopt the measured peak broadening ( $B_{\text{obs}}$ ) of grossular as the instrumental broadening ( $B_{\text{inst}} = B_{\text{gro}}$ ), the real peak broadening of garnets ( $B_{\text{real}}$ ) can be determined as follows:

$$B_{\text{real}} = B_{\text{obs}} - B_{\text{inst}} \approx B_{\text{obs}} - B_{\text{gro}}. \quad (2)$$

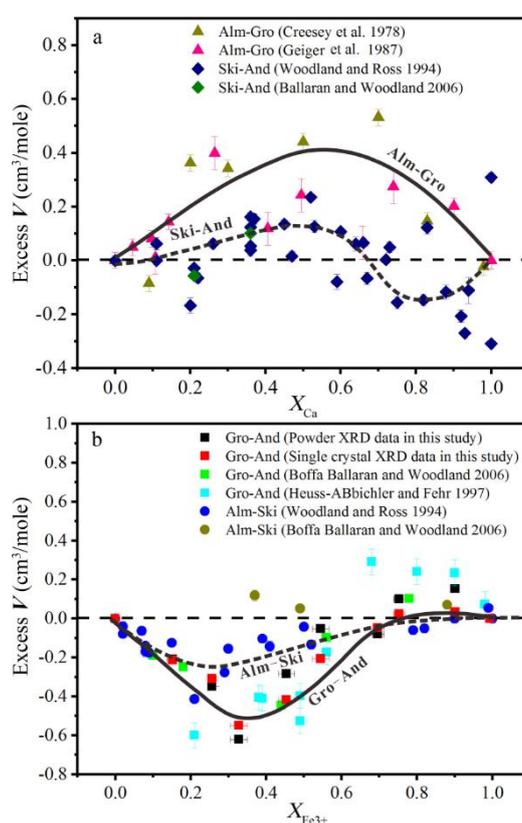
## 4. Discussion

### 4.1. Excess Volume and Distortion

When incorporating cations with different sizes in a certain site, the structure of the mineral has to adjust to larger or smaller unit-cell volume. As a result, excess volume can represent the size of the lattice distortion.

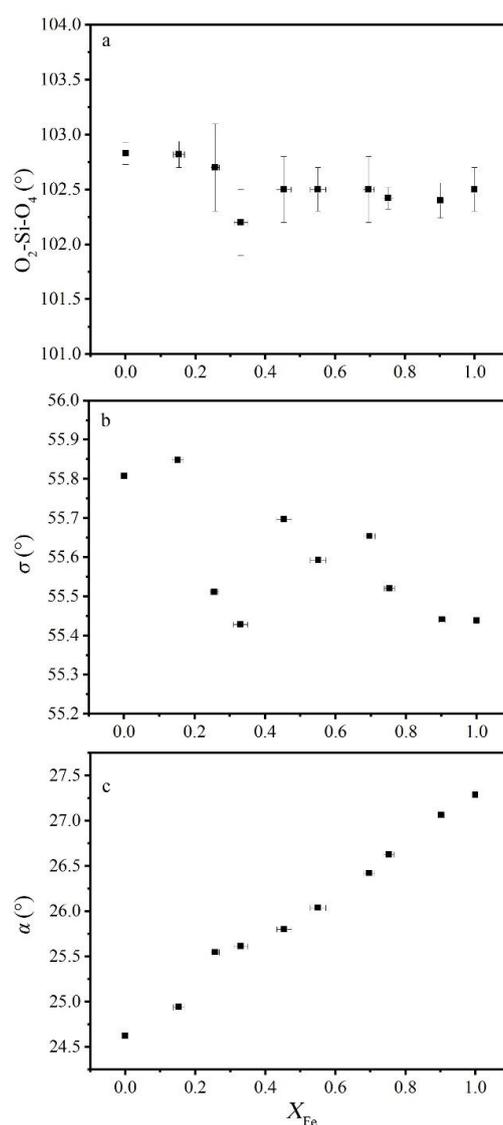
After comparing the excess volume in Pyr–Gro, Alm–Gro, and Pyr–Alm garnet solid solutions, Ungaretti et al., (1995) [13] proved that the radius difference of cations in the dodecahedron site will

affect the distortion size. As shown in Figure 7a, at fixed octahedral cation, Alm–Gro garnet had higher excess volume than Ski (skiagite,  $\text{Fe}_3\text{Fe}_2\text{Si}_3\text{O}_{12}$ )–And garnet (data from [17,19,26,28,32]). Excess volume along the solid solution with Mg– $\text{Fe}^{2+}$  exchanging on dodecahedron was also affected by the type of cation on the octahedron. Due to large radius difference between  $\text{Al}^{3+}$  and  $\text{Fe}^{3+}$  ( $\Delta r = 0.11 \text{ \AA}$  [33]), excess volume also appeared for Gro–And and Alm–Gro solid solutions (Figure 7b). The results showed that an asymmetric negative excess volume toward the Al-bearing end-member was apparent in both solid solutions (Figure 7b).



**Figure 7.** Excess volumes of mixing in (a) almandine–grossular and skiagite–andradite solid solutions and (b) grossular–andradite and almandine–skiagite solid solutions.

The excess volume along the Gro–And join reached the minimum value ( $-0.6 \text{ cm}^3/\text{mol}$ ) at nearly 33% And, which might indicate maximum structural deformation. The volume of tetrahedron was almost constant in the calculated X, Y, and Z polyhedral volume (Figure 5a). This could also be proven by the fixed value of the Si–O bond length (Figure 5b) and the  $\text{O}_2\text{–Si–O}_4$  angle (Figure 8a). When Al was completely substituted by Fe, the octahedron volume and Y–O length in Gro–And series decreased from  $\sim 15.79\%$  to  $\sim 4.99\%$  (Figure 5a), whereas the bond length in Alm–Ski series increased about 4.74% [28]. An approximately linear variation could be observed in the octahedron volume and the Y–O bond length, indicating that the octahedron had no obvious distortion (Figure 5). The main distortion seemed to occur in the dodecahedron, even though the volume only increased by 2.67%. Corresponding to excess volume, negative deviations of both polyhedral volume and X–O bond length were observed in Al-bearing samples. Similar to the Alm–Ski solid solution in Woodland and Ross (1994) [28], two inequivalent dodecahedral Ca–O bond had 0.48% and 1.68% growth rate for long and short bonds, respectively. This could be one of the reasons for the dodecahedral distortion.



**Figure 8.** (a) Angle variance of  $O_2-Si-O_4$  in tetrahedron, (b) the octahedral distortion parameter  $\sigma$  (Euler and Bruce (1965) [34]), and (c) the tetrahedral rotation parameter  $\alpha$  (Born and Zemann (1964) [34]) as a function of composition for grossular-andradite solid solution.

Distortion of octahedron can be described by index  $\sigma$  (Figure 8b), which is the angle between the 3-axis and the Y–O bond [34]. When  $\sigma$  is equal to  $54.75^\circ$ , octahedron has the standard form. In Gro–And solid solution,  $\sigma$  slightly changed and became larger than  $54.736^\circ$ , which meant the octahedral site of garnet always had a flat shape and little structural distortion. Unexpectedly, Woodland and Ross (1994) [28] showed that  $\sigma$  changed from  $53.34^\circ$  to  $55.42^\circ$  on Ski–And solid solution, indicating greater influence of dodecahedral cation substitution. As the octahedron shared edges with neighboring dodecahedral sites, distortion was greater when cation substitutions occurred on neighboring dodecahedral sites than on the octahedral site itself. The opposite also seemed to be true for Gro–And solid solution, with the dodecahedral sites having more obvious distortion compared with octahedron in this study (Figure 5a).

For tetrahedron, two factors should be considered: (1) expansion or contraction along the 4-axis and (2) rotation around the 4-axis. Variation of  $O_2-Si-O_4$  angle (Figure 8a) can be used to represent the degree of expansion or contraction distortion. Limited angle variation indicates this kind of distortion is low and can be ignored. Degree of rotation can be represented by the rotation of tetrahedron  $\alpha$  [35], which is defined as the angle between the crystallographic a-axis and the polyhedral edge

shared between the tetrahedra and dodecahedra. As shown in Figure 8c,  $\alpha$  increased with increasing andradite content and had a negative deviation in Al-rich samples. The change in  $\alpha$  might have been caused by Al–Fe substitution in the octahedral sites because the tetrahedra shared each corner with the octahedra. With the increase in  $\text{Fe}^{3+}$  content, the linked tetrahedra rotated to higher angles of  $\alpha$  in an attempt to accommodate the larger octahedral cation into the network. In addition, the tetrahedra also shared edges with parts of the dodecahedra. Therefore,  $\alpha$  has important structural implications on dodecahedra, such as inequivalent Ca–O<sub>(2)</sub>/Ca–O<sub>(4)</sub> bond increase [6,27,35–37], which can also cause distortions in the dodecahedral sites.

As structural distortion mainly occurred in the dodecahedral site for the Gro–And solid solution, the cation species substituted in dodecahedral site would affect the magnitude of the distortion (Figure 7b). Compared to  $\text{Fe}^{2+}$  with a smaller radius, dodecahedral site occupied by  $\text{Ca}^{2+}$  with a larger polyhedral volume and a longer X–O bond length seemed to be more easily distorted structurally.

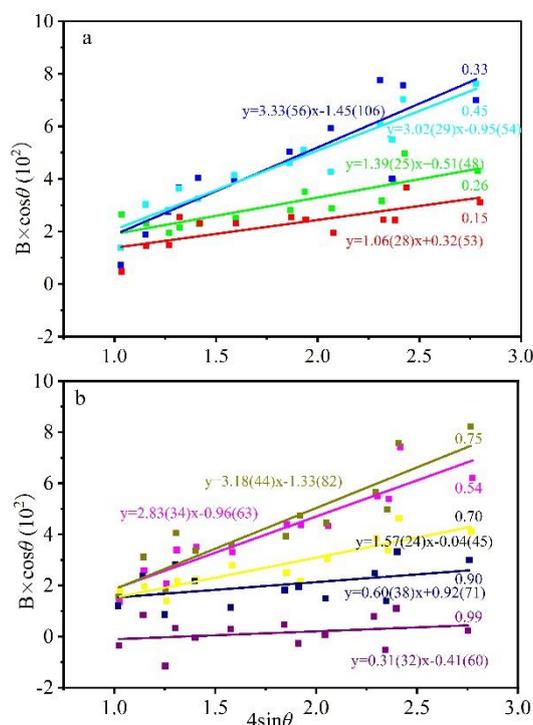
#### 4.2. Microstrain

Williamson–Hall plot [38] establishes the relationship between peak broadening and microstrain and grain size:

$$B_{\text{garnet}} = B_{\text{size+strain}} = \frac{K\lambda}{\langle L \rangle \cos\theta} + 4\eta \tan\theta \quad (3)$$

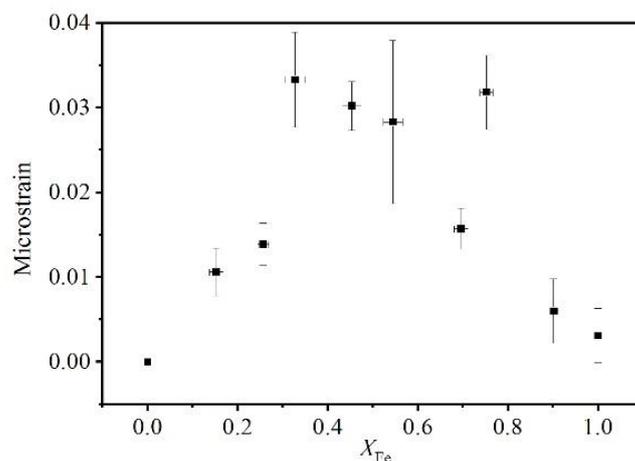
where  $K = 0.94$ ,  $L$  is related to crystal size,  $\eta$  is the microstrain factor,  $\lambda$  is the wavelength of the X-ray, and  $\theta$  is the Bragg angle. The formula can be transformed into (Figure 9):

$$B \times \cos\theta = \frac{K\lambda}{\langle L \rangle} + 4\eta \sin\theta \quad (4)$$



**Figure 9.** Williamson–Hall plot for XRD peak widening data of grossular–andradite solid solution. The number represents the mole fraction of andradite component. The slope of equation represents the microstrain factor ( $\eta$ ). (a) grossular-rich samples; (b) andradite-rich samples.

When the peak broadening was completely caused by the instrument, i.e., when the composition was purely grossular, the garnet showed little microstrain (Figure 10). The microstrain showed two maxima when the garnet had 33% and 75% andradite, respectively, and the value between the peaks was slightly lower. The double peak shape pattern was considered to be related to nonuniform lattice distortion (Figure 10), which was similar to the microstrain relationship for Pyr–Gro solid solution [15].



**Figure 10.** Microstrain inside garnet structure calculated from X-ray peak broadening along the pyrope–grossular join.

The size of the microstrain is related to defects [31]. Dapiaggi et al., (2005) [31] proposed that the microstrain in pyrope–grossular garnets may be related to excess enthalpies because elastic strain has an important contribution to mixing enthalpies [39]. However, Becker and Pollok (2002) [40] predicted nearly symmetric positive excess behavior of enthalpies in Gro–And solid solution. Due to complex contributions, such as elastic, electronic, and magnetic effects to thermodynamic properties, further research is necessary.

## 5. Conclusions

The analysis of XRD diffraction data along the grossular–andradite solid solution series showed that Al-rich samples had small negative excess volume. The dodecahedral volume and Ca–O bond length had a negative deviation at about 30–40% And, which seemed to be responsible for excess volume in the solid solution. The tight edge connection between the octahedral and dodecahedral sites as well as the nonlinear increase in the tetrahedral rotation angle  $\alpha$  might be the reasons for dodecahedral distortion. Contrary to symmetric positive trend in excess enthalpies, microstrain calculated from powder X-ray diffraction data showed double peaks, which needs further investigation.

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**Conflicts of Interest:** The authors declare no conflict of interest.

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