

Article

# Underground Vertical Seismic Profiling with Conventional and Fiber-Optic Systems for Exploration in the Kylylahti Polymetallic Mine, Eastern Finland

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**Abstract:** Seismic reflection methods have been used for the exploration of mineral resources for several decades. However, despite their unmatched spatial resolution and depth penetration, they only have played a minor role in mineral discoveries so far. Instead, mining and exploration companies have traditionally focused more on the use of potential field, electric and electromagnetic methods. In this context, we present a case study of an underground Vertical Seismic Profiling (VSP) experiment, which was designed to image a (semi-)massive sulfide deposit located in the Kylylahti polymetallic mine in eastern Finland. For the measurement, we used a conventional VSP with three-component geophones and a novel fiber-optic Distributed Acoustic Sensing (DAS) system. Both systems were deployed in boreholes located nearby the target sulfide deposit, and used in combination with an active seismic source that was fired from within the underground tunnels. With this setup, we successfully recorded seismic reflections from the deposit and its nearby geological contrasts. The recording systems provided data with a good signal-to-noise ratio and high spatial resolution. In addition to the measurements, we generated a realistic synthetic dataset based on a detailed geological model derived from extensive drilling data and petrophysical laboratory analysis. Specific processing and imaging of the acquired and synthetic datasets yielded high-resolution reflectivity images. Joint analysis of these images and cross-validation with lithological logging data from 135 nearby boreholes led to successful interpretation of key geological contacts including the target sulfide mineralization. In conclusion, our experiment demonstrates the value of in-mine VSP measurements for detailed resource delineation in a complex geological setting. In particular, we emphasize the potential benefit of using fiber-optic DAS systems, which provide reflection data at sufficient quality with less logistical effort and a higher acquisition rate. This amounts to a lower total acquisition cost, which makes DAS a valuable tool for future mineral exploration activities.

**Keywords:** mineral exploration; reflection seismic imaging; hardrock seismics; vertical seismic profiling; distributed acoustic sensing

## 1. Introduction

Seismic reflection methods have been used for mineral exploration and mine planning for several decades (e.g., [1–4] and the references therein). Traditionally, mineral prospecting has mainly concentrated on the use of electromagnetic and potential field methods. However, the growing demand for mineral resources in modern day society continuously fuels the search for new deposits. This drives mineral exploration into increasingly deeper levels, which exceed the coverage of the traditional techniques [5,6]. Reflection seismics, on the other hand, offers depth penetration and high spatial resolution down to several kilometers of depth. Although the application of seismic methods to mineral exploration poses various challenges, several successful cases, including surveys from Australia [7], Canada [8–11], Scandinavia [12–20] and South Africa [21–23], have demonstrated the feasibility of the method.

Despite these successes, some challenges still remain, which have restrained the use of seismic methods on a wider scale for industrial applications. One important reason is that mineral deposits are often located in hardrock rather than sedimentary environments. Hardrock settings are usually dominated by igneous and metamorphic rocks that have undergone severe deformation due to tectonic processes [1]. This causes various challenges for seismic methods: (1) The predominant rock types usually exhibit high seismic velocities ( $\approx 4\text{--}8$  km/s [24]), causing the spatial resolution to be lower compared to sedimentary environments. (2) The rocks often show less variation in terms of their elastic properties [6], which results in weaker reflective signals arising from the geological interfaces (i.e., low Signal-to-Noise Ratio (SNR)). (3) The tectonic activities produce irregularly-shaped geological interfaces, which cause the seismic wavefield to be dominated by similar proportions of diffractions and reflections [5,25,26]. (4) The tectonic processes often result in steeply-dipping geological units, which are difficult to image with surface-based acquisitions.

As a result of these factors, seismic sections obtained in hardrock mining and mineral exploration settings often lack well-pronounced, continuous reflectors, but are characterized by piecewise, disconnected reflectivity patterns, which can be difficult to interpret.

In order to tackle the above-mentioned challenges and to make seismic reflection methods more attractive for mineral exploration, various attempts have been made to increase the overall efficiency and to lower the costs of seismic surveys. One way is to integrate seismic reflection methods with other geophysical and geological techniques that are conventionally used in mineral exploration (e.g., [22,27]). Another promising approach is to combine strategically different forms of seismic acquisition. 2D surveys can be carried out for exploration at a regional-scale, followed a combination of a smaller-scale 3D survey and borehole seismic measurements for more focused exploration and precise resource delineation [6,16,28]. For resource delineation, borehole seismic surveys have been proven to be particularly useful, implementing either crosshole tomography approaches [29–31] or reflectivity imaging [32–37] using Vertical Seismic Profiling (VSP) methods. Additionally, the application of new technologies, such as passive seismic imaging methods [38] or the use of fiber-optic technologies [39], provides new possibilities for seismic surveying.

The presented work is part of a larger international research collaboration, which aims at the development of COst-effective Geophysical Imaging Techniques for supporting Ongoing MINeral exploration in Europe (COGITO-MIN). The COGITO-MIN project involved acquisition and integration of multi-scale seismic methods including surface-based 2D and 3D active-source and passive seismic measurements and an underground in-mine VSP survey using conventional and fiber-optic technologies. This work addresses the VSP measurements that were carried out inside the Kylylahti polymetallic mine.

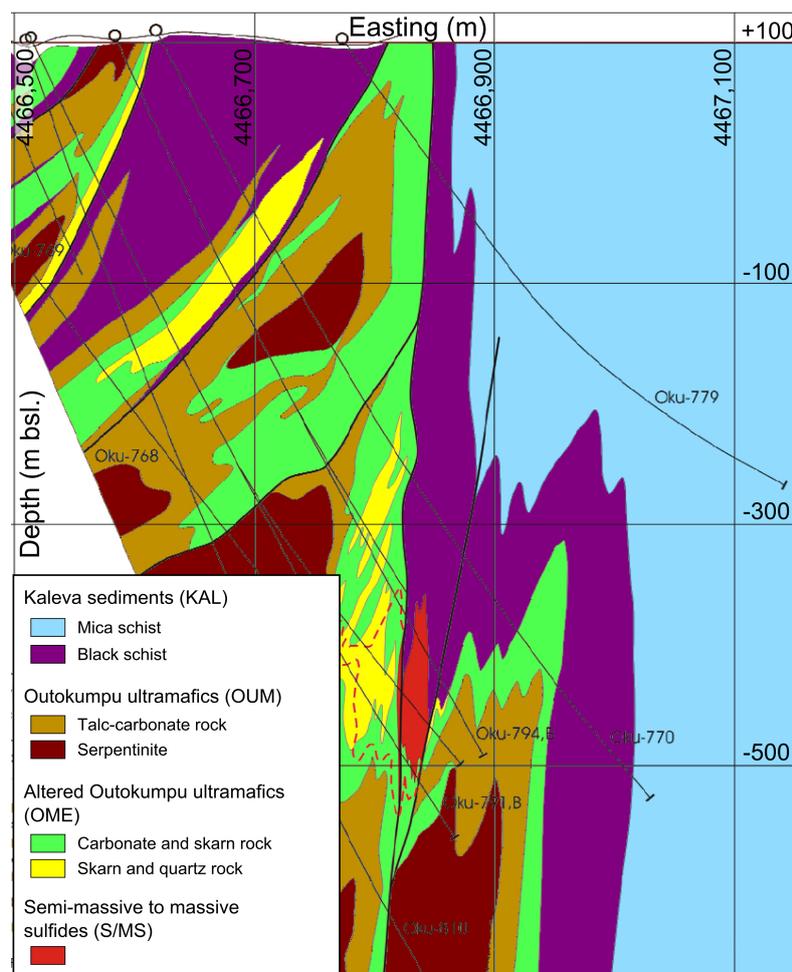
For this purpose, we first explain the geological setting of the survey area and its petrophysical characteristics. Next, we describe the experimental setup of the VSP survey inside the Kylylahti mine and compare the used VSP systems in terms of their technological characteristics and data quality. To facilitate interpretation of the acquired VSP data, we then present a realistic seismic forward modeling study, which helps to identify target reflections and link them to the predominant geological

contrasts. Finally, we present reflectivity images of the acquired and synthetic VSP datasets and discuss their geological interpretation in relation to borehole data.

## 2. Geological and Petrophysical Setting

### 2.1. Geological Setting

The Kylylahti mine is located in the Outokumpu mining district, where the first copper mine was opened in 1913. The predominant geology of this area is described in detail by [40]. In the Kylylahti area, the geology can be characterized in terms of four units (see Figure 1): (1) The Kylylahti deposit, a Semi-massive to massive sulfide (S/MS) mineralization, is associated with (2) ultramafic rocks Outokumpu ultramafics (OUM) that mainly consist of serpentinite and talc-carbonate rocks. These are fringed by (3) alteration zones of a few to several tens of meters thickness composed of carbonate-skarn-quartz rocks, the Altered Outokumpu ultramafics (OME). In combination, the nearly N-S trending and near-vertical OUM and OME units are called the Outokumpu assemblage rocks, which are embedded in (4) the regional Kaleva Sedimentary Belt (KAL). The Kaleva sediments consist of mica schist and black schist. They form a near-vertical contrast with the Outokumpu assemblage rocks (OUM + OME). Moreover, black schist interlayers also occur in the severely folded Outokumpu assemblage package, where various rock contacts are repeated several times.



**Figure 1.** Vertical cross-section from west to east showing the predominant geological units of the Kylylahti mine area (modified after [40]). The red dashed line is the 1000 ppm Cu isopleth that constrains the extent of the ore-grade Co-Cu sulfide disseminations flanking the massive-semimassive sulfides.

The Kylylahti S/MS deposit is located in a steeply-dipping contact between black schists and carbonate-skarn-quartz rocks (Figure 1). It has the shape of a  $\approx 1.3$  km-long, N-S trending lens, which reaches into a depth of  $\approx 750$  m below sea level. The maximum thickness and height are 30 m and 170 m, respectively. The shape of the S/MS deposit in combination with the severely folded and near-vertical geological contrasts makes this deposit a very challenging imaging target for surface-based seismic surveys. For this reason, the in-mine VSP measurements provide the most promising tool for accurate delineation of the deposit.

## 2.2. Petrophysics

Whether or not a mineral deposit is detectable using the seismic reflection method depends on the distribution of the elastic properties throughout the predominant geological units. These properties are the density  $\rho$  and the seismic P- and S-wave velocities  $v_P$  and  $v_S$ , respectively. Those form the seismic impedances:

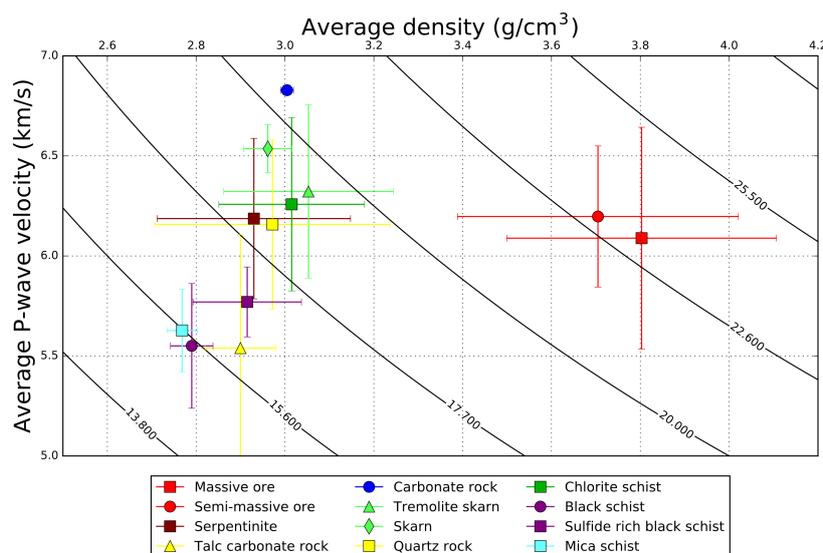
$$Z_{P,S} = \rho \cdot v_{P,S} \quad (1)$$

which are related to the seismic reflectivity:

$$R = \frac{Z_2 - Z_1}{Z_2 + Z_1}. \quad (2)$$

Equation (2) describes the seismic reflection coefficient that is caused by the impedance contrast of two subsequent geological layers 1 and 2. The coefficient can be computed for either P- or S-waves. It is directly proportional to the amount of seismic energy that is reflected back from an impedance contrast (the reflector) under normal incidence. The more general cases of arbitrary incidence are mathematically described by the Zoeppritz equations [41,42].

The reflective contrast between many mineral deposits and their host-rocks has long been recognized [1,43–45] and has led to successful application of the seismic reflection method for mineral exploration, in particular for the exploration of sulfide deposits [24,46–48]. For the Kylylahti mine area, detailed petrophysical laboratory measurements of about 200 rock samples revealed a strong impedance contrast between the S/MS mineralization and its various host rock units [49] (Figure 2).



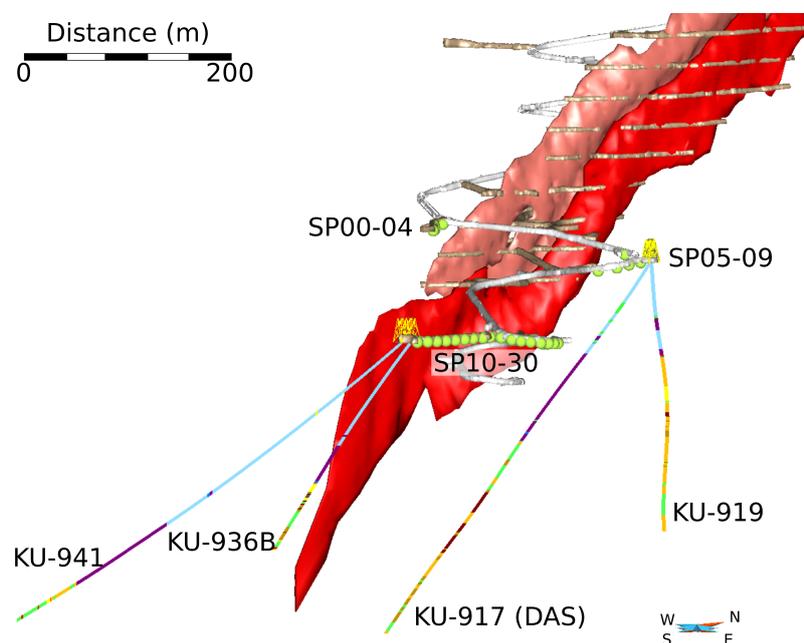
**Figure 2.** Results of petrophysical laboratory analysis of rock samples from the Kylylahti mine area (modified after [49]). Error bars indicate respective standard deviations. Contour lines denote the P-wave (acoustic) impedance. The line spacing corresponds to a reflection coefficient of 0.06, which is necessary to produce a detectable reflection [24].

In addition, the petrophysical analysis reveals that various weaker reflections can be expected from within the Outokumpu assemblage unit. In particular, the talc-carbonate rocks exhibit a significantly lower P-wave velocity than the serpentinites. Both rock types belong to the OUM and form a clear and distinctive contrast evidenced by various borehole logs. On the other hand, the OME rocks are much more difficult to distinguish lithologically, which complicates precise logging and the geological interpretation of the obtained VSP images (see Section 6).

### 3. Vertical Seismic Profiling at the Kylylahti Mine Site

#### 3.1. In-Mine VSP Acquisition

The Kylylahti survey campaign was launched in early fall 2016. As part of this survey, the underground VSP measurements were carried out inside the Kylylahti polymetallic mine (Figure 3). The imminent goal of the VSP survey was to image a massive-to-semi-massive sulfide mineralization.



**Figure 3.** Geological model of the known Kylylahti sulfide deposit (red/pink). Boreholes used for VSP measurements (KU-917, KU-919, KU-936B, KU-941) are color-coded according to logged geology (see Figure 1). Seismic sources are indicated by green spheres in the mine tunnels.

In comparison to surface-based seismic acquisition, the underground VSP measurements have several advantages in this setting. First, their sources and receivers are located within the bedrock and in direct vicinity of the target mineralization. This means that the deteriorating effect of the near-surface is avoided by these types of measurements. Second, they operate at much higher dominant frequency ( $\approx 180$  Hz), which provides a significantly higher spatial resolution than the surface-based measurements. Third, their acquisition geometry is more appropriate to image the sulfide deposit, which embodies a nearly vertical lens of only a few tens of meters thickness [40]. The VSP measurements illuminate the deposit from the side, and therefore, they are much more likely to record a pronounced reflection. For these reasons, the underground VSP measurements represent the most promising tool for detailed imaging of the sulfide deposit amongst the various seismic methods carried out in the Kylylahti survey.

The acquisition configuration of the VSP experiment is shown in Figure 3 and summarized in Table 1. In total, four boreholes from prior exploration drilling were used, which are located adjacent to and inclined towards the target mineralization (Figure 3). The VSP recording systems were deployed in these boreholes. A VIBSIST-200 system was used as a seismic source, which represents a multi-impact, time-distributed source [50,51]. The advantages of this system consist of its mobility, the high frequency of the source signal (180 Hz) and the high signal-to-noise ratio that is achieved by stacking of multiple impacts. The source was fired from 31 distinct shot locations inside the mine tunnels. The locations can be grouped into three different depth levels: the first five shots (SP00–04) were located in a tunnel above the mineralization at  $\approx 550$  m depth; shots SP05–09 were positioned at  $\approx 580$  m along the main decline; and shots SP10–30 were located in a tunnel at 635 m depth below mean sea level. Data recorded from the first group of shots were found to have almost no reflected energy, which is most likely due to the acquisition geometry. Thus, these shots were excluded from the detailed analysis and interpretation later on.

**Table 1.** Borehole receiver configurations for the Kylylahti Vertical Seismic Profiling (VSP) survey. The receiver range is given in downhole depth. KU-917 (DAS) refers to the the Distributed Acoustic Sensing (DAS) fiber-optic system. All other configurations used conventional three-component geophones.

Borehole	Starting Depth (m bsl.)	Total Borehole Length	Receiver Range	Receiver Spacing
KU-917	580.6 m	523 m	46.3 m to 365.6 m	2.5 m
KU-919 <sup>1</sup>	580.4 m	544 m	72.4 m to 142.4 m	10 m
KU-936B	632.6 m	396 m	2.4 m to 392.4 m	5 m
KU-941	632.2 m	544 m	46.3 m to 537.2 m	5 m
KU-917 (DAS)	580.6 m	523 m	0.0 m to 451.38 m	1 m

<sup>1</sup> Data from KU-919 were only acquired for data quality control and not used for interpretation.

All the boreholes were equipped with a conventional VSP system using three-component (3C) geophones as seismic receivers. Varying geophone spacings along the cable were applied (see Table 1). Depending on the spacing, the length of the entire recording system varied between 160 m and 240 m. Hence, in order to survey one complete borehole length, the geophone string had to be lowered successively, followed by a repetition of all the associated shots for each new receiver configuration.

### 3.2. Distributed Acoustic Sensing

In addition to the conventional VSP acquisition, a fiber-optic Distributed Acoustic Sensing (DAS) measurement was also carried out in borehole KU-917. DAS represents a new technology based on Optical Time-Domain Reflectometry (OTDR) [39,52,53]. In contrast to conventional VSP systems, DAS uses one single fiber-optic cable as a continuous sensor, which is connected to a data interrogator unit. The interrogator emits a laser pulse down the cable, which is back-scattered by small material fluctuations within the fiber. If the cable is subjected to elastic strain, this results in a phase difference in the backscattering response, which is measured over a spatial interval along the cable, the gauge length. The phase difference taken over the gauge length is the optical measuring quantity. It is proportional to the subjected strain rate, which is internally computed and output by the interrogator.

The Kylylahti DAS experiment was carried out using the commercial iDAS<sup>TM</sup> system provided by Silixa Ltd. (Elstree, UK) in combination with a single straight fiber. Due to operational reasons, the cable could not be cemented in the borehole, but rested against the borehole wall. The borehole was filled with water. The borehole inclination ranged between 37 and 42 degrees from the vertical.

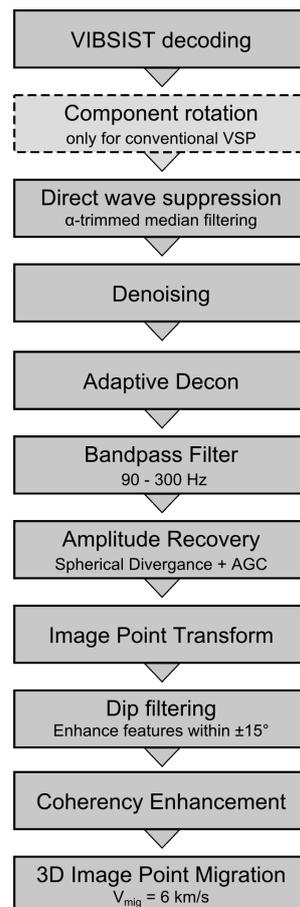
In practice, the sampling resolution in the fiber-optic cable is determined by the speed of light throughout the fiber. In our case, the sampling interval was set to 1 m. However, as mentioned above, the system's output signal represents an average over a spatial window, the gauge length, which depends on the width of the laser pulse. In our study, the gauge length was set to a default value of 10 m, which represents the effective spatial resolution of the system. In principle, this value

can be decreased; however, the choice of the optimal gauge length is a trade-off between resolution and signal-to-noise ratio (e.g., [54]).

DAS allows simultaneous measurement over one cable ranging up to several tens of kilometers. For VSP applications, this implies that no shot repetition is required to survey one borehole. Thus, the reduced logistical effort and high acquisition speed make DAS a promising low-cost alternative for conventional geophone-based systems. So far, DAS has been used in various geoscientific applications, which primarily include monitoring of carbon-sequestration storage sites (e.g., [55–57]), hydro-carbon reservoir monitoring (e.g., [58–60]) and, recently, noise recording for near-surface traffic monitoring [61] and in-mine safety applications [62]. However, to our knowledge, no applications for hardrock mineral exploration had been conducted before this experiment.

#### 4. Processing and Comparison of DAS versus 3C VSP Data

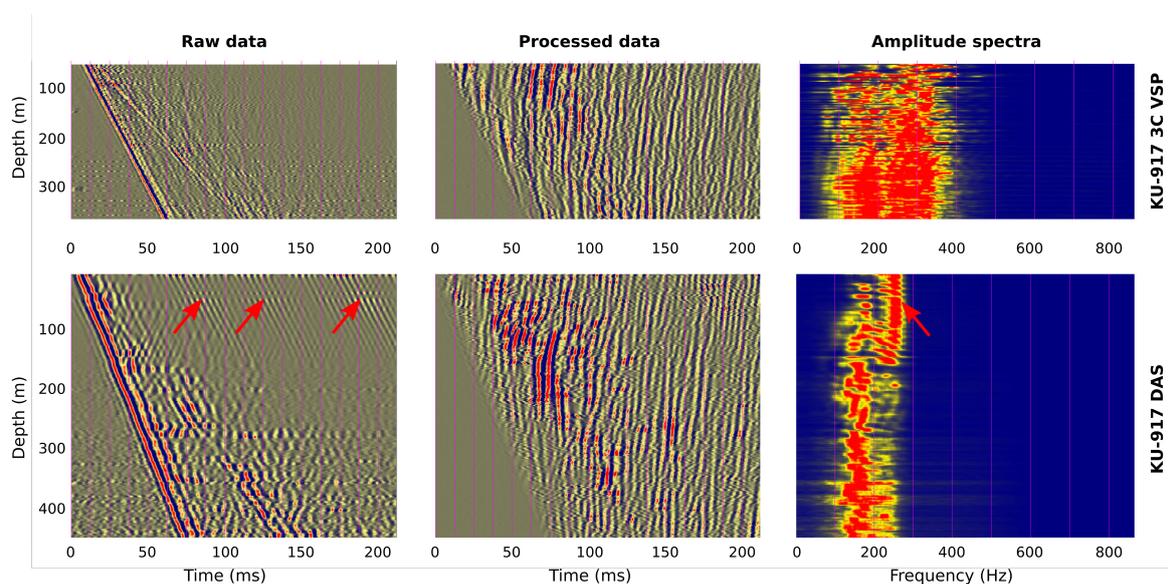
After successful acquisition, geophone VSP and DAS VSP data were processed using a VSP-specific processing flow (Figure 4). Most elements in the flow were commonly applied to both datasets, except for the component rotation, which was only done for the geophone VSP data. For this, A 3C rotation was used, based on polarity analysis of the first arrival maximum energy direction of incidence. Commonly-applied processing steps include suppression of direct P- and S-wave arrivals, denoising, the time-varying deconvolution method based on adaptive linear filtering techniques, bandpass filtering and amplitude recovery. The specific parameterizations of these steps varied slightly between the DAS and VSP records due to the different nature of the data.



**Figure 4.** Processing flow applied to geophone VSP and DAS VSP data. Component rotation was only applied to geophone VSP data. Dip filtering and coherency enhancement were carried out in the Image Point (IP) space [63].

The Image Point (IP) transform [63] represents a crucial processing step, which allows separation and extraction of target wave energy (e.g., P-wave reflections) in the IP space. Moreover, it provides the basis for further processing steps such as dip filtering and coherency enhancement, which are also carried out in the IP space. Dip filtering was used to emphasize features within a dip range of  $\pm 15$  deg from the vertical, which is based on the a priori knowledge of the orientation of the geological contrasts in the area. The processing flow is concluded with 3D IP migration, which provides reflectivity images in 3D space that are used for geological interpretation.

Figure 5 shows a comparison between the geophone VSP and DAS VSP data for shot point SP06 (zero-offset shot; see Figure 3). It compares the respective shot records before and after processing, as well as the raw amplitude spectra. It has to be noted that the obtained records do not show the same physical quantity, i.e., they show particle velocity for 3C geophones versus strain rate for DAS. Hence, the comparison is mainly restricted to the kinematics of the data. The amplitudes can be compared on a qualitative, but not on a quantitative level.



**Figure 5.** Comparison of conventional 3C geophones (top row) against fiber-optic DAS data (bottom row). The left column shows the raw data. The center column shows the processed data (before migration). The right column shows the amplitude spectra after processing was applied. The Z component (along the borehole) is shown for the 3C raw data and the P-component (particle movement towards the source location) for the processed 3C data. Strain rate is shown for the DAS data. Red arrows indicate noise in the upper part of the fiber-optic cable due to oscillations in the uncemented borehole.

Comparing the DAS versus geophone records reveals that both datasets share many similarities, but also exhibit some important differences. The raw data show a clear direct P-wave, a slightly less pronounced direct S-wave and some lower-amplitude indications of backscattered energy for both records. In that, the geophone data appear to have a higher signal-to-noise ratio and a higher spatial resolution, which results in an overall crisper appearance of the seismic events. In addition, the DAS record is subject to a strong high-frequency noise signal that is persistent in the top 150 m along the cable. This signal is caused by a lack of cementation of the borehole, which resulted in the fiber resting loosely against the borehole wall and thus producing the observed oscillations. The lack of borehole cementation also may account for the lower signal-to-noise ratio of the DAS due to the inferior coupling.

VSP-specific processing of the raw shot records enhanced the overall signal quality, particularly P-wave reflections, which represent the target signals. In addition, unwanted energy such as direct waves, reverberations and various noise are suppressed. At this point, we observed a slight difference ( $\approx 5\%$ ) between the average velocities measured from the first arrivals of the geophone and the DAS VSP records. We believe that this is due to uncertainties in the positioning of the virtual DAS receivers. Applying the same processing sequence to geophone VSP and DAS VSP shots significantly narrows the gap between the two datasets in terms of overall quality. Still, the reflections appear slightly more continuous in the geophone VSP record, particularly in the top part of the borehole.

A third comparison is made using the amplitude spectra from both datasets. This confirms the superior bandwidth of the geophone VSP data compared to the DAS records, which particularly affects the high frequency range. In addition, the cable oscillations observed along the top 150 m of the borehole in the shot record also manifest as a clear artifact in the spectrum.

The observation of shorter bandwidth in the DAS might be slightly unexpected and thus deserves more detailed consideration. The geophone VSP data were acquired with a very small geophone spacing of 2.5 m in borehole KU-917, which provides shot records of high resolution. In the case of DAS VSP, the gauge length was set to a default value of 10 m, which represents the effective spatial resolution of the system. In the presented study, the medium velocity along the borehole equals about 6 km/s, as obtained from the first arrivals. From these values, we obtain a maximum detectable frequency of about 300 Hz for the fiber optic system in our case. This corresponds well to the maximum recorded frequency that is observed in the raw amplitude spectra (Figure 5).

## 5. Seismic Forward Modeling

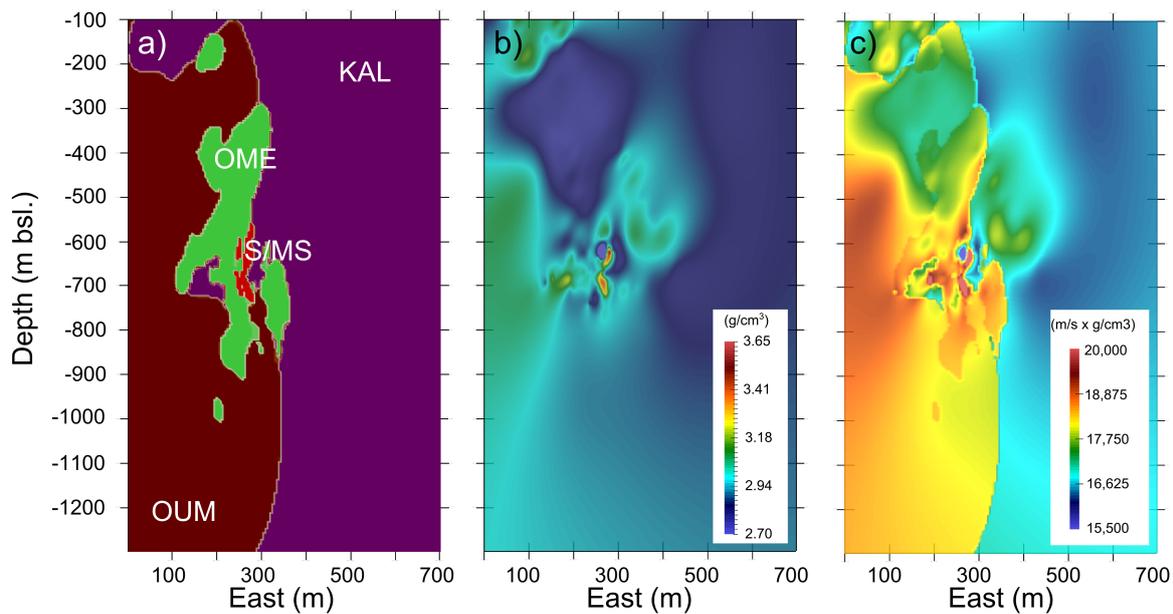
Even though seismic processing significantly enhanced the quality of the acquired VSP shots, the understanding and interpretation of the recorded reflections is not a straightforward task. This is mainly due to the geometrical ambiguity inherent to VSP surveys in general, and the complexity of the geological setting at Kylylahti mine site specifically.

For this reason, we conducted a seismic forward modeling experiment to study the wave propagation in the area of interest and to identify key reflections with their geological causes (i.e., rock contrasts). Seismic forward modeling has been utilized to better understand seismic data in hardrock environments in various studies (e.g., [13,26,64,65]). However, in order to conduct a realistic modeling study, an accurate input model is required.

### 5.1. Geological and Petrophysical Models

For our study, we used a detailed geological model as a starting point (Figure 6a). This model was derived from about 1250 boreholes, out of which about 900 are located in the immediate vicinity of the sulfide deposit, and thus provides very detailed information about the predominant lithology. The model divides the geology into the larger dominant geological units that were discussed above: the regional sedimentary rocks (KAL), the Outokumpu assemblage rocks (OME and OUM) and the semi-massive to massive sulfide mineralization (S/MS). For these units, a P-wave velocity model was derived using the average P-wave velocities of the respective dominant rock type for each unit, as determined by the petrophysical laboratory measurements [49]. The resulting velocity model does not distinguish between the individual rock types, but only between the larger geological units.

In addition to geological logging, specific gravity measurements were routinely made on borehole core samples. Those provide a detailed 3D density model (Figure 6b) that was used directly in the seismic forward modeling. Both quantities, P-wave velocity and density, form the acoustic impedance (Figure 6c), which represents the physical cause for seismic reflections. The petrophysical parameters of the modeling domain are summarized in Table 2.



**Figure 6.** Vertical sections through geological and petrophysical models used for seismic forward modeling study. (a) Model of the larger geological domains that were used to define the P-wave velocities. (b) Kriging interpolated density model obtained from specific gravity measurements of borehole core logs. (c) Acoustic impedance model calculated from the P-wave velocity model (a) and density model (b).

**Table 2.** Overview of geological and petrophysical parameters used for seismic forward modeling. Density and acoustic impedance values represent the mean average values within the respective geological units. More detailed explanations are given in the text.

Geological Unit	Dominant Rock Type	P-Wave Velocity	Density	Acoustic Impedance
KAL	Black schist	5800 m/s	2.893 g/cm <sup>3</sup>	16,781 m/s × g/cm <sup>3</sup>
OME	Tremolite Skarn	6300 m/s	2.921 g/cm <sup>3</sup>	18,402 m/s × g/cm <sup>3</sup>
OUM	Serpentinite	6200 m/s	2.886 g/cm <sup>3</sup>	17,895 m/s × g/cm <sup>3</sup>
S/MS	Massive Sulfide	6100 m/s	3.239 g/cm <sup>3</sup>	19,756 m/s × g/cm <sup>3</sup>

## 5.2. Simulation Parameters and Results

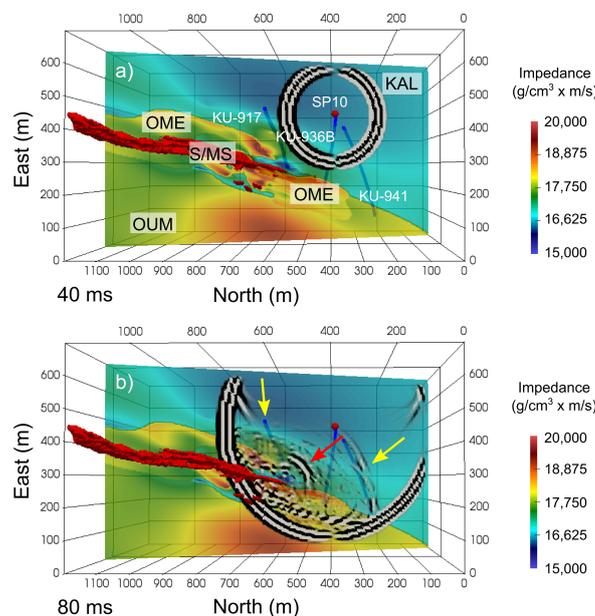
For our forward modeling study, we used the free software SOFI3D [66], which is a viscoelastic, three-dimensional modeling tool using a parallel Finite-Difference (FD) approach. The modeling code SOFI3D (git-master-275248a88, 14 July 2016, Karlsruhe Institute of Technology, Karlsruhe, Germany) is capable of viscoelastic modeling; however, we only did elastic modeling, i.e., no attenuation was included in this study. Furthermore, no petrophysical data were available for S-wave velocities. Thus, those were scaled from the P-wave velocity data by a constant factor of  $1/\sqrt{3}$ . We combined the constructed input models with the VSP acquisition geometry that was used in the Kylylahti survey (Figure 3). With this setup, we computed synthetic VSP shot records along boreholes KU-917 (DAS and 3C VSP configurations), KU-936B and KU-941. As a source signal, we used a Ricker wavelet with a central frequency of 180 Hz, which equals the dominant frequency of the acquired VIBSIST-200 records. All key parameters are summarized in Table 3.

**Table 3.** Summary of technical parameters used for the seismic forward modeling experiment. FD, Finite-Difference.

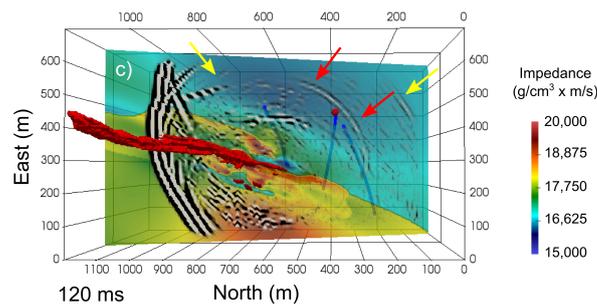
Modeling Parameter	Configuration
Modeling software	SOFI3D
Computation mode	Elastic
FD scheme	8th order, Holberg stencil
Source wavelet	Ricker, 180-Hz central frequency
Model extent	1200 m × 700 m × 1200 m
Grid spacing	1 m
Minimum P-wavelength	16.1 m
Simulated shot length	360 ms
Time sampling	0.05
CPU time	26 min on 1000 CPUs

The forward modeling experiment provides synthetic VSP shot records and wavefield snapshots that can be tracked through the input model. Figure 7 shows three snapshots at three different time steps. They illustrate how the seismic wavefield propagates from one of the VIBSIST-200 shot locations (SP10) towards the mineralization zone and gets reflected back from the various geological contrasts. The simulation helps to understand the complexity of the recorded seismic wavefield, identify the most dominant reflections and link them to their geological causes.

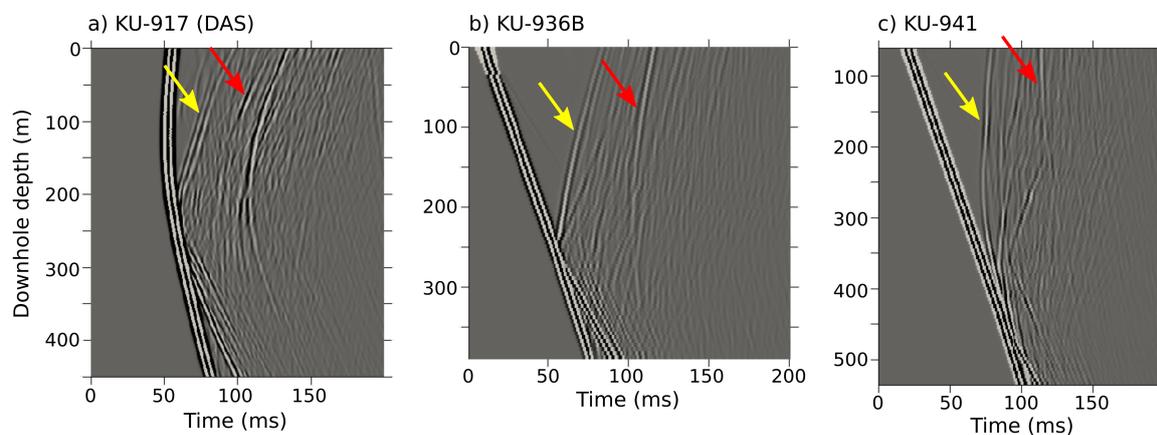
A selection of synthetic shot records obtained from the simulation is shown in Figure 8. In combination, the synthetic data and snapshots reveal that the first strong signal that is recorded by the VSP receivers (yellow arrow in Figures 7 and 8) originates not directly from the target S/MS deposit, but from the contrast between the Kaleva sediments (KAL) and the Outokumpu alteration fringes (OME). The sulfide reflection (red arrow in Figures 7 and 8) appears later in the synthetic shot record and is more difficult to identify than the previous one. These observations help in guiding the geological interpretation of the migrated shot records.



**Figure 7.** Cont.



**Figure 7.** Wavefield snapshots over acoustic impedance model at time steps of 40 ms (a), 80 ms (b) and 120 ms (c). The semi-massive to massive sulfide mineralization is depicted as a 3D red surface. Virtual receivers for boreholes KU-917, KU-936B and KU-941 are shown as blue solid lines. The source position for shot SP10 is shown as a red sphere. Yellow arrows in (b,c) highlight the reflection from the KAL-OME contrast, whereas red arrows denote a reflection arising from the S/MS deposit. The synthetic shot records are shown in Figure 8.



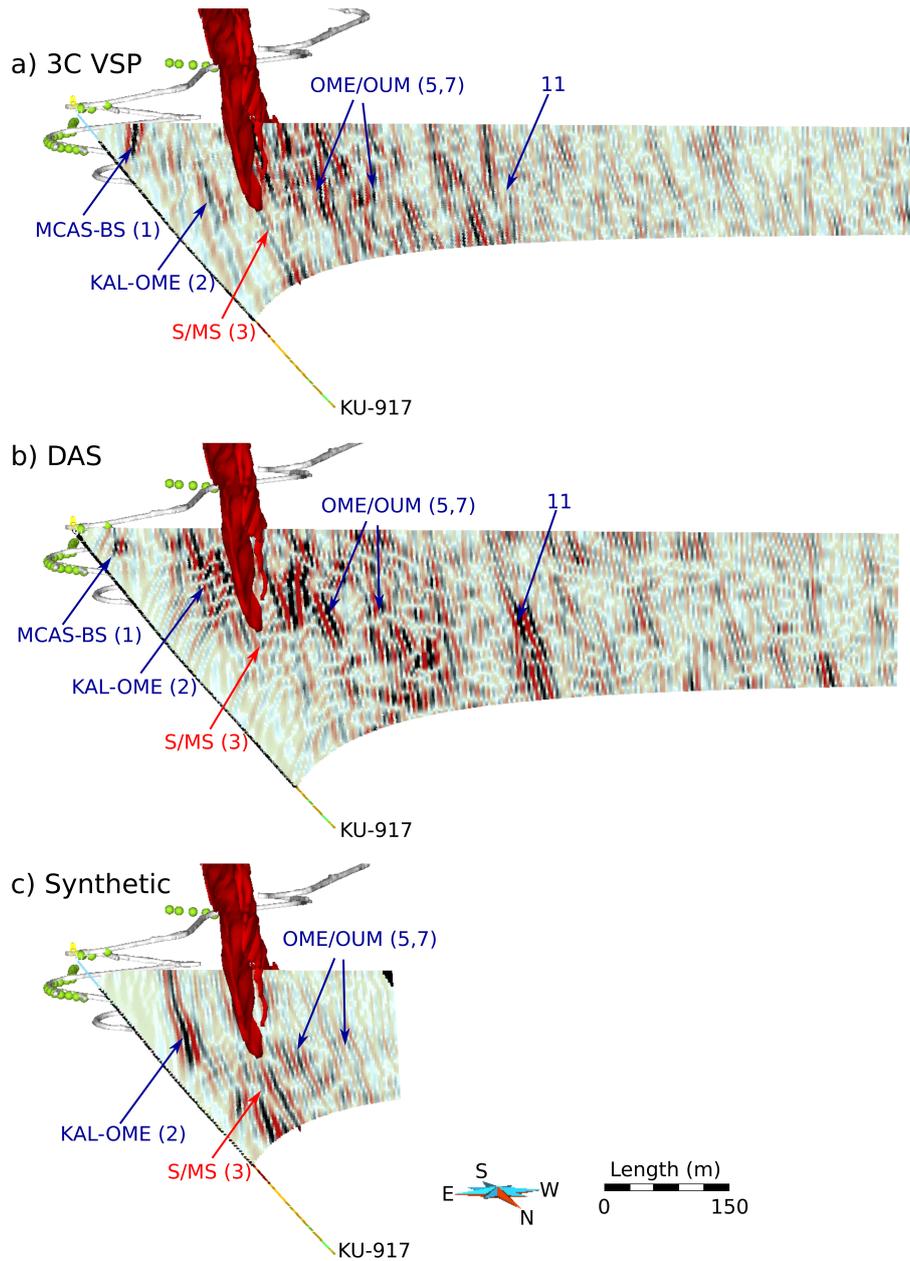
**Figure 8.** Synthetic VSP shots obtained from seismic forward modeling for virtual receivers along borehole KU-917, DAS configuration (a), KU-936B (b) and KU-941 (c). Arrows indicate reflections from the KAL-OME contrast (yellow) and the S/MS deposit (red) (compare to Figure 7).

## 6. VSP Imaging and Interpretation

### 6.1. Comparison of 3C, DAS and Synthetic Migrated Images

The synthetic VSP shots were processed alongside the acquired data and finally migrated using 3D image point migration. The VSP migration maps the processed VSP records from the time-domain into 3D space, in a similar way as common-shot pre-stack depth migration does for surface-based seismic data. We used a constant velocity of 6 km/s, which represents the average velocity that was derived from the VSP first-arrivals. As a result, we obtain migrated VSP records for each shot-borehole configuration. These individual images could not be unambiguously stacked together due to the uneven shot layout, which is limited to accessible mine infrastructure. Thus, the obtained images were primarily interpreted separately.

Figure 9 shows a comparison of resulting images from shot SP06 that were recorded using the conventional geophone and the DAS VSP, respectively. In addition, these recorded data are compared to a synthetic VSP shot that was processed and migrated in the same manner. All images are directed in an approximately perpendicular direction towards the location of the known sulfide deposit.



**Figure 9.** Comparison between conventional geophone (a) and DAS VSP (b) migrated images from acquired data and from synthetic VSP data (c) as viewed from north to south. The synthetic image is limited towards the west, due to the lack of geological information in that area. Blue labels denote interpreted rock contrasts. The reflector numbering is explained in the next section (Table 4). Seismic sources are indicated by green spheres in the mine tunnels.

The comparison reveals similarities and differences between the results of the two recording technologies. In the area between the shot position and the sulfide deposit, both acquired VSP images evidence reflections from the predominant geological contrasts KAL-OME and S/MS, which are also contained in the synthetic image. Moreover, the acquired data reveal a contrast between the mica schist and black schist (MCAS-BS) units, which is not contained in the forward model and therefore does not exist in the synthetic image.

The reflection from the target sulfide deposit (S/MS) is evidenced by all images; however, it does not appear as clearly in the acquired datasets as it is in the synthetic shot. In general, the synthetic reflections appear more coherent, which naturally can be attributed to the simplifications of the

model (compared to reality), the absence of mine tunnel structures and the absence of noise in the synthetic data.

West of the S/MS mineralization lens, the acquired images show several reflections, which originate from rock contrasts within the Outokumpu assemblage zone (OME and OUM). As far as the existing boreholes extend to the west, the imaged contacts were checked against the available logging data.

Here, the agreement between the synthetic and the acquired data images is less clear, since the velocity model does not differentiate between rock types in that area. Thus, synthetic reflections are mainly caused by variations in the density model and therefore less pronounced.

In addition to the reflections that occur in the close vicinity of the sulfide mineralization, there are a number of other features located further to the west in the acquired images (e.g., Reflection 11). These features get weaker with further distance from the source locations due to the loss of seismic energy. As the coverage of boreholes rapidly decreases west of the target deposit, the geological model only contains very little information in that area. For that reason, the synthetic migrated image does not extend into the far west.

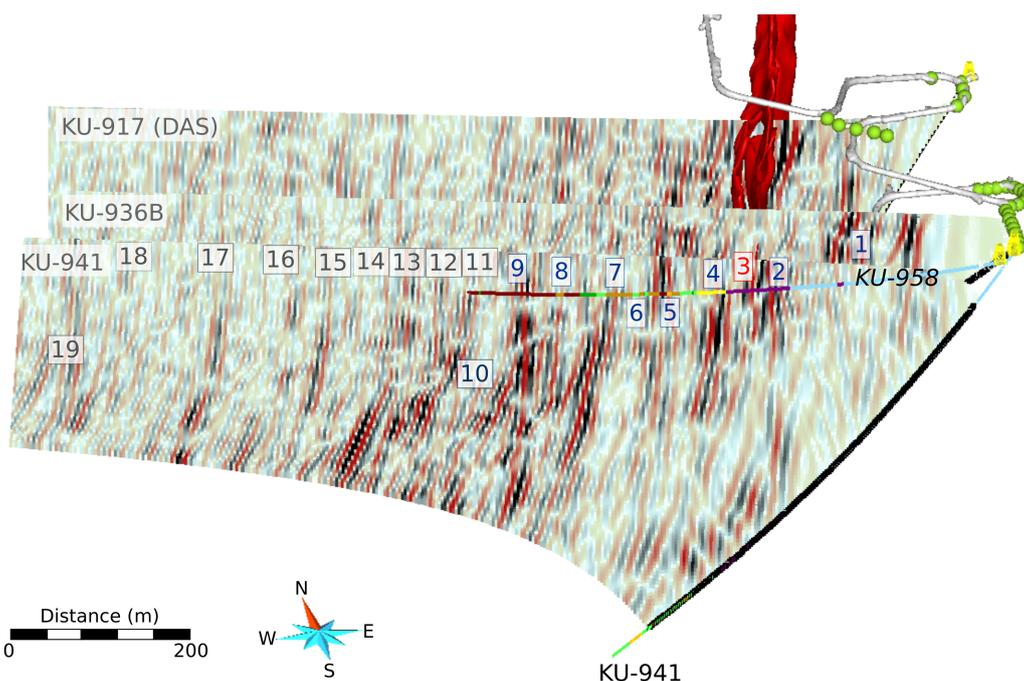
In terms of quality, the differences between finally migrated images obtained from the conventional geophone and the DAS VSP are less significant than between their respective raw shot gathers as a result of the applied data processing. The DAS VSP extends a bit further in depth due to the longer borehole coverage of the fiber-optic cable. The 3C data still appear slightly superior in terms of spatial resolution than the DAS. In particular, the target sulfide reflection is more pronounced in the geophone data. Moreover, the MCAS-BS and KAL-OME reflections are overprinted by coherent noise artifacts in the DAS image, which result from the cable ringing in the top parts of the borehole, as already described in the raw data comparison.

## 6.2. Geological Interpretation and Correlation with Borehole Data

In total, 83 individual migrated shot records were obtained from the acquired VSP data. These images were analyzed jointly to trace predominant reflections through the illuminated area parallel to the target sulfide mineralization (Figure 10). In this way, 19 distinct reflections were identified that were evidenced by the majority of the computed reflectivity images. In particular, 14 out of 19 contrasts were evidenced by at least 65%, and six out of 19 (Reflections 3–8) were evidenced by at least 75% of the migrated images. Most of the less supported features occur further to the west.

Direct relation of the identified reflections to their petrophysical causes is often ambiguous due to the highly complex and rapidly changing geology that particularly characterizes the Outokumpu assemblage rocks. Nevertheless, cross-validation of the migrated images with logging data from 135 intersecting or adjacent boreholes found that Reflections 1–8 and 10 can be attributed to lithological contrasts (Table 4). Reflections 11–19 are too far away from the target sulfide deposit, and thus, no logging data are available to support these features.

The majority of the boreholes used for correlation were drilled before the Kylylahti survey was conducted in order to explore the immediate vicinity of the target sulfide lens. However, some boreholes were drilled after the campaign. In particular, borehole KU-958 was drilled nearly horizontally through the formation in order to validate the predicted features that had been obtained from the initial interpretation of the VSP survey.

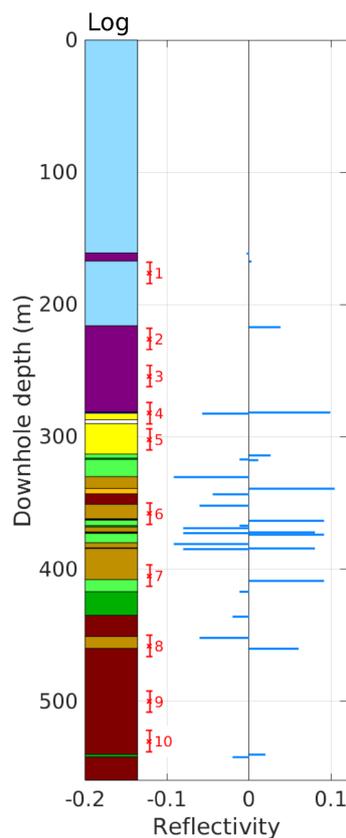


**Figure 10.** Interpretation of migrated reflection images obtained from all boreholes. The displayed images were obtained from shots SP26 (KU-917 and KU-936B) and SP13 (KU-941). Seismic sources are indicated by green spheres in the mine tunnels. Boxed numbers mark the identified reflections; blue numbers (1–10) are backed up by logging data, whereas gray numbers (11–19) represent new features. Reflector 3 (red) represents the target sulfide lens. Borehole KU-958 was drilled for validation purposes after the initial interpretation was obtained.

**Table 4.** Evidence for identified Reflections 1–10 from 135 borehole logs. Boreholes were used for validation if they intersect or are located closely to a migrated VSP image. If the log contained a lithological contrast within one wavelength of the identified reflection, it was counted as evidence. Otherwise, it counted as a contradiction.

Reflector	Associated Contrast	Evidences	Contradictions
1	Mica schist-black schist (MCAS-BS)	54	0
2	Black schist-carbonate rock (KAL-OME)	56	0
3	Black schist-carbonate rock-massive sulfide (S/MS)	59	2
4	Quartz rock-carbonate-skarn rock	71	3
5	Skarn-quartz rock-chlorite schist	28	2
6	Serpentinite-(talc-)carbonate rock	11	0
7	Skarn-serpentinite-quartz-talc-carbonate rock	5	0
8	Serpentinite-talc-carbonate rock	1	0
9	No evidence	0	1
10	Serpentinite-chlorite schist	1	0

Figure 11 shows a comparison between the KU-958 log, a synthetic reflectivity series derived from petrophysical measurements of the logged rock types, and the identified reflections from the closest migrated image. It demonstrates how most reflections identified from VSP migration results correspond to lithological contrasts evidenced in the borehole log. Very closely-spaced interfaces (e.g., between Reflections 4 and 7) could not be individually resolved by the VSP method due the inherent limitations of spatial resolution. The dominant P-wavelength (16.1 m) is depicted by error intervals in Figure 11. Moreover, Reflectors 2, 3 and 9 are not evidenced in the log.



**Figure 11.** Comparison between the KU-958 borehole log and synthetic reflectivity series derived from petrophysical velocities and densities. Red numbers denote the reflectors picked from the VSP image obtained from borehole KU-941 (Figure 10); error intervals are equal to the dominant P-wavelength.

Feature 2 most likely corresponds to a black schist to carbonate rock contrast; Feature 3 represents the target sulfide mineralization. Due to the rapidly changing and complex geology in that zone, it is possible that these contrasts simply do not extend far enough to the south to “reach” borehole KU-958, and hence, they are not evidenced in the log. However, their existence is supported by numerous other boreholes (Table 4).

In contrast, Reflector 9 is located too far to the west, and thus, no other exploration boreholes are available to be used for correlation.

## 7. Discussion and Conclusions

The Kylylahti underground VSP experiment presented in this study provides several important conclusions. In more general terms, it demonstrated the potential value of underground VSP measurements for mineral exploration and resource delineation, which is due to various reasons. (1) The underground VSP measurements are conducted in close proximity to the imaging target. (2) They are not affected by the overburden layer. (3) The use of borehole acquisition provides better illumination of steeply-dipping geological features. (4) The data provide very high spatial resolution, which allows precise mapping of geological contrasts.

However, the inherent directional ambiguity of VSP measurements also requires complementary information from geological logging and potentially surface-seismic data. These data provide important input for optimizing VSP processing parameters in order to constrain the dips and azimuths of recorded reflected signatures.

In this study, migrated VSP images from 83 shot records were analyzed individually and cross-validated against 135 boreholes that had been drilled in the target area. This joint analysis

enabled identification of 19 pronounced reflectors from the VSP images that area supported by the majority of the migrated shot records. Ten out of the 19 reflectors are within the reach of existing boreholes and thus are additionally supported by borehole logging data. The remaining nine reflectors are located outside of the explored area and thus cannot be validated by boreholes, but may represent future drilling targets. The analysis confirms that underground VSP surveys are capable of precise mapping of geological contrasts and resource delineation and thus can be used complementary to surface-based seismic methods, which usually operate on a larger scale.

More specifically, this study focuses on the application and comparison of a conventional geophone VSP and a fiber-optic DAS recording system. Both technologies successfully recorded seismic reflections arising from nearby geological contrasts. The obtained data records were processed and imaged alongside each other using the same workflow. Based on the results, we conducted a qualitative comparison of the raw and the processed data obtained from both recording systems.

As a general conclusion, the conventional VSP data were found to be superior with respect to the signal-to-noise ratio and bandwidth, which results in a better spatial resolution of the conventional 3C geophone system. The resolution of the DAS is effectively governed by the gauge length applied during the survey. In addition, the DAS system was more prone to noise in the absence of borehole cementation. Furthermore, the DAS naturally cannot provide three-component data, but only outputs one component. Despite these differences, data processing narrowed the gap in terms of overall quality between the systems. Eventually, both technologies provided valuable reflectivity images that were used for joint geological interpretation.

The main benefit of the DAS technology is its low logistical effort and ability to acquire data over the whole length of the borehole at once, which makes the data acquisition faster, reduces the required number of shots and therefore the cost for surveying a complete borehole. For this reason, DAS can either be used to accomplish a VSP survey in much less time or it can be used to survey additional boreholes in order to increase the illumination of the imaging target. We believe that the value of acquiring additional data from different boreholes outweighs the benefits of the three-component nature of conventional VSP data in terms of target illumination. Of course, this requires the availability of additional boreholes and shot point locations, which would improve azimuthal coverage of the target.

We conclude that DAS VSP surveys provide a very promising tool for mineral exploration and mine planning. We believe that the application of in-mine VSP surveys could generally be used to plan ongoing exploration drilling more strategically and thus potentially reduce the number of required boreholes. With the use of DAS technology, this approach would become even more efficient.

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## Abbreviations

The following abbreviations are used in this manuscript:

3C	Three-component
BS	Black schist
DAS	Distributed Acoustic Sensing
FD	Finite-difference
KAL	Kaleva sedimentary belt
MCAS	Mica schist
OME	Altered Outokumpu ultramafics
OUM	Outokumpu ultramafics
S/MS	Semi-massive to massive sulfide
VSP	Vertical Seismic Profiling

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