

Article

Importance of Magmatic Water Content and Oxidation State for Porphyry-Style Au Mineralization: An Example from the Giant Beiya Au Deposit, SW China

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Abstract: The Beiya Au deposit is the largest Cenozoic Au deposit in the Jinshajiang-Ailaoshan porphyry metallogenic belt. Numerous studies document that high water content and fO_2 are vital factors for the generation of Au mineralization. In this belt, only the Wandongshan and Hongnitang districts are considered to be of economic importance, while the other districts, such as Bailiancun, are barren. So in order to reveal the importance of water content and oxidation state for Beiya porphyry-style Au mineralization, the amphiboles and zircons compositions are used to evaluate the physicochemical conditions (e.g., pressure, temperature, fO_2 , and water content) of the Wandongshan ore-ferile porphyries and Bailiancun ore-barren porphyries observed in the Beiya Au deposit. The results show that the water content of the Wandongshan parent magma ($\leq 4.11 \pm 0.4$ wt %) are slightly higher than those of the parent magma at Bailiancun ($\leq 3.91 \pm 0.4$ wt %), while the emplacement pressure of the Wandongshan parent magma (31.5–68.6 MPa) is much lower than that of the parent magma at Bailiancun (142.3–192.8 MPa), indicating that the Wandongshan magma reached water saturation earlier. In addition, the Wandongshan porphyries crystallized from more oxidized magma (average of $\Delta FMQ = +3.5$) with an average temperature of 778 °C compared to the Bailiancun porphyries (average of $\Delta FMQ = +1.5$) with a mean magmatic temperature of 770 °C. The Ce^{4+}/Ce^{3+} ratio of zircon in the Wandongshan ore-related intrusions (average Ce^{4+}/Ce^{3+} of 62.00) is much higher than that of the Bailiancun barren porphyries (average Ce^{4+}/Ce^{3+} of 23.15), which further confirmed Wandongshan ore-related magma is more oxidized than the Bailiancun barren magma. Therefore, melts that are more enriched in water and with a high oxidation state will be more fertile to form an economic porphyry-style Au system.

Keywords: oxidation state; hydrous melts; constraints on mineralization; Beiya Au deposit; SW China

1. Introduction

Porphyry systems supply three-quarters of the world's Cu and one-fifth of the world's Au resources [1]. It's widely accepted that high water content (>4 wt %) [2,3] and a high oxidation state ($>FMQ + 1.3-2$) [4,5] of the magma are critical factors for the generation of Cu-Au mineralization. The solubility of Cu and Au in hydrous silicate melts was proven to be affected by the amount and proportion of dissolved sulfate and sulfide in S-bearing magma, whereas fO_2 controls the speciation of sulfur [5–7]. Under high fO_2 conditions, the majority of sulfur exists as SO_2 or SO_4^{2-} with very low concentrations of S^{2-} in a magma [6], which prevents saturation of immiscible Cu- and Au-sulfides during magmatic fractionation [4,6]. High magmatic water content of the magma ($H_2O > 4$ wt %) is

also an essential factor for the generation of porphyry Cu-Au deposits [8–11]. The high water content of the magma not only promotes magmatic water reaching saturation earlier [10], but also increases the partitioning of Au from melt to aqueous volatile phases [12,13]. However, the impacts of water content and magma oxidation state on Au mineralization are still debated.

The Beiya Au deposit is the largest Cenozoic Au deposit in the Jinshajiang-Ailaoshan porphyry metallogenic belt [14,15], with proven Au reserves of more than 377 t, and by-products of Cu, Fe, Ag, Pb, and Zn [16]. Numerous studies have documented the petrography, chronology, geochemistry, and fluid inclusions of the porphyry intrusion from the Beiya Au deposit [17–19]. These studies have revealed that the monzogranite porphyry is directly related to skarn formation and mineralization [17,18]. The ore-forming monzogranite porphyry at Wandongshan is water-rich [13,19] and has a high oxidation state [20,21]. However, systematic studies on the relationship of water content and oxidation state on Au mineralization at Beiya are lacking. Comparing the water content and oxidation state of the mineralized and barren porphyries in the Beiya Au deposit is important for understanding the importance of water content and oxidation state for Beiya porphyry-style Au mineralization.

Both the monzogranite porphyries at Wandongshan and Bailiancun contain amphibole phenocrysts [22], even though the Wandongshan is considered to be of economic importance and the Bailiancun is barren [22–24]. Thus, we choose the Wandongshan and Bailiancun as the focus of our study. We investigated the mineralogical and geochemical characteristics of the amphiboles and zircons from the Wandongshan and Bailiancun monzogranite porphyries. In order to determine the relationship of fO_2 and water content on Au mineralization, we compared the water content and oxidation state of the mineralized Wandongshan ore-bearing and Bailiancun barren monzogranite porphyries.

2. Geological Setting

The Sanjiang (Nujiang-Lancangjiang-Jinshajiang) region is an important part of the Tethys tectonic domain, located in Southwestern China [25,26]. The Jinshajiang-Ailaoshan suture zone is located in the eastern Himalaya-Tibetan plateau of the Tethyan domain and is bounded by the western Yangtze Craton to the southeast (Figure 1). The regional tectonic evolution includes subduction of the Paleo-Tethyan oceanic slab, and magmatic arc formation followed by arc-continent collision during the Paleozoic-Mesozoic era, and then Indo-Asian continental collision during the Cenozoic [27,28]. The collision between the Indo and Asian continents have caused more than 300 km of strike-slip movement along the Ailaoshan-Red River shear zone since the Paleocene [29,30]. As a consequence of this strike-slip, lithospheric-scale extension has occurred, causing a set of secondary NNW-trending strike-slip faults and folds, and the emplacement of numerous alkali-rich intrusions along the Jinshajiang suture zone [31–33]. These intrusions form a 2000-km-long and 50- to 80-km-wide alkaline magmatic belt, and are associated with several important porphyry and skarn polymetallic (Cu, Au, Mo, Zn, Pb, and Ag) deposits [34,35] in the Jinshajiang-Ailaoshan metallogenic belt (Figure 1b) [7], from north to south: the Yulong Cu, the Beiya Au, the Machangqing Cu-Mo-Au, the Yao'an Au, the Habo Cu-Mo-Au, and the Tongchang Cu-Mo deposits.

The Beiya Au deposit is located in the center of the Jinshajiang-Ailaoshan magmatic belt (Figure 1a), which covers an area of ~800 km² and is distributed along the limbs of the N-S trending Beiya syncline [16]. The Beiya Au deposit consists of two zones [22,23], with the Bijiaoshan (BJS), Weiganpo (WGP), and Guogaishan (GGS) ore districts to the east, the Wandongshan (WDS), Hongnitang (HNT), and Jingouba (JGB) ore zones to the west, and the Bailiancun (BLC) occur around the periphery of the Beiya Au deposit (Figure 2). The formation hosting the Wandongshan district are dominated by Permian basalt, Triassic sandstone and limestone, with Quaternary alluvial cover, while the strata hosting the Bailiancun district are mainly Permian basalt and Quaternary alluvial cover. There are abundant Tertiary alkaline porphyry intrusions distributed in the Wandongshan and Bailiancun districts. The Tertiary alkaline porphyry intrusions include monzogranite porphyries, biotite monzogranite porphyries, and lamprophyres (Figure 2a). The monzogranite porphyries

represent the volumetrically most important igneous rock and are genetically related to the mineralization both at Wandongshan and Bailiancun districts [15,25]. Though the monzogranite porphyries at Wandongshan and Bailiancun districts have similar zircon U-Pb ages (36.2 ± 0.6 Ma and 36.1 ± 0.15 Ma, respectively) [15,17] and the same magma source [22], their mineralization types are different. The Wandongshan monzogranite porphyries form economical porphyry-skarn Fe-Cu-Au mineralization [22–24], while the Bailiancun monzogranite porphyries form hydrothermal Pb-Zn veins [22].

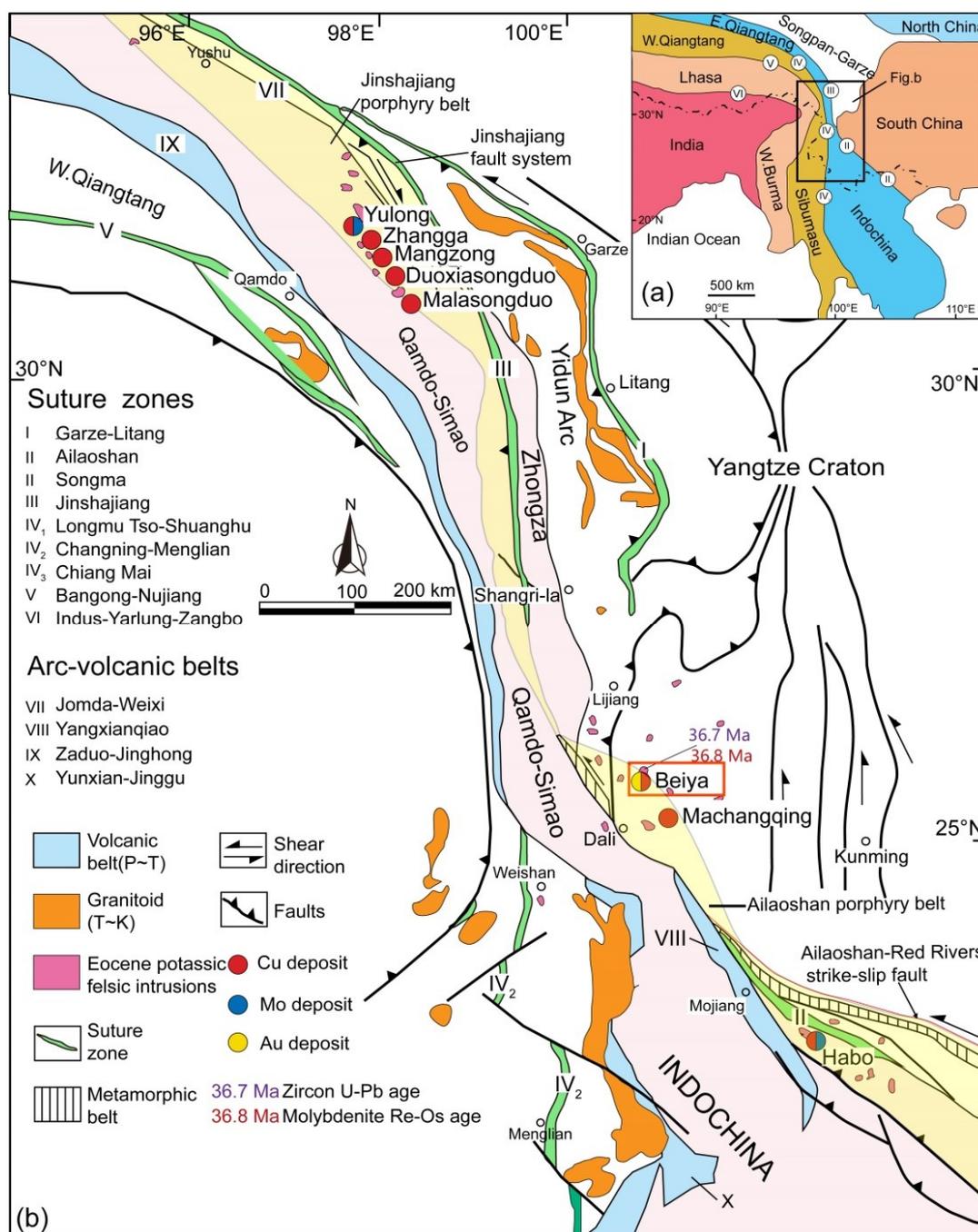


Figure 1. Regional tectonics of the Jinshajiang-Ailaoshan alkaline porphyry belt showing the distribution of Cenozoic deposits [17].

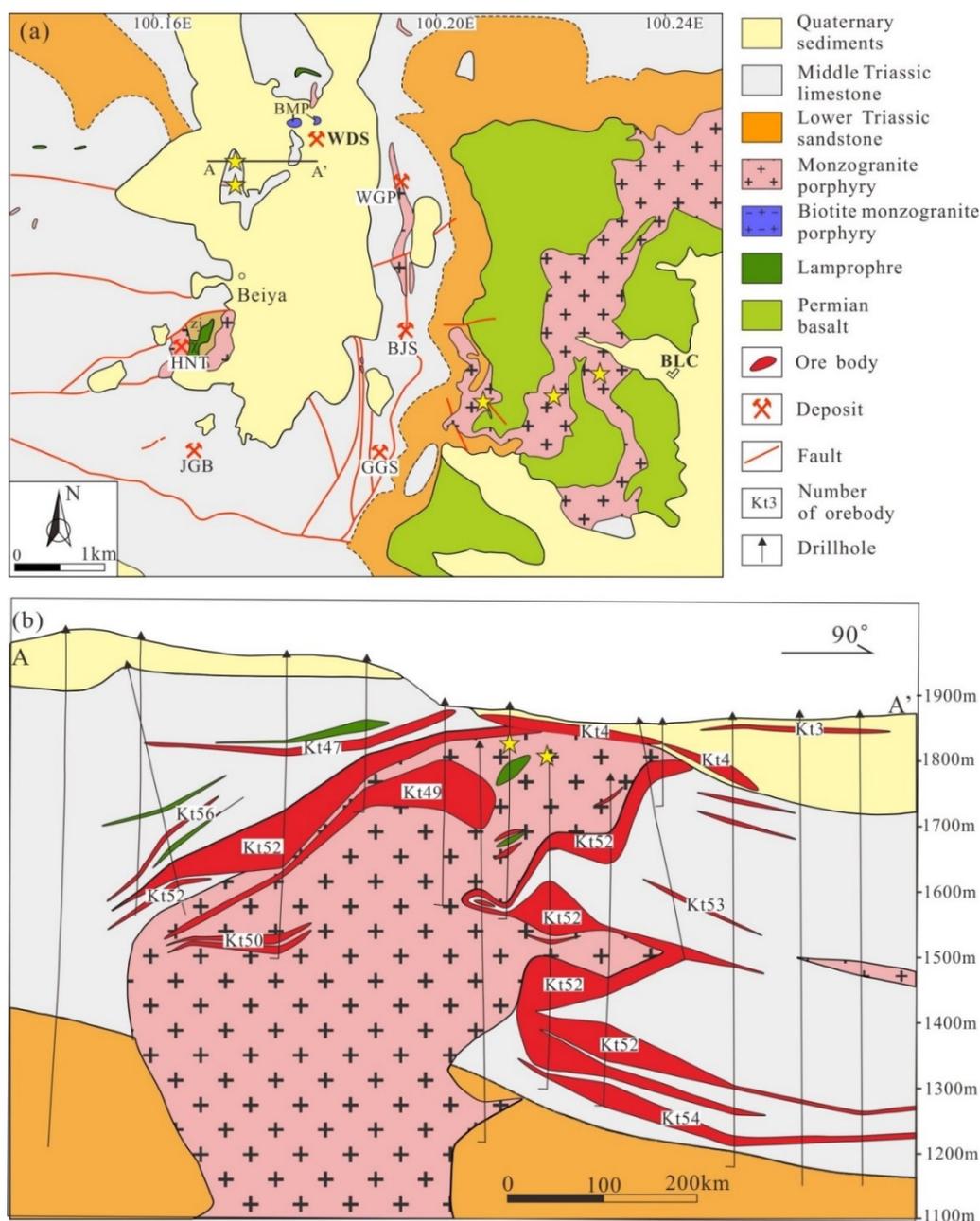


Figure 2. Geology of the Beiya Au deposit (a) and representative crosssection of Wandongshan district showing the location of the skarn ore body (b) [23]. WDS: Wandongshan district; HNT: Hongnitang district; JGB: Jin'gouba district; WGP: Weiganpo district; GGS: Guogaishan district; BJS: Bijiashan district; BLC: Bailiancun district.

3. Petrography of Alkaline Intrusions

3.1. The Wandongshan Monzogranite Porphyry

The Wandongshan monzogranite porphyry intrusion has a gray-white color and porphyritic textures. Phenocrysts form about 50 vol % of the rock, and include K-feldspar, plagioclase, quartz, biotite, and minor amphiboles (Table 1). The K-feldspar phenocrysts make up about 5–15 vol % and usually have euhedral to subhedral shapes that range in size from 0.2 to 0.5 mm, with occasional Carlsbad twinning (Figure 3a). The plagioclase phenocrysts (1–5 mm) are mostly euhedral in shape, showing tabular biotite inclusions (Figure 3c). The quartz phenocrysts are irregular, with a circular

shape. The amphibole phenocrysts are mostly euhedral, usually measuring 0.4 mm in size (Figure 3b). The groundmass comprises fine-grained feldspar and quartz. The accessory minerals are apatite, magnetite, titanite, and zircon. The mineralization is dominated by pyrite and chalcopyrite within quartz veins, as well as disseminated in the monzogranite.

Table 1. Petrography of the monzogranite porphyry from the Wandongshan and Bailiancun ore districts.

Ore Deposits	Rock Type	Phenocrysts	Groundmass	Accessory
Wandongshan	Monzogranite porphyry	Kfs + Qtz + Pl + Bt + Amp	Pl + Qtz	Ap + Zrc + Ttn + Mag
Bailiancun	Monzogranite porphyry	Qtz + Kfs + Pl + Amp	Kfs + Qtz + Amp	Ap + Zrc

Qtz: quartz; Kfs: K-feldspar; Pl: plagioclase; Bt: biotite; Amp: amphibole; Ap: apatite; Zrc: zircon; Ttn: titanite; Mag: magnetite.

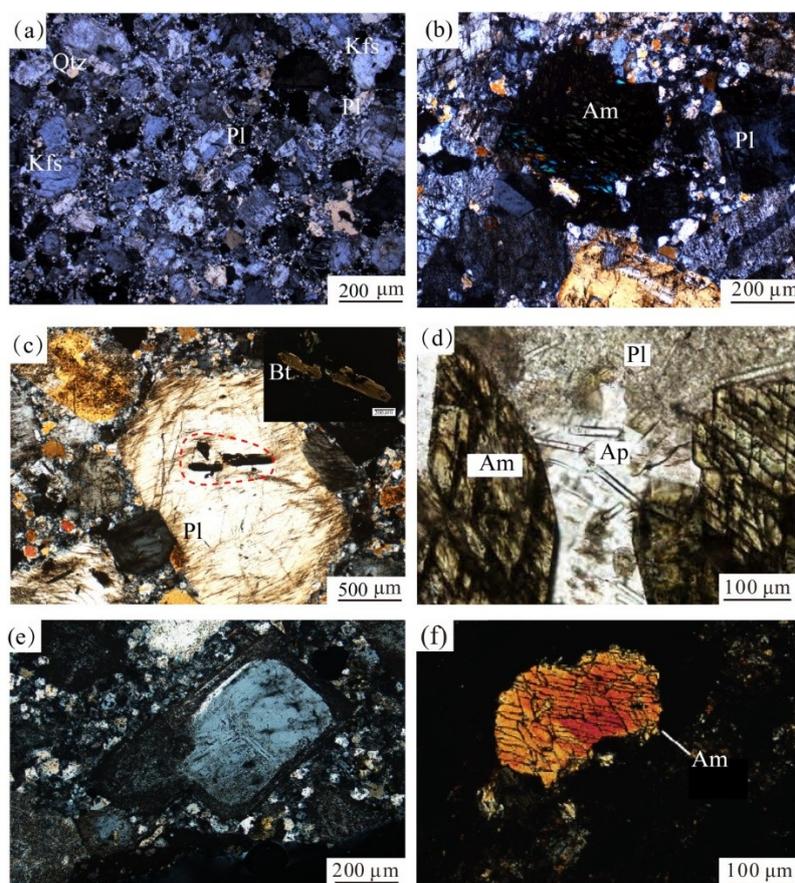


Figure 3. Photomicrographs of the monzogranite porphyries at the Wandongshan and Bailiancun districtdistricts. (a) The monzogranite porphyries at the Wandongshan deposit; cross-polarized light; (b) Plagioclase and amphibole phenocrysts of the Wandongshan monzogranite porphyries; cross-polarized light; (c) Biotite is inside the plagioclase; cross-polarized light; (d) Dioritic enclaves from Wandongshan; cross-polarized light [23]; (e,f) are photographs of monzogranite porphyries at the Bailiancun district; cross-polarized light. Qtz: quartz; Kfs: K-feldspar; Pl: plagioclase; Bt: biotite; Am: amphibole; Ap: apatite.

3.2. The Bailiancun Monzogranite Porphyry

The monzogranite porphyry intrusion at the Bailiancun deposit is characterized by porphyritic texture. Phenocrysts consist of K-feldspar, quartz, plagioclase, and amphibole. The groundmass includes K-feldspar, quartz, and amphiboles (Table 1). The accessory minerals are dominated by apatite and zircon. The K-feldspar phenocrysts make up about 30 vol % with a euhedral to subhedral

habitus, and measuring 0.4 mm in size (Figure 3e). The quartz phenocrysts have subrounded circular shapes. The plagioclase phenocrysts account for 25 vol % and are mostly euhedral. The amphibole phenocrysts are small (0.3 mm), with euhedral shapes (Figure 3f).

4. Analytical Methods and Results

4.1. Analytical Methods

Electron microprobe analysis was conducted on amphiboles using a JEOL JXA-8800 Superprobe at the Institute of Mineral Resources, Chinese Academy of Geological Sciences. The microprobe was operated at an accelerating voltage of 15 kV, a beam current of 20 nA, and a beam size of 5 μm . Matrix corrections were performed using the ZAF correction program supplied by the instrument manufacturer. Chemical compositions of amphiboles can be used to evaluate the physicochemical conditions (e.g., pressure, temperature, $f\text{O}_2$, and water content) of the parental magma from which the amphibole crystallized [25,36,37].

Zircon trace elements were determined by laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) at the Key Laboratory of Crust-Mantle Materials and Environments, Chinese Academy of Sciences (CAS), Hefei. The LA-ICP-MS analyses used a Finnigan Neptune MC-IPS-MS with a high-performance interface coupled with a New-Wave UP-213nm Nd: YAG UV laser. The diameter of the laser ablation pit was 45 μm and the mean power output was 4 W. Signal and background measuring durations were about 70 and 40 s, respectively. Further analytical details are given in Liu et al. [38]. Zircon is a widespread accessory mineral that retains its primary chemical and isotope compositions from the time of crystallization [39], which means that the magma oxygen state during crystallization can be recorded by the zircon composition. Zircon solubility is mainly controlled by the temperature and major oxide concentrations of the magma [40]. Titanium exists as Zr^{4+} and Si^{4+} isotopes, so the Ti content are controlled by the temperature of the zircon crystallization [41–43]. This study uses the Ti content of the zircon crystals in zircon-saturated melts to estimate the temperature of the melt ($\log(\text{Ti}) = 6.01 \pm 0.03 - (5080 \pm 30)/T(\text{K})$) [42].

4.2. Mineral Chemistry

4.2.1. Amphibole Compositions

Major element compositions and chemistry of amphibole from the monzogranite porphyry at Wandongshan and Bailiancun districts are reported in Supplementary Materials, Table S1.

The amphiboles from the Wandongshan ore deposit are characterized by high Mg (MgO = 12.46–14.58 wt %), high Ca (CaO = 10.50–12.31 wt %), high Fe (FeO = 13.61–16.29 wt %), but low K (K_2O = 0.33–0.94 wt %) and Na (Na_2O = 0.55–2.18 wt %) content (Table S1). Compared to the Wandongshan amphiboles, the Bailiancun amphiboles are characterized by lower content of Mg (MgO = 9.57–9.90 wt %) and Ca (CaO = 9.54–10.85 wt %), higher content of Fe (FeO = 20.20–21.16 wt %) and K (K_2O = 1.22–1.44 wt %). The amphiboles from the Wandongshan porphyry are all edenites (Figure 4a) with high $\text{Mg}/(\text{Mg} + \text{Fe}^{2+})$ ratios (0.67–0.79) based on the classification of Leake et al. [41], whereas the amphiboles from the Bailiancun porphyry are all silicic edenites with low $\text{Mg}/(\text{Mg} + \text{Fe}^{2+})$ ratios (0.50–0.52; Table S1).

The emplacement pressure of the parental magma at Wandongshan and Bailiancun can be estimated using the Al-in-amphibole geobarometer [25,36]. These results show that the emplacement pressures of Wandongshan porphyries range from 31.5 MPa to 68.6 MPa, which corresponds to a paleodepth of 1.2–2.4 km (average 1.8 km) under lithostatic pressure (Table 2) [36]. In contrast, the emplacement pressures of Bailiancun porphyries are significantly higher (142.3–192.8 MPa), which corresponds to a deeper paleodepth of 5.4–7.3 km (average 6.7 km). Furthermore, water content are directly measure by the component of Wandongshan amphiboles ($\text{H}_2\text{O}_{\text{melt}} = 5.215^{[6]} \text{Al}^* + 12.28$) [25]. These results show that the water content of the Wandongshan parent magma exhibit

a wider range of 3.05 ± 0.4 wt % to 4.11 ± 0.4 wt %, whereas those of the Bailiancun parent magma have a tight range of 3.46 ± 0.52 wt % to 3.91 ± 0.4 wt % (Table 2, Figure 4b). Moreover, the amphibole composition is affected by the oxidation state of its parental melt, thus they allow the direct analysis of the oxygen state of their parental magmas [25]. The oxidation state deduced from amphiboles of the Wandongshan porphyry range from $\Delta\text{NNO} + 1.00$ to $\Delta\text{NNO} + 1.89$ (NNO is the nickel-nickel oxide oxygen buffer), slightly higher than that of the Bailiancun porphyries ($\Delta\text{NNO} + 0.39$ to $\Delta\text{NNO} + 0.67$, Table S1).

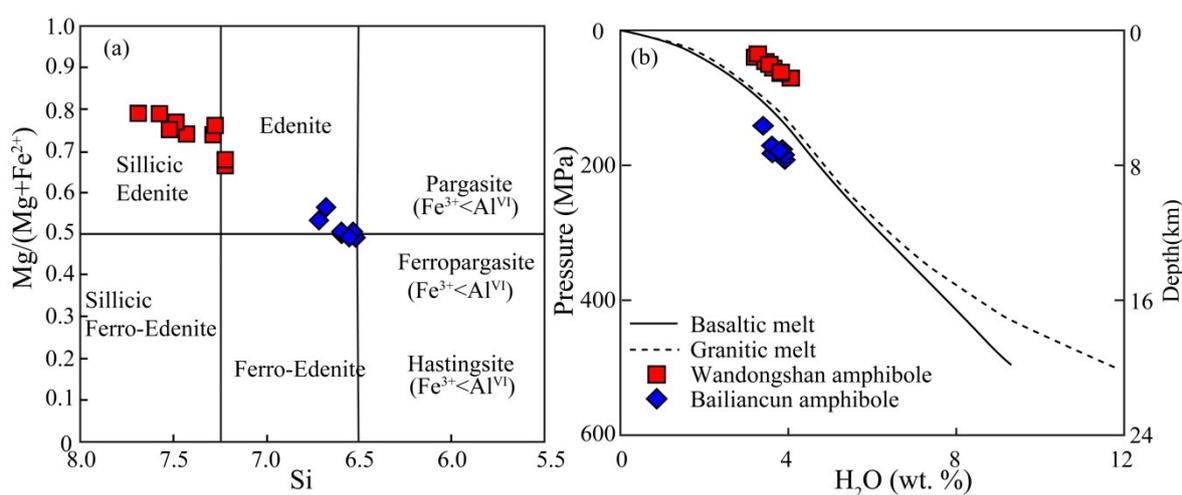


Figure 4. Composition of amphiboles from the Wandongshan and Bailiancun monzogranite porphyries, using the nomenclature of Leake et al. [42] (a), and experimentally determined solubilities of H_2O in silicate melts as a function of granitic pegmatite pressure (b) [44]. WDS: Wandongshan; BLC: Bailiancun.

Table 2. The water content and oxidation state of the parental magmas from the Wandongshan and Bailiancun districts.

Magmatic Geochemistry	Amphibole Phenocryst		Zircon			
	Pressure (MPa)	Water Content (wt %)	Oxidation State Index			
			ΔNNO	ΔFMQ	$\text{Ce}^{4+}/\text{Ce}^{3+}$	δCe
Wandongshan	31.5~68.6	3.05~4.11	+1.62	+3.5	12.1~144.6	9.25~203.83
Bailiancun	142.3~192.8	3.46~3.91	+0.5	+1.5	11.5~39.4	11.42~53.49

4.2.2. Zircon Trace Element Composition

Zircons from the Wandongshan monzogranite porphyries have ΣREE ranging from 149.0 to 1096.6 ppm, LREE/HREE (Low Rare Earth Elements/Heavy Rare Earth Elements) ratios varying from 0.01–0.15 (averaging 0.06), Ti content of 2.83–48.3 ppm, Ce anomalies varying from 9.25 to 203.83 (averaging 76.11), and Eu anomalies varying from 0.24 to 0.87 (averaging 0.6). However, zircons from the Bailiancun monzogranite have a lower average LREE/HREE ratio (0.04), average Ce anomaly (30.98), and higher average Eu anomaly (0.7) (Figure 5d, Table S2).

The zircon $\text{Ce}^{4+}/\text{Ce}^{3+}$ ratios (12.1–144.6, averaging 62.00) of the Wandongshan monzogranite porphyry intrusions are much higher than those of the Bailiancun monzogranite porphyry (11.5–39.4, averaging 23.15; Figure 5d). Moreover, the average ΔFMQ value of Wandongshan monzogranite porphyry is $\Delta\text{FMQ} + 3.5$, slightly higher than that of the Bailiancun monzogranite porphyry ($\Delta\text{FMQ} + 1.5$; Figure 5a). In addition, the Ti-in zircon temperatures of the Wandongshan monzogranite porphyry intrusions yield a wider range of 670.84–988.42 °C, averaging 778 °C, whereas those of the Beiya porphyry system exhibit a tight range of 680.17–818.53 °C, averaging 770 °C (Table 2).

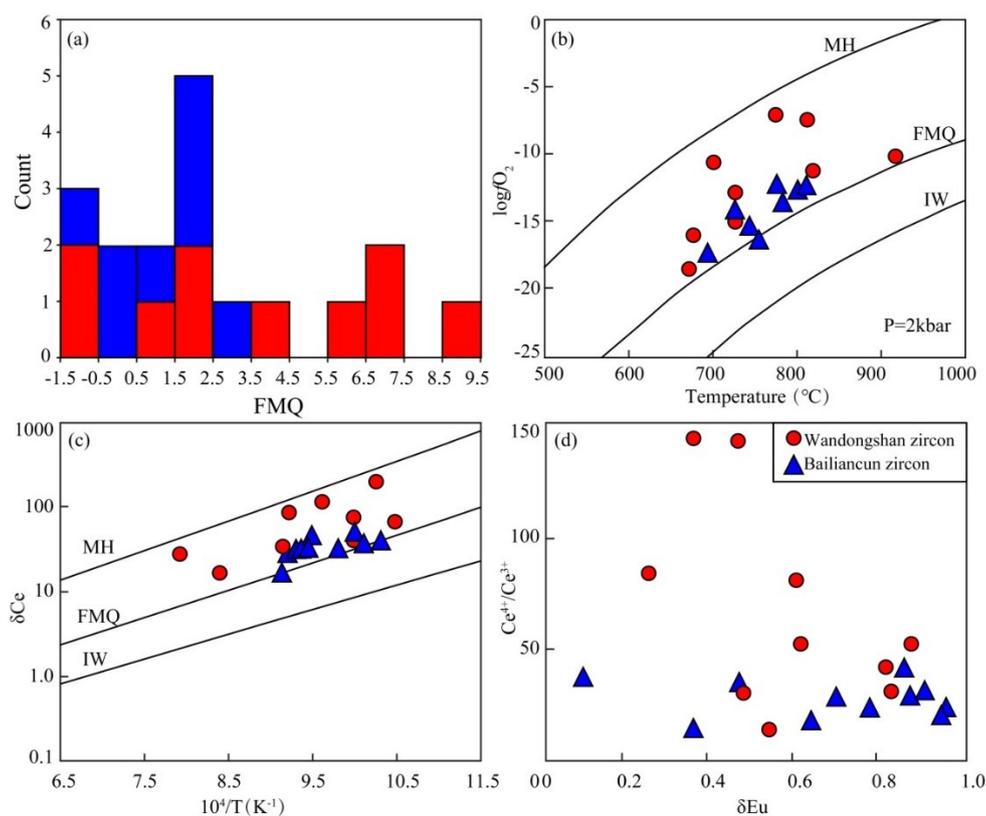


Figure 5. Magmatic oxidation state of intrusions from the Wandongshan and Bailiancun district, estimated from zircon. (a) Magma oxidation state of the Wandongshan and Bailiancun porphyries [41]; (b) Zircon $\log fO_2$ vs. temperature; (c) linear relations for zircon Ce anomaly vs. $10^4/T$; (d) zircon Ce^{4+}/Ce^{3+} vs. δEu [44]. MH: magnetite-hematite buffer curve; NNO: nickel-nickel oxide buffer curve; FMQ: fayalite-magnetite-quartz buffer curve; IW: iron-wustite.

5. Discussion

As mentioned above, the Wandongshan Au-bearing porphyries and Bailiancun barren porphyries have different physicochemical conditions (e.g., pressure, temperature, fO_2 , and water content). It is widely accepted that high oxidation state (high fO_2) and high H_2O content are two key factors responsible for Au mineralization [4,12,13].

5.1. The Water Content of the Parental Magmas

A high magmatic fluid flux ensures the exsolution of an aqueous volatile phase, which is considered as the sine qua non for magmatic hydrothermal ore-forming systems [2]. Both the Wandongshan and Bailiancun monzogranite porphyries contain amphibole phenocrysts. The amphibole phenocrysts at Wandongshan and Bailiancun are mostly euhedral in shape, and are characterized by embayed textures indicative of resorption (Figure 3b,f), suggesting that these amphiboles crystallized early. The presence of amphibole during the early stages of crystallization demonstrates that the parental magma at Wandongshan and Bailiancun districts are hydrous melts (>4 wt % H_2O) at elevated pressures [45,46]. Furthermore, the water content of the Wandongshan parent magma are characterized by a wider range of 3.05 ± 0.4 wt % to 4.11 ± 0.4 wt %, whereas those of the Bailiancun parent magma have a tight range of 3.46 ± 0.52 wt % to 3.91 ± 0.4 wt % (Table 2), which further confirmed the parental magma at Wandongshan and Bailiancun districts are relatively water-rich.

The emplacement pressure of the parental magma at Wandongshan and Bailiancun can be estimated using the Al-in-amphibole geobarometer [25,36]. These results show that the emplacement

pressures of Wandongshan porphyries range from 31.5 MPa to 68.6 MPa, but those of Bailiancun porphyries are significantly higher, ranging from 142.3 MPa to 192.8 MPa (Table 2). The solubility of H₂O in silicate magmas is mainly controlled by pressure and to a lesser extent by temperature [45]. The Wandongshan parent magma could reach water saturation at 31.5–70.9 MPa if they contain ~2.5% H₂O (Figure 4b). By contrast, the Bailiancun parental magma would need more water when they reach water saturation at 142.3–192.8 MPa. Therefore, it is expected that the Wandongshan parental magma will achieve water saturation easier, and the fluid exsolution would take place at 1.2–2.4 km, which consistent with the estimated trapping depth for Stage I S-type inclusions (~2 km depth under lithostatic conditions) [17].

5.2. Magmatic Oxidation State

The $\log(fO_2)$ vs. T and Ce anomaly vs. $1/T$ diagrams are divided into several oxygen fugacity fields by curves, which can be used to estimate the oxidation state of the magma. The data points of the Wandongshan monzogranite porphyry are mainly plotted within the field between magnetite-hematite buffer curve (MH) and fayalite-magnetite-quartz buffer curve (FMQ), while the fO_2 of the intrusions from the Bailiancun area are lower than those of the Wandongshan porphyries (Figure 5b,c). This result confirms that the porphyries from the Wandongshan district are more oxidized than those the porphyries from the Bailiancun district.

Furthermore, direct measurement of fO_2 also implies that the Wandongshan porphyries are more oxidized, the fO_2 deduced from amphibole chemistry of the Wandongshan porphyry range from $\Delta NNO + 1.00$ to $\Delta NNO + 1.89$ (NNO is the nickel-nickel oxide oxygen buffer), slightly higher than that of the Bailiancun porphyries ($\Delta NNO + 0.39$ to $\Delta NNO + 0.67$, Figure 6a). This result is consistent with the values calculated by zircon chemistry. The estimated fO_2 using zircon chemistry for the Wandongshan porphyry is $\Delta FMQ + 3.5$, slightly higher than that of the Bailiancun monzogranite porphyry ($\Delta FMQ + 1.5$; Figure 5a). Moreover, the Wandongshan monzogranite porphyries have an average zircon Ce^{4+}/Ce^{3+} ratio of 62.00 (Table 2), and an average zircon Ce anomaly of 76.11 (Figure 5c), much higher than those of the Bailiancun monzogranite porphyry (23.15, 30.98, respectively). Those differences above indicate the formation of Au-bearing porphyries requires elevated fO_2 .

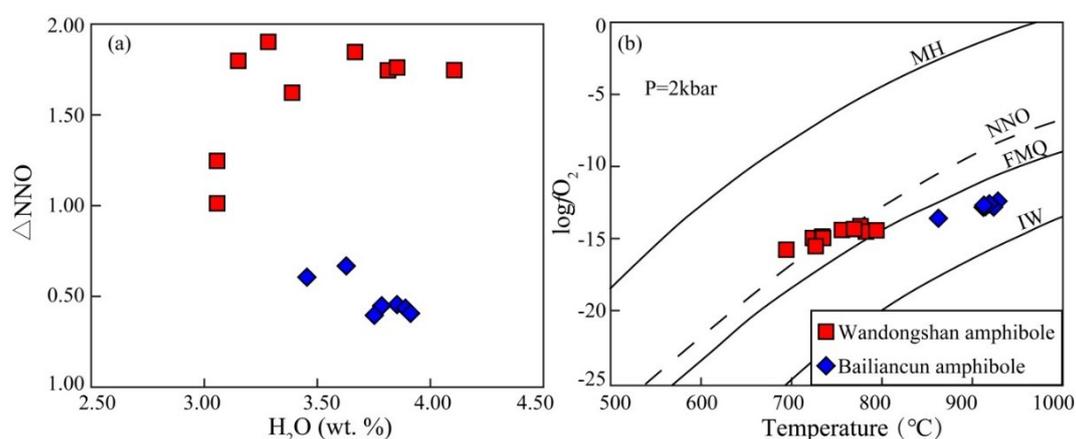


Figure 6. Magmatic oxidation state of intrusions in the Wandongshan and Bailiancun districts, estimated from amphibole compositions. (a) Magma oxidation state (ΔNNO) of Wandongshan and Bailiancun porphyries; (b) magmatic fO_2 ($\log fO_2$) vs. magmatic temperature. MH: magnetite-hematite buffer curve; NNO: nickel-nickel oxide buffer curve; FMQ: fayalite-magnetite-quartz buffer curve; IW: iron-wustite.

5.3. Implications for Au Mineralization

Beiya is one of the largest Au deposits in China, with by-products Cu, Fe, Ag, Pb, and Zn [23–25]. Au and other ore metals were derived from the alkaline melt as documented by Yang et al. [47], and the

mineralized fluids at Beiya were mainly derived from magmatic water [16,17]. Those indicate that the Beiya alkaline melt provides ore-forming fluids and metallic elements for the formation of porphyry Au mineralization [16].

In contrast to Bailiancun porphyry, the Wandongshan porphyries are characterized by higher water content and higher fO_2 , but lower emplacement pressure (Table 2, Figure 7). Au solubility was investigated as a function of fO_2 in sulfur-saturated silicate melts [12], indicating Au solubility in sulfur-free silicate melt increases with increasing of fO_2 [48]. In addition, high fO_2 favors sulfur as sulfate dissolved in the evolved magmas and suppresses magmatic sulfide precipitation [4,7], which can cause sequestration of metals before they partition into the aqueous phase [10]. On the other hand, a high H_2O content and a low pressure results in the Wandongshan parental magma reaching water saturation easier, because water solubility increases with increasing pressure [44], hence ensuring exsolution of aqueous volatile phases and the partitioning of Au between sulfides and silicate melt efficiently [12,49,50]. Therefore, melts that are more enriched in water and with a high oxidation state will be more fertile to generate Au mineralization.

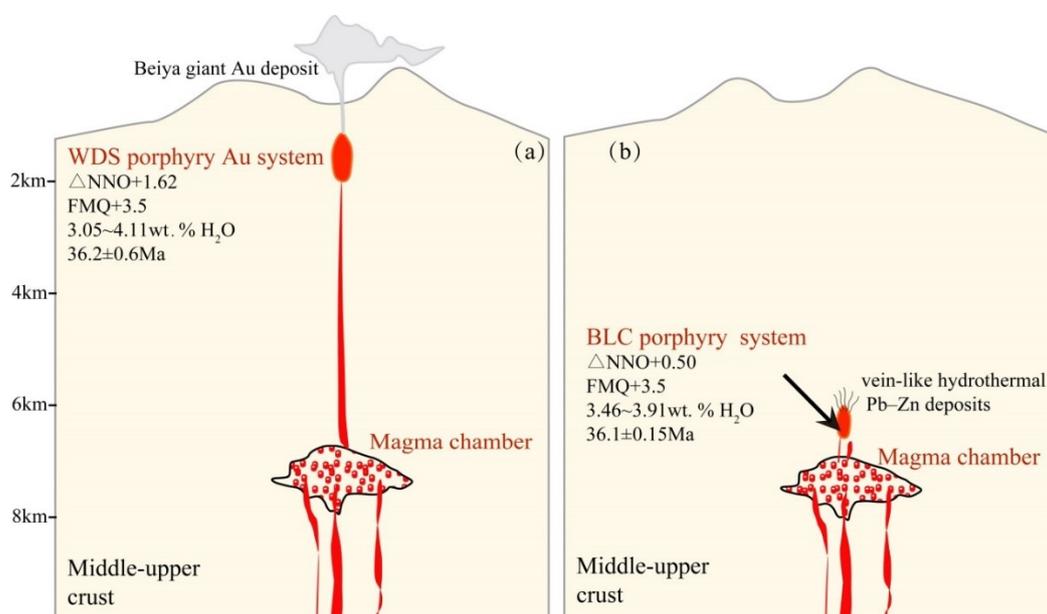


Figure 7. Cartoons illustrating the different water content, magmatic oxidation state, and emplacement depths of the parental magmas at Wandongshan (a) and Bailiancun (b), respectively. Previous studies show that the parent magma at Wandongshan (a) and Bailiancun (b) districts are derived from the same magma source [22]. This study reveals that the magma of Wandongshan experienced fluid saturation at about 1.19–2.68 km, while the magma of Wandongshan experienced fluid saturation at about 6.48–7.28 km. The water content and magmatic oxidation state of Wandongshan parental magmas are both higher than those of the Bailiancun parental magmas.

6. Conclusions

Chemical compositions of zircons and amphiboles are used to evaluate the physicochemical conditions (e.g., pressure, temperature, fO_2 , and water content) of the parental magma. We found the water content of the Wandongshan parent magma ($\leq 4.11 \pm 0.4$ wt %) are slightly higher than those of the parent magma at Bailiancun ($\leq 3.91 \pm 0.4$ wt %). Furthermore, the fO_2 deduced from amphibole chemistry of the Wandongshan porphyry range from $\Delta NNO + 1.00$ to $\Delta NNO + 1.89$, slightly higher than that of the Bailiancun porphyries ($\Delta NNO + 0.39$ to $\Delta NNO + 0.67$). This result is consistent with the values calculated by zircon chemistry. The Wandongshan porphyry has higher zircon Ce^{4+}/Ce^{3+} ratio (average of 62) and high fO_2 value (average of $\Delta FMQ = +3.5$) than those of Bailiancun porphyry (average of $Ce^{4+}/Ce^{3+} = 23.15$; $\Delta FMQ = +1.5$). Obviously, the Wandongshan ore-related porphyry has

higher water content and more oxidized than the Bailiancun barren porphyry. Therefore melts that are more enriched in water and with a high oxidation state will be more fertile to form an economic hydrothermal system especially when emplaced at shallow levels.

Supplementary Materials: The following are available online at <http://www.mdpi.com/2075-163X/8/10/441/s1>, Table S1: The compositions of amphiboles from the monzogranite porphyries at Wandongshan and Bailiancun, respectively, Table S2: The Zircon trace elements in the monzogranite porphyry from Wandongshan and Bailiancun districts.

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