

Article



# Reservoir Characteristics and Main Factors Controlling Carboniferous Volcanic Rocks in the Well CH471 Area of the Hongche Fault Zone: Northwest Margin of Junggar Basin, China

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Abstract: Nearly 100 million tons of reserves have been explored in the Well 471 area of the Hongche Fault zone. The Carboniferous volcanic rock reservoir is the main oil-bearing reservoir in the well CH471 area and is the main target of exploration and development. The characteristics of the Carboniferous volcanic rock reservoir are studied through core, thin section, physical property, logging, and other data, and its main controlling factors are analyzed in combination with actual means of production. The lithologies of the volcanic reservoir in the study area are mainly volcanic breccia, andesite, and basalt. The matrix physical properties of volcanic rock reservoirs are mediumporosity and ultralow-permeability, among which volcanic breccia has the best physical properties. The reservoir space mainly comprises primary pores, secondary dissolution pores, and fractures, resulting in a dual medium pore-fracture-type reservoir. Combined with production data analysis, the lateral distribution of oil and gas is controlled by lithology and lithofacies, with explosive volcanic breccia being the best, followed by the basalt and andesite of overflow facies, which are vertically affected by weathering and leaching and distributed within 50~300 m from the top of the Carboniferous system. The area with densely developed fractures was conducive to developing high-quality reservoirs. The tectonic movement promoted the formation of weathering and controlled the development of faults. Based on a comprehensive analysis, it is believed that the formation of Carboniferous volcanic oil and gas reservoirs in the study area was controlled and influenced by the lithology, lithofacies, weathering, leaching, faults (fractures), and tectonics.

**Keywords:** Hongche fault zone; Carboniferous; volcanic rock; reservoir characteristics; main controlling factor

# 1. Introduction

Volcanic oil and gas reservoirs have been explored for more than 130 years [1], and great achievements have been made worldwide thus far [2–5]. Volcanic reservoirs have become an important field of oil and gas exploration [6–9]. Volcanic oil and gas reservoirs were first discovered in the northwest margin of the Junggar Basin in China in the 1950s. To date, breakthroughs have been made in 14 petroliferous basins [10–14]. The Junggar Basin hosts one of the largest volcanic oil and gas reservoirs discovered [15]. Since the 1950s, several single wells in the Junggar Basin with Triassic and Jurassic as the target layers have encountered industrial oil flow in the Carboniferous strata, which has opened the prelude to the exploration of Carboniferous volcanic oil and gas in this area [16]. Volcanic



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**Copyright:** © 2023 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). oil and gas reservoirs have been found in Shixi, Kebai fault, Chepaizi, and other areas, with proven reserves of more than 100 million tons [17,18]. The Carboniferous strata of the Hongche fault zone in the northwest margin of the Junggar Basin contain one of the most important oil and gas enrichment zones in this area [19], and the object of this study is the Carboniferous volcanic reservoir of the CH471 well area in the middle part of the Hongche fault zone.

Reservoir identification and development is the main task of volcanic oil and gas exploration and development. The reservoir characteristics and main controlling factors are the foci of volcanic rock reservoir research [20,21]. Early studies on volcanic rock reservoirs mainly focused on the distribution of lithology and lithofacies and the influence of different lithologies and lithofacies on the relative reservoir space [22–27]. Some areas are limited by the strong heterogeneity of volcanic rocks, which makes their study difficult, and the effect on exploration and development is not ideal. In recent years, it has been gradually realized that the formation of volcanic rock reservoirs depends on late reconstruction, so the research direction has gradually changed to research on weathering, leaching, and the controlling range of tectonic fractures, and some achievements have been made [28–35].

Since the discovery of the Carboniferous volcanic rock oil and gas reservoir in the CH471 well area in the middle section of the Hongche Fault Zone in 1985, the exploration and development approach has been guided by the distribution of lithology and lithofacies. The particularities of volcanic reservoirs have not received sufficient attention, and systematic research on lithology, lithofacies, fractures, weathering, and other controlling factors is insufficient, which leads to difficulty in achieving the expected output. This study studied the characteristics and main factors controlling volcanic rock reservoirs using core and logging data combined with thin section observations and physical property experiments to provide a theoretical reference for reservoir prediction and efficient development in this area.

## 2. Geological Setting

The study area is located in the southern Chepaizi oilfield on the northwest margin of the Junggar Basin and is tectonically located in the upper disk-shaped segment of the Hongche fault belt on the northwest margin of the Junggar Basin [36,37] (Figure 1). The volcanic rocks in the study area were formed by the collision and closure of the Junggar Ocean Basin during the Carboniferous period under the influence of Hercynian movement and are typical continental island arc volcanic rocks at the active edge of converging plates [38–41]. After the formation of the Carboniferous volcanic mechanism in the study area, the area experienced the superimposed transformation of the Hercynian, Indosinian, Yanshan, and Himalayan tectonic movements, and the resulting internal structure is extremely complex [42–45]. At the end of the Hercynian movement, volcanic rocks in the study area experienced tectonic inversion from regional stretching to thrust nappe, and the Hongche fault zone began forming (Figure 2). The Indochinese period began to be transformed into a southwest–northeast oblique extrusion deformation, and the Hongche fault zone began to be transformed into a right-lateral strike-slip extrusion and continued to act in the Yanshanian period and Himalayan period [46,47].

Laterally, the Carboniferous system mainly contains two groups of nearly S-N and E-W reverse faults [48]. Among them, the north–south fault is the main fault in the study area, which was formed in the thrust-napping process of the middle Hercynian to late Hercynian. The Indosinian and Yanshan Epochs mainly inherited and formed synsedimentary faults based on the original structure. The overall formation time was early, and the faults cut many horizons over a large scale. The other group of east–west faults was formed in the middle to late Indosinian period. The right-lateral strike-slip fault occurred along the Hongche fault zone, and the thrust fault perpendicular to the main fault was formed under the action of extrusion. The main disconnected beds were Carboniferous beds, with faults with small fault spacing and late formation times that cut through the early north–south faults (Figure 1) [44]. Longitudinally, the Carboniferous

strata are dominated by igneous rocks, which are in unconformable contact with the lower Wuerhe Formation ( $P_2w$ ). Influenced by the thrust nappe of the Hongche fault zone (middle Hercynian to late Hercynian), the Carboniferous strata gradually uplifted from east to west and underwent denudation, and a weathering crust formed at the top boundary of the Carboniferous system.



Figure 1. Regional tectonic map and Carboniferous strata histogram of the study area.



Figure 2. Schematic diagram of E-W strike section in the study area.

## 3. Samples and Methods

In this paper, core observation and description are carried out in 17 wells with a total of 250 m and 33 samples. Preparation of 0.03 mm-thick cast thin sheets for more than 30 samples from the research area was performed, and microscopic studies of these thin sections were carried out using LB102 optical microscope produced by Laite Company. At the same time, 50 cast thin slices, 72 fluorescence thin slices, and 140 scanning electron microscopes were collected. Various types of pore spaces were evaluated through tissue observation and thin-section studies of resin-impregnated samples.

Petrophysical analysis was performed on more than 200 porosity samples from 17 wells in the CH471 area to measure the range of porosity and permeability. According to China National Petroleum Industry Standard (SY/T 5830-1993) [49] requirements, the porosity, and permeability were measured using a PoreMaster60 automatic mercury injection instrument. The analyzer measures the porosity of the rock by the mercury injection method. By pressurizing the mercury into the rock sample, the mercury volume change in the different hole grades can be measured when the pressure of the hole is continuously changed, and the porosity can be obtained by comparing the volume of the rock sample.

Many types of information on oil and gas exploration from Xinjiang Oilfield Company, PetroChina, including logging, mud logging, and oil and gas testing, were used to study and evaluate the characteristics of volcanic rock reservoirs.

# 4. Experimental Results

## 4.1. Volcanic Reservoir Characteristics of Lithology and Lithofacies

Through observations of more than 200 m of Carboniferous volcanic rock cores and identification of corresponding rock thin sections, combined with a comparative analysis of logging data, the volcanic facies in the study area are mainly divided into volcanic sedimentary facies, overflow facies, and explosive facies. The lithology mainly consists of tuffaceous sandstone, tuff, andesite, basalt, and volcanic breccia. Laterally, the Carboniferous volcanic rocks are distributed in a southwest–northeast strip shape with good continuity and rapidly pinch out from east to west (Figure 1).

The tuff in the study area is mainly cemented by volcanic debris. The interior is mostly composed of debris, crystal pyroclasts, glass shards, flame, and volcanic ash, and the appearance is mostly dark gray or gray-black (Figure 3a). The material in tuffaceous sandstone volcanic rock is lower than that of tuff, and the material is mixed with normal

sediments, such as sediment and gravel. Under the microscope, the particle size of the tuff debris is less than 2 mm, and the particles have a tuff structure (Figure 3b). Andesite is a neutral extrusive rock; fresh surfaces are brownish gray; after weathering or alteration, it is generally purple-red, and the cracks in the core are filled by white material (Figure 3c). The andesite is mainly composed of medium feldspar and labradorite, with some dark minerals, and the matrix shows an interweaving structure (Figure 3d). Basalt is a basic extrusive rock. The rock core color is usually dark black or dark green, with pores and almond-shaped pores. The pores are mostly filled with chlorite and calcite in the later stage, affected by weathering or alteration (Figure 3e). Under the microscope, the basalt is composed of basic plagioclase and pyroxene, and it has a porphyritic structure (Figure 3f,g). The volcanic breccia in the area is mainly basaltic andesitic volcanic breccia. The fresh surface of the core is mostly grayish black, with a volcanic breccia structure and block structure. The breccia content is more than 50%, the particle size is more than 1 cm, and some breccia surface pores are developed (Figure 3h). Under the microscope, the matrix is mainly strip-shaped plagioclase, with a small amount of chlorite and pyroxene and a porphyritic structure (Figure 3i).



**Figure 3.** Lithological characteristics of Carboniferous volcanic rocks. (**a**) Tuff, CH545, 2097.84–2098.3 m; (**b**) Tuff structure and microfractures, CH545, 2097.84 m; (**c**) Andesite, CH47, 2770.97–2771.15 m; (**d**) Andesite with interwoven structure with visible medium feldspar, CH47, 2770.97 m. (**e**) Basalt with visible stomatal development, CH471, 2630.59–2630.72 m. (**f**) Half-filling fractures and oil spots can be seen in basalt, CH471, 2630.59 m. (**g**) Basalt, Porphyritic structure, CH471, 2633.75 m. (**h**) Some breccia surface pores are developed, CH472, 2627.35–2627.69 m. (**i**) The basaltic andesitic breccia has breccia and interweaving structures, CH472, 2627.35 m.

### 4.2. Volcanic Reservoir Pore Types and Characteristics

Through core observations, thin section identification, and scanning electron microscope observation and analysis, it is suggested that the reservoir space of the Carboniferous volcanic reservoir in the area can be divided mainly into two categories, namely, pore type and fracture type, which can be further subdivided into four categories: primary pore, secondary pore, primary fracture, and secondary fracture [50–55].

The main pores of the reservoir in the study area are mainly primary pores (Figure 4a). The upwelling and decompression of molten magma causes a "degassing" effect of volatile components, such as water or carbon dioxide, which erupts to the surface and cools quickly and then is fixed in the rocks to form pores; these pores are developed mostly in the upper parts of lava, such as basalt and andesite [56,57]. The observations under the microscope showed that the stomatal distribution is uneven, the pore shape is mainly round, and some pores have irregular shapes with different sizes. The larger pores are visible to the naked eye at approximately 1 cm, and the smaller pores need to be observed under the microscope. Some pores are filled with calcite, chlorite, and other minerals, forming semifilled pores or amygdules, resulting in poor reservoir storage. The primary pores that developed in the volcanic rocks are mostly isolated, and their connectivity is poor. In the later stage, the transformation of tectonism and weathering dissolution is needed to form effective reservoir space. Secondary pores are dominated by various dissolution pores, mainly pores formed by the complete or partial dissolution of the volcanic rock matrix or soluble minerals in the later stage. The phenocryst dissolved pores, intragranular dissolved pores, and sieve pores are pores formed via hydrothermal dissolution in the later soluble mineral stage (Figure 4b–d). The volcanic matrix's dissolution forms the Cave pores (Figure 4e). Moldic pores (Figure 4f) are formed through mineral shedding [58].



**Figure 4.** Pore types of volcanic reservoirs. (a) Gas pore of basalt, CH541, 2079.05 m. (b) Porphyritic dissolved pore, CH541, 2055.89 m. (c) Dissolved pores in grains, CHF061, 2546.55 m. (d) Cribriform pore, CH541, 2055.89 m. (e) Cave pore, CH542, 2090.86 m. (f) Moldic pore, CH541, 2057.31 m.

The proportion of primary fractures in the study area is small, being mainly condensation shrinkage fractures [59]. As a typical rock-forming fracture formed in the process of volcanic rock diagenesis, there are two ways to form condensation shrinkage fractures. One method of formation is the process of magma condensation and crystallization, which forms inside clastic particles (Figure 5a). The other formation method is via the tensile stress caused by the cooling and shrinkage of volcanic rocks, which makes the surface of the rock mass rupture and form cold shrinkage joints (Figure 5b). The secondary fractures mainly include structural fractures and dissolution fractures. Structural fractures are caused by tectonic stress, which has the characteristics of long extension, directionality, and regularity (Figure 5c). Some structural fractures are affected by breccia and develop along the edge of breccia to form intergravel fractures (Figure 5d). A dissolution fracture is based on an original fracture or joint, which is dissolved by dissolution fluid, and dissolution expansion occurs at the edge of the fracture (Figure 5e,f).



**Figure 5.** Fracture types of volcanic reservoir. (a) Microfracture and condensation shrinkage fractures, CH30, 2287.08 m. (b) Condensation shrinkage fractures and tension fractures, CH471, 2625.09–2625.38 m. (c) Structural fractures, CH472, 2956.78 m. (d) Intergravel fractures, CH48, 3165.22–3165.41 m. (e) Dissolution fractures, CH48, 2726.67–2726.97 m. (f) Half-filled dissolution fractures, CHF6, 2506.11 m.

#### 4.3. Reservoir Physical Properties

The experimental results of the porosity and permeability analysis of Carboniferous volcanic rock core samples in the study area show that (Figure 6) the porosity and permeability of volcanic breccia are generally high, the porosity values are between 7.54% and 21.1%, with an average of 14.87%, and the permeability values are between 0.01 mD and 2.05 mD, with an average of 0.4 mD; andesite has the second highest porosity and permeability, with porosity values of 1.2%~18.6% and an average of 10.65%, and permeability values of 0.01 mD~1.06 mD, with an average of 0.24 mD; basalt is the densest, with porosity values of 1.2%~17.2%, and a mean value of 6.1%, permeability values of 0.01 mD~0.612 mD, and a mean value of 0.06 mD.

The overall porosity values of volcanic rocks in the study area are between 1.2% and 21.1%, with an average of 11.08%, and the average permeability values are less than 0.3 mD, which indicates medium-porosity and ultralow-permeability rocks. Because the study area is located on the hanging wall of the Hongche fault zone, the regional structure is complex, the faults and cracks are extremely developed, and the volcanic rock itself is highly heterogeneous, experiments on the porosity and permeability cannot show the improvement in the physical properties of the volcanic rock matrix caused by large cracks. The porosity and permeability of the layers with microcracks, pores, and dissolved pores are much higher than those of the undeveloped layers, which causes the overall experimental results to have a poor correlation with the real porosity and permeability, and the actual porosity and permeability are higher [60–62]. The experimental results show that without



considering other factors, the volcanic breccia matrix has the best physical properties, followed by andesite, and basalt is the densest.

Figure 6. Volcanic rock porosity and permeability distribution histograms.

# 5. Discussion

Volcanic rocks have strong heterogeneity, and it is impossible to track reservoirs simply by the distribution of the dominant lithology. As the basis of reservoir development, the dominant lithology is not a decisive factor [63,64]. The main body of the study area is located on the hanging wall of the Hongche fault zone. After forming the Carboniferous volcanic rocks, many cracks are generated inside the rock mass under the influence of thrust nappe. At the same time, they are uplifted and subjected to weathering and denudation so that the isolated pores inside the volcanic rocks can be connected, and finally, a high-quality reservoir is formed. The above comprehensive analysis considers the lithology the basis for developing volcanic reservoirs in the area. The later transformation of fractures, weathering and leaching is the key to reservoir formation. Tectonic action controls the development of faults and promotes the occurrence of weathering. The above points are the main factors controlling reservoir development. Based on the above control factors, the control influence range and degree of each factor on the reservoir can be divided, respectively, and the advantage area can be superimposed to track the favorable reservoir.

#### 5.1. Lithology Is the Foundation of Reservoirs

Physical properties experiments show that the volcanic rock matrix has medium-to low-porosity and ultralow-permeability physical properties. Some pores are present in the matrix, but there is a lack of connectivity. The formation of reservoirs is more dependent on later fracturing, weathering, and leaching, but the matrix pores of different lithologies are different, and the later transformation is different and difficult [65]. Among them, the volcanic breccia matrix of the explosive facies has the best physical properties, and internal pores, such as gravel fissures, are conducive to transformation via later weathering and dissolution and fracture development. The physical properties of andesite and basalt are relatively poor, and basalt is denser, but as a basic magmatic rock, basalt is more susceptible to weathering and dissolution. The acidic andesite has a strong anti-weathering ability, and the effect on later transformation is poor. Therefore, volcanic breccia is the best lithology for reservoirs, followed by basalt, and andesite is a poor reservoir lithology.

According to the statistics of 13 single wells with a cumulative production of more than 3000 tons and oil test data in the study area, the cumulative total reservoir thickness is 586 m (Table 1): the thickness of the volcanic breccia reservoir accounts for 63%, that of andesite accounts for 9.5%, and that of basalt accounts for 27.5%; the oil test production values are 192 t/d, 27 t/d, and 103 t/d, respectively; and the statistical results are consistent with the research findings; that is, lithology cannot be used as a decisive factor, but the matrix porosity and permeability are better, and the lithology with poor weathering resistance is more likely to become a reservoir.

| Well  | Volcanic Breccia<br>Reservoir (m) | Volcanic Breccia<br>Oil Test<br>(t/Day) | Andesite Reservoir (m) | Andesite Oil Test<br>(t/Day) | Basalt Reservoir (m) | Basalt Oil Test<br>(t/Day) |
|-------|-----------------------------------|---|------------------------|------------------------------|----------------------|----------------------------|
| CH471 | 0                                 | 0                                       | 10                     | 4.45                         | 24                   | 10.69                      |
| CH472 | 68                                | 31.45                                   | 0                      | 0                            | 14                   | 4.01                       |
| CH476 | 23                                | 10                                      | 0                      | 0                            | 0                    | 0                          |
| CH480 | 0                                 | 0                                       | 4.5                    | 2.92                         | 28.5                 | 18.21                      |
| CH481 | 47                                | 38.47                                   | 0                      | 0                            | 0                    | 0                          |
| CH482 | 0                                 | 0                                       | 0                      | 0                            | 29                   | 26.26                      |
| CH488 | 0                                 | 0                                       | 0                      | 0                            | 30                   | 18.71                      |
| CH489 | 64                                | 14.72                                   | 0                      | 0                            | 0                    | 0                          |
| CH47  | 39                                | 25.94                                   | 42.1                   | 20.075                       | 0                    | 0                          |
| CH50  | 18                                | 13.86                                   | 0                      | 0                            | 0                    | 0                          |
| CHF6  | 38                                | 21.09                                   | 0                      | 0                            | 14                   | 3.89                       |
| CH303 | 64                                | 18.66                                   | 0                      | 0                            | 11                   | 17.58                      |
| CH304 | 27                                | 18.02                                   | 0                      | 0                            | 9                    | 4.3                        |
| Sum   | 388                               | 192.21                                  | 56.6                   | 27.445                       | 159.5                | 103.65                     |

According to the results of a single-well volcanic reservoir, the lateral distribution of Carboniferous volcanic reservoirs with different lithologies and lithofacies is reconstructed and analyzed by using Gxplorer software. The distribution of volcanic breccia reservoirs with explosive facies is relatively concentrated, and the thickness is large. This type is mainly located in the eastern part of the study area and is distributed in a north—south strip (Figure 7a). This feature occurs because the explosive facies is mainly developed near the crater. The basalt and andesite reservoirs with the overflow facies are widely distributed laterally, have small thicknesses, are mainly located in the central and western parts of the study area, and are distributed in a north—south strip (Figure 7b).



**Figure 7.** The distribution map of reservoir plane thickness in the study area. (**a**) Plane thickness distribution diagram of explosive facies volcanic breccia reservoir; (**b**) plane thickness distribution diagram of overflow facies basalt and andesite reservoir.

## 5.2. Fracturing, Weathering, and Leaching Are the Key to Reservoir Formation

The primary pores in the early stage of volcanic rock formation are often isolated and disconnected from each other. The reservoir needs to be reformed through fracturing, weathering, and leaching to improve the connectivity and increase the reservoir space before it can become an effective reservoir.

## 5.2.1. Effect of Fracturing on the Reservoir

The development of fractures is very important for forming volcanic reservoirs [51,66]. The development of fractures not only improves the connectivity and increases the reservoir space but also provides a channel for hydrothermal infiltration of later weathering and leaching, accelerating the weathering of volcanic rocks and forming effective reservoirs. The study area is located on the front edge of the hanging wall of the Hongche fault zone. In the Hercynian period, multiple faults were formed under the influence of thrust nappe tectonic activity, which caused extreme development of the cracks in the Carboniferous volcanic rocks in the area and provided conditions for the formation of reservoirs. In this paper, the influence of fractures on reservoirs is further verified by core observations and imaging logging combined with production data.

A total of more than 200 m of statistics and observations of volcanic rock cores in the study area show that the filling degree of core fractures in the oil layer is low, with unfilled fractures accounting for more than 60%, semifilled fractures accounting for 27%, and fully filled fractures accounting for only 11%. The filling degree of the fractures in the core in the non-oil layer is high, the sum of fully filled and half-filled fractures accounts for nearly 95%, and the unfilled fractures account for only approximately 5% (Figure 8). The effectiveness of the fractures in the volcanic reservoir section is significantly higher than that in the non-oil section.



**Figure 8.** Core fracture filling degree distribution histograms. (a) Histogram of distribution of oil layer core filling degree; (b) histogram of distribution of non-oil layer core filling degree.

In the imaging logging data comparison, three high-production Wells CH476, CH481, and CH480 are taken as examples. The first is CH476: the depths of the perforated section are 2558~2561 m, the lithology is volcanic breccia, the fracture density is 2.6 bar/m, the permeability is 4.72 mD, and the oil test output is 1.6 t/d. The depths of the perforated section are 2567~2570 m, the lithology is volcanic breccia, the fracture density is 7 bar/m, the permeability is 13.97 mD, and the test oil production is 6.9 t/d. The second is CH481: the depths of the perforated section are 2530~2533 m, the lithology is volcanic breccia, the fracture density is 1.3 bar/m, the permeability is 1.46 mD, and the oil test output is 14.2 t/d. The depths of the perforated section are 2668~2671 m, the lithology is volcanic breccia, the fracture density is 2.6 bar/m, the permeability is 8.11 mD, and the test oil production is 24.27 t/d. The second is CH480: the depths of the perforated section are  $2569 \sim 2572$  m, the lithology is basalt, the fracture density is 1 bar/m, the permeability is 2.72 mD, and the oil test output is 0.49 t/d. The depths of the perforated section are  $2869 \sim 2872 \text{ m}$ , the lithology is basalt, the fracture density is 3 bar/m, the permeability is 15.25 mD, and the test oil production is 21.13 t/d (Figure 9). In summary, it is concluded that the test oil production is relatively high in the layers with high fracture density.



Figure 9. CH476, CH481, and CH480 imaging logging figure.

According to the analysis of the above research results, based on the same lithology in the same single well, the high fracture development density can effectively improve the reservoir permeability and obtain relatively high oil test data, the density and effectiveness of fractures are key factors controlling reservoir development, and the area with dense development of effective fractures is conducive to the development of high-quality reservoirs.

## 5.2.2. Effect of Weathering and Leaching on Reservoirs

The influence of weathering and leaching on volcanic reservoirs is similar to that of fractures, which mainly improves connectivity and increases reservoir space. The difference is that the influence of weathering and leaching on volcanic reservoirs is vertical and regular [67,68]. The study area is located in the middle section of the hanging wall of the Hongche fault zone. The Carboniferous volcanic rocks in the area were affected by the late Hercynian thrust nappe after formation. After uplift, they were exposed at the surface and affected by weathering and leaching for a long time, forming an ancient weathering crust at the top of the Carboniferous strata.

According to core observations and statistical work, the volcanic rocks have obvious weathering and leaching in the depth range of 200–300 m from the top boundary of the Carboniferous system. This depth difference should be due to the different weathering resistances of different lithologies. Taking well CH48 as an example, at the top boundary of the Carboniferous strata (2631.10 m), the core is the most affected by weathering, and some parts of the core have been broken. However, with increasing distance from the top boundary, the degree of weathering of the core gradually decreases. The core above 100 m (2726.67 m) from the top boundary is still affected by weathering to a certain extent, with obvious development of dissolution pores and fractures. Until the depth reaches more than 250 m (2885.71 m), the core has no obvious weathering (Figure 10).

Based on the above understanding, a map of the longitudinal distribution of the weathering degree in the study area was established (Figure 11). The range within 50 m from the top boundary of the Carboniferous system is part of the weathered clay layer. In this range, the reservoir is blocked by clayization due to excessive weathering. The range of 50~200 m from the top boundary of the Carboniferous strata is a strong weathering zone. The closer to the top of this range, the stronger the weathering is, and the more favorable the formation of high-quality reservoirs. In the range of 200~300 m, the strata are part of the weak weathering zone, the reservoir development is relatively poor, and the reservoir physical properties are general. The unweathered zone greater than 300 m basically does not contain reservoirs.



Figure 10. Core photo of well CH48.



Figure 11. The longitudinal distribution map of weathering degree in the study area.

To further verify the influence of weathering and leaching on the reservoir, the distance between the oil-bearing intervals of 12 wells in the study area and the top boundary of the Carboniferous system is statistically analyzed. The results are as follows (Figure 12). The oil-bearing layers in the study area are mainly concentrated in the range of 50~300 m from the top boundary of the Carboniferous strata, and the cumulative production of some single wells can reach more than 20,000 tons. The analysis shows that the oil-bearing layers are too close to the top surface of the weathering crust (less than 50 m), and the weathering effect is too strong, which causes the layers to be easily blocked by clay and have difficultly forming a reservoir. However, far away (more than 300 m), the weathering effect is weak, the transformation effect is poor, and it is also difficult to form high-quality reservoirs. In the range of 50~300 m, the weathering and leaching effect is not strong enough to destroy the reservoir but can also complete the transformation of the reservoir, thus forming a high-quality reservoir. The strong weathering zone within 50~200 m is the optimal reservoir interval, and the closer to the top, the easier it is to form high-quality reservoirs.



Figure 12. The depth of oil-bearing interval from the top boundary of Carboniferous in the study area.

## 5.3. Tectonic Movement Is the Basis of Volcanic Reservoir Formation

Volcanic reservoirs are different from sedimentary reservoirs. Their formation is inseparable from the influence of tectonic movement. In the long period of geological history after its formation, volcanic rocks experience multiple stages and various forms of superposition due to tectonic movement, which promotes the formation of a large number of structural fractures in volcanic rocks and increases connectivity. The high part of the volcanic rock mass is uplifted via tectonic movement and subjected to weathering and denudation to form a weathering crust, forming many dissolution pores and dissolution fractures in the volcanic rock, further improving the connectivity and increasing the reservoir space.

The Carboniferous volcanic rocks in the study area are located on the hanging wall of the Hongche fault zone. Since their formation, they have been superimposed by a series of tectonic movements, such as the Hercynian, Indosinian, Yanshanian, and Himalayan movements. It is mainly divided into the following three stages: (1) The volcanic rocks in the late Hercynian movement were uplifted and subjected to weathering and denudation under the action of thrust and napper. A large number of fractures occur in the rock mass during the tectonic inversion from tension to extrusion. (2) Under the influence of the Indosinian tectonic movement, the Hongche fault began to have right-lateral strike-slip extrusion movement, accompanied by a large number of fractures, while the weathering and denudation weakened and received deposition again. (3) The Yanshan Stage and Himalayan Stage inherited the original tectonic activity, but the intensity gradually weakened.

The influence of tectonic movement on the reservoir in the area mainly manifests in the following two points. The first point is that the superposition of multistage tectonic movements causes a large number of structural fractures to be formed inside the tight volcanic rocks, which increases the connectivity and provides a channel for the infiltration of the dissolution fluid to promote dissolution. The second point is that the rock mass is uplifted and subjected to weathering and denudation under the influence of the thrust nappe structure, and the weathering crust forms on top of the volcanic rock, which greatly improves the reservoir physical properties of the volcanic rock within a certain range from the weathering crust.

## 6. Conclusions

(1) The volcanic reservoir lithology in the study area mainly consists of volcanic breccia, andesite, and basalt. Among them, the volcanic breccia matrix has the best physical properties and the largest reservoir thickness, and it is the dominant lithology of reservoir development, followed by basalt. Andesite is the least dominant lithology.

(2) The reservoir space of the Carboniferous volcanic reservoir in the study area is mainly composed of primary pores, secondary structural fractures, and dissolution pores. The primary pores depend on the later transformation of secondary pores to form effective reservoirs.

(3) Due to the uplift of the Carboniferous system in the study area by a thrust nappe, the surface is exposed to weathering and denudation, and the weathering crust is formed at the top. The top weathering crust of the Carboniferous strata is divided into a clay layer, strong weathering layer, weak weathering layer, and unweathered area from top to bottom. The reservoirs with the best physical properties are mainly found in the interval of 50~300 m from the top of the Carboniferous strata.

(4) The formation of Carboniferous volcanic reservoirs in the CH471 well area is controlled by lithological facies, weathering and leaching, fractures, and tectonic action. Vertically, the reservoir is controlled by weathering and leaching within 400 m from the top boundary of the Carboniferous. Laterally, the volcanic reservoir is controlled by the distribution of the dominant lithology and the area of dense fracture development. The fractures mainly develop near the fault zone. Tectonic movement promotes the development of faults and the formation of weathering and leaching.

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