





## Article

# Evaluation of the Gulf of Aqaba Coastal Water, Jordan

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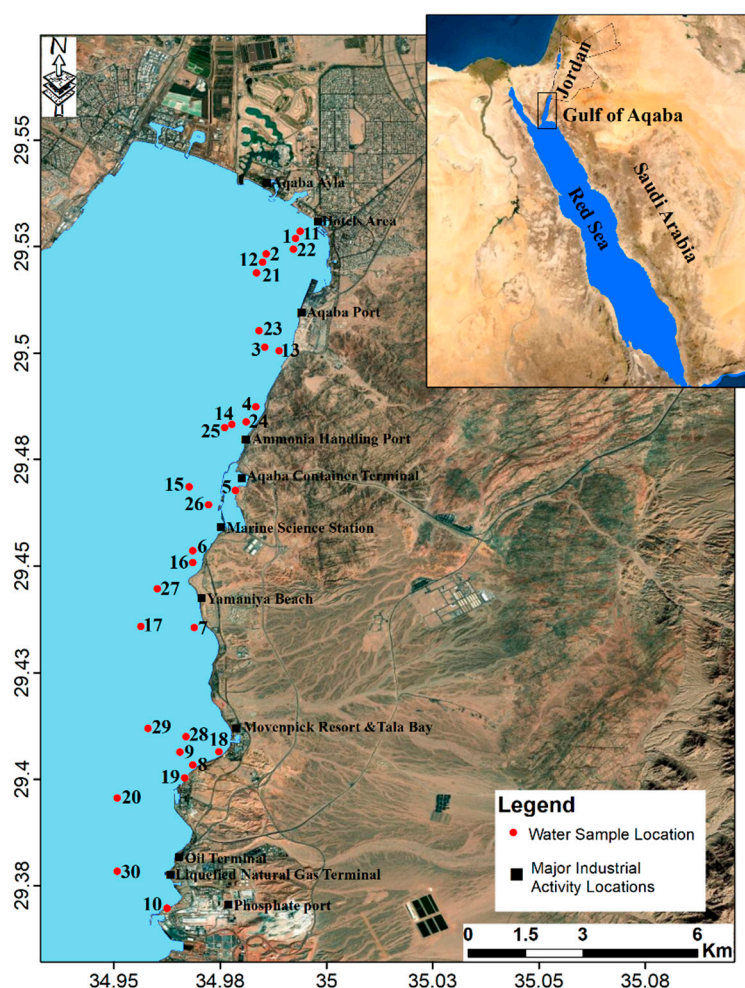


**Abstract:** (1) Background: The Gulf of Aqaba (GoA) supports unique and diverse marine ecosystems. It is one of the highest anthropogenically impacted coasts in the Middle East region, where rapid human activities are likely to degrade these naturally diverse but stressed ecosystems. (2) Methods: Various water quality parameters were measured to assess the current status and conditions of GoA seawater including pH, total dissolved solids (TDS), total alkalinity (TA),  $\text{Cl}^-$ ,  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$ ,  $\text{PO}_4^{3-}$ ,  $\text{NH}_4^+$ ,  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Na}^+$ ,  $\text{K}^+$ , Sr, Cd, Co, Cr, Cu, Fe, Mn, Pb, and Zn. (3) Results: The pH values indicated basic coastal waters. The elevated levels of TDS with an average of about 42 g/L indicated highly saline conditions. Relatively low levels of inorganic nutrients were observed consistent with the prevalence of oligotrophic conditions in GoA seawater. The concentrations of  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Na}^+$ ,  $\text{K}^+$ , Sr,  $\text{Cl}^-$ , and  $\text{SO}_4^{2-}$  in surface layer varied spatially from about 423–487, 2246–2356, 9542–12,647, 513–713, 9.2–10.4, 22,173–25,992, and 317–407 mg/L, respectively. The average levels of Cd, Co, Cr, Cu, Fe, Mn, Pb and Zn ranged from 0.51, 0.38, 1.44, 1.29, 0.88, 0.38, and 6.05  $\mu\text{g/L}$ , respectively. (4) Conclusions: The prevailing saline conditions of high temperatures, high evaporation rates, the water stratification and intense dust storms are major contributing factors to the observed seawater chemistry. The surface distribution of water quality variables showed spatial variations with no specific patterns, except for metal contents which exhibited southward increasing trends, closed to the industrial complex. The vast majority of these quality parameters showed relatively higher values compared to those of other regions.

**Keywords:** water quality; coastal area; metals; pollution source; Gulf of Aqaba; Jordan; Red Sea

## 1. Introduction

The Gulf of Aqaba (GoA) is the upper northeastern segment of the Red Sea [1]. It is a partially-isolated, narrow and deep coastal water body. The Strait of Tiran connects GoA with the Red Sea (Figure 1). Despite the extreme environmental conditions, the GoA supports unique aquatic ecosystems and biodiversity, and is a habitat for one of the world's richest coral communities [2,3].



**Figure 1.** Location map of the Gulf of Aqaba (GoA) and sampling sites.

The GoA is one of the high anthropogenically impacted coasts in the Middle East region [4]. The expansion in economic and industrial activities in the Gulf's bordering countries have contributed to the degradation of naturally stressed coastal and marine ecosystems. They are subjected to various impacts and sources of pollution including dredging and reclamation activities, coastal construction development, industrial waste, ports, oil spills, and domestic sewage, among others [5].

GoA is the only marine port for Jordan, and is highly vulnerable to pollution, where all marine-related activities are concentrated within a few kilometers of the coast (27 km). In addition, many economic, industrial, and recreational activities are taking place along the Jordanian coastline, many of which are of potential environmental impacts [6–9]. Additionally, the region plans to have a number of large coastal projects (such as the Red-Dead Sea conduit, new resorts, and ports relocation), which will certainly accelerate the degradation cycle of existing environmental conditions and threaten these unique marine communities [7]. Signs of human impacts were reported [10–12].

In addition to human impacts, the GoA is subject to regular dust storm events that contribute metals and other chemicals to the GoA coastal water [8,9]. Aeolian dust flux to GoA is likely to influence seawater chemistry [13], where atmospheric dry deposition in the GoA is considered an important external source of trace metals [8,9,14–16]. The mineral dust rate on GoA region is one of the highest on Earth [8,9,13,17]. It is believed that the frequency of dust storm events will become more common in the GoA, due to increase in regional aridity and dust fluxes [18,19].

The relatively small volume and absence of significant wave action along with the low rate of water circulation and renewal (between GoA and the Red Sea), render the Gulf particularly vulnerable to pollution. The residence time of water in the Gulf averaged 1–3 years [20,21].

The growing concern over the sustainability of these unique aquatic ecosystems of GoA has recently gained momentum and became a priority issue in Jordan. The impact of intense and widespread human activities from the Gulf's bordering countries poses imminent threats to GoA coast, which requires a proper monitoring plan. The objective of the present study is to assess the current status of surface water quality along the coastal region of the Jordanian GoA coast. It also intends to evaluate the spatial distribution of a variety heavy metals and to identify potential sources of contamination. This assessment will help develop a sustainable management plan for coastal water resources.

## 2. Materials and Methods

### 2.1. Description of Study Area

The GoA is the Red Sea's northeastern extension. It is a partially-enclosed, narrow and deep coastal water body. The Strait of Tiran connects the Red Sea to the GoA (Figure 1). The GoA extends approximately 180 km southward with a width ranging from 5–25 km (the average of 16 km maximum), and a maximum depth of about 1800 m (the average is about 800 m). Only 27 km of the eastern coast belong in Jordan, and the remaining coastline, unpopulated, and largely underdeveloped, lie in the Saudi territory.

The GoA is influenced by prevailing subtropical conditions with extremely high temperatures, high evaporation rate (about 400 cm/year) and negligible rainfall (of less than 2.2 cm/year) [22]. Surface water flow in the Gulf is nonexistent or limited solely during rare intense rainstorms occurring as flash floods in winter. The average water temperature in the upper 200 m varies seasonally from 20 °C in winter to 28 °C in summer, whereas the average air temperature ranges between  $32.20 \pm 3.16$  °C in summer and  $17.60 \pm 3.46$  °C in winter [23]. The maximum sea level of 154.30 cm was recorded in 2013 [23].

These conditions result in a high salinity in surface water layer, ranging from 40.3 to 40.8‰ in winter and from 40.5 to 46.6‰ in summer [9,24–26]. The surface coastal water of the GoA is extremely oligotrophic, because of its nutrient-poor water originating from the Red Sea surface waters through the Straits of Tiran. The surface water is shallow with stable thermocline throughout the year, except in wintertime, when a wind-driven convective mixing occurs between the deep (nutrient-rich) and surface waters. Water stratification occurs in spring. However, the oceanographic characteristics of extensive solar irradiance, high transparency, deep sunlight penetration, and warm water created unique aquatic ecosystems and biodiversity, with one of the world's richest coral communities [2,3].

The northerly wind, with a high speed and activity during summertime, is the prevailing wind direction and is responsible for the majority of aeolian deposition events in the region. However, Khamaseen winds blowing in springtime account for most sand and dust storms in southern Jordan and the adjacent areas [27]. They deliver dust from the interior of the Sahara Desert in north Africa.

### 2.2. Sampling and Analysis

Surface water samples were collected in September 2017 from 30 different locations along the coastal areas of GoA, Jordan (from north to south), sampling sites are presented in Figure 1. Coastal water samples were collected in 1-L precleaned polyethylene containers pre-rinsed with 10% HCl and 2 mL of HNO<sub>3</sub>. Samples were labeled and measured in the field for pH, electrical conductivity (EC, mS/cm at 25 °C), and total dissolved solids (TDS) using pH-meter (Sensions 5, HACH portable case), and EC/TDS-meter (ECOSCAN-hand held series, EUTECH instruments). Water samples were kept refrigerated at 4 °C and transported to a water laboratory (Yarmouk University, Jordan) for subsequent chemical analyses. Sample preparation and analysis followed APHA [28] procedures.

In the laboratory, all samples were filtered by Whatman filter paper (No. 42) and analyzed for total alkalinity (TA), Cl<sup>−</sup>, NO<sub>3</sub><sup>−</sup>, SO<sub>4</sub><sup>2−</sup>, PO<sub>4</sub><sup>3−</sup>, NH<sub>4</sub><sup>+</sup>, Ca<sup>2+</sup>, Mg<sup>2+</sup>, Na<sup>+</sup>, K<sup>+</sup>, Cd, Co, Cr, Cu, Fe, Mn, Pb, Sr, and Zn, as follows: 50 mL of filtered samples were used to determine the concentrations of Na<sup>+</sup>, K<sup>+</sup>,

$\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ , and Cd, Co, Cr, Fe, Mn, Pb, Sr, Zn using flame atomic absorption spectrophotometer FAAS (NOVAA 300 Analytica JENA AJ with detection limits varying from 0.001–0.02  $\mu\text{g/L}$ ). Each sample was analyzed in duplicate. The accuracy and precision of the analytical method was evaluated by the analysis of a reference material (NASS-5), with recoveries ranging between 98.02–104.01%. A total of 5 mL of filtered samples was used to measure  $\text{NH}_4^+$ ,  $\text{Cl}^-$ ,  $\text{NO}_3^-$ ,  $\text{SO}_4^{2-}$ ,  $\text{PO}_4^{3-}$  by ion chromatography (IC) (Dionex ICS 1600, Thermoscientific). A total of 25 mL of filtered samples were titrated with 0.02 N  $\text{H}_2\text{SO}_4$ , using phenolphthalein and methyl orange as indicators to determine total alkalinity (TA) of the water samples. There are two pH endpoints corresponding to the above indicators at 8.3 and at 4.3. TA was calculated using the following equation:

$$\text{TA} = (\text{volume of acid used} * \text{normality of acid} * 50,000) / \text{volume of sample}.$$

The average ionic mass balance for water quality data was  $-0.02\%$  indicating a high level of accuracy.

### 3. Results and Discussion

The results of dissolved metals and physicochemical properties of coastal water are tabulated in Table 1, and presented in Figures 2–5. The pH values of surface water layer ranged between 8 and 8.49, with a mean value of 8.26 (Table 1 and Figure 2). They indicate a slightly basic coastal water. They showed a slight spatial variability with no distinct trends. This is likely related to the calcium carbonate buffering capacity of water [23].

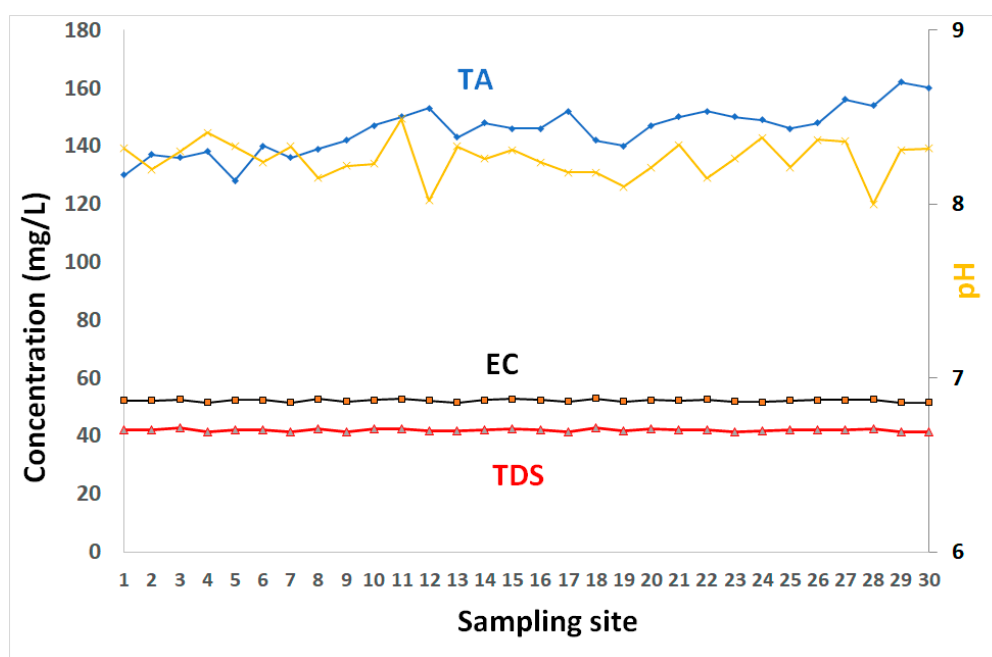


Figure 2. The electrical conductivity (EC), total dissolved solids (TDS), and pH of GoA coastal water.

**Table 1.** Characteristics of surface seawater layer, GoA, Jordan.

S.N.	pH	EC	TDS	TA	Cl <sup>−</sup>	NO <sub>3</sub> <sup>−</sup>	SO <sub>4</sub> <sup>2−</sup>	PO <sub>4</sub> <sup>3−</sup>	NH <sub>4</sub> <sup>+</sup>	Ca <sup>2+</sup>	Mg <sup>2+</sup>	Na <sup>+</sup>	K <sup>+</sup>	Sr	Cd	Co	Cr	Cu	Fe	Mn	Pb	Zn
		mS/cm	g/L																			
									mg/L										µg/L			
1	8.32	52.27	41.97	130	24,592.68	15.29	382.26	0.27	16.57	442.45	2295.70	11,165.80	591.78	9.88	0.20	0.38	1.01	0.91	0.94	0.76	0.19	3.93
2	8.20	52.12	41.89	137	24,992.25	13.80	382.12	0.27	15.19	443.91	2283.90	10,759.90	587.75	9.98	0.28	0.31	0.97	0.76	1.23	0.74	0.23	5.10
3	8.30	52.48	42.74	136	25,392.13	14.26	381.29	0.28	15.35	486.99	2355.90	11,736.90	602.31	10.40	0.40	0.31	0.96	1.18	0.75	0.76	0.26	5.23
4	8.41	51.56	41.34	138	23,72.94	13.50	407.45	0.26	14.15	433.9	2253.90	9891.20	547.82	9.44	0.49	0.33	1.32	0.87	1.40	0.70	0.17	4.45
5	8.33	52.40	42.10	128	25,632.05	12.65	396.85	0.29	13.08	444.76	2312.40	11,571.90	579.97	9.93	0.42	0.37	1.38	0.69	1.29	0.70	0.22	5.70
6	8.24	52.41	42.15	140	25,192.19	14.01	401.12	0.26	15.46	468.79	2320.90	11,610.50	570.23	10.03	0.38	0.38	1.38	1.31	0.84	0.86	0.34	5.51
7	8.33	51.42	41.22	136	23,792.62	12.78	386.94	0.27	14.83	423.32	2246.20	9541.50	581.90	9.19	0.26	0.34	1.28	1.71	1.11	0.73	0.23	6.41
8	8.15	52.74	42.34	139	24,792.31	13.88	379.78	0.25	14.06	470.41	2330.20	11,309.00	594.68	10.02	0.50	0.37	1.35	1.11	0.97	0.87	0.27	6.44
9	8.22	51.90	41.41	142	22,253.10	12.35	380.60	0.25	13.67	432.60	2258.90	10,065.40	558.09	9.40	0.46	0.29	1.29	1.30	1.28	0.87	0.21	5.14
10	8.23	52.42	42.34	147	24,792.31	13.80	392.86	0.24	14.55	480.36	2350.30	10,520.90	711.29	10.04	0.44	0.34	1.38	1.52	1.41	0.88	0.24	4.87
11	8.49	52.68	42.34	150	25,392.13	14.48	381.29	0.22	15.02	473.00	2349.10	11,358.90	580.34	10.03	0.48	0.39	1.25	0.91	1.16	0.84	0.28	5.40
12	8.02	52.13	41.78	153	23,992.56	13.75	359.81	0.17	14.98	444.55	2284.60	10,394.50	619.86	9.97	0.48	0.29	1.30	0.81	1.00	0.77	0.27	4.02
13	8.33	51.48	41.54	143	22,172.88	14.82	395.20	0.20	15.93	438.58	2261.00	9784.60	604.60	9.55	0.50	0.38	1.19	0.79	0.78	0.76	0.30	3.63
14	8.26	52.29	42.10	148	22,593.46	13.41	400.98	0.19	14.72	459.73	2315.50	10,456.20	567.31	10.37	0.44	0.31	1.39	1.50	0.94	0.83	0.26	6.90
15	8.31	52.69	42.42	146	25,591.75	15.33	391.07	0.20	16.44	429.90	2251.70	12,646.90	555.55	9.62	0.52	0.36	1.30	1.30	1.26	0.79	0.26	7.05
16	8.24	52.40	42.10	146	25,792.00	15.29	396.85	0.22	16.91	458.40	2313.50	12,437.90	712.91	10.08	0.40	0.34	1.40	0.99	1.12	0.84	0.28	5.52
17	8.18	51.89	41.34	152	24,292.47	12.73	387.35	0.18	13.65	429.54	2251.60	9724.70	598.42	9.64	0.43	0.34	1.40	1.30	1.05	0.68	0.28	3.78
18	8.18	52.87	42.70	142	25,991.94	13.58	390.52	0.19	14.24	483.88	2354.80	12,420.90	579.69	10.15	0.44	0.28	1.22	1.31	0.80	0.79	0.28	4.54
19	8.10	51.91	41.62	140	24,592.37	12.22	402.91	0.16	14.44	439.57	2270.70	9745.90	572.09	9.62	0.44	0.36	1.41	1.30	1.37	0.76	0.26	5.40
20	8.21	52.45	42.34	147	25,991.32	12.99	405.53	0.16	14.44	484.87	2350.90	10,592.90	512.84	10.42	0.42	0.38	1.30	1.12	1.22	0.81	0.32	5.37
21	8.34	52.15	41.92	150	24,592.37	13.20	351.27	0.16	15.00	441.12	2280.10	10,608.90	559.62	9.17	0.44	0.39	1.80	0.99	0.82	1.08	0.40	6.11
22	8.15	52.51	42.24	152	24,792.31	15.24	316.99	0.22	13.33	468.72	2327.10	10,740.70	600.64	10.27	0.62	0.41	1.75	1.01	1.36	1.03	0.49	7.26
23	8.26	51.83	41.48	150	22,253.41	12.78	398.64	0.16	13.92	438.41	2259.00	9889.90	612.91	9.56	0.67	0.46	1.70	1.06	1.62	0.96	0.47	7.63

Table 1. Cont.

S.N.	pH	EC	TDS	TA	Cl <sup>-</sup>	NO <sub>3</sub> <sup>-</sup>	SO <sub>4</sub> <sup>2-</sup>	PO <sub>4</sub> <sup>3-</sup>	NH <sub>4</sub> <sup>+</sup>	Ca <sup>2+</sup>	Mg <sup>2+</sup>	Na <sup>+</sup>	K <sup>+</sup>	Sr	Cd	Co	Cr	Cu	Fe	Mn	Pb	Zn
		mS/cm	g/L	mg/L										µg/L								
24	8.38	51.69	41.62	149	22,293.09	14.26	391.34	0.22	16.48	438.21	2269.30	10,147.90	695.98	9.58	0.75	0.50	1.90	1.38	1.79	0.99	0.66	7.93
25	8.21	52.14	41.99	146	24,592.37	13.63	393.82	0.16	15.29	444.22	2295.40	11,180.10	606.13	9.99	0.68	0.51	1.73	1.50	1.66	1.02	0.59	8.22
26	8.37	52.44	42.22	148	25,191.57	13.46	364.49	0.16	16.08	463.92	2325.70	11,720.90	597.54	10.26	0.75	0.46	1.74	1.69	1.87	1.05	0.73	8.23
27	8.36	52.40	42.12	156	24,991.01	15.5	386.11	0.22	15.38	460.75	2317.40	11,536.90	577.79	10.18	0.76	0.48	1.66	1.70	1.87	1.04	0.79	7.58
28	8.00	52.57	42.46	154	23,792.62	13.16	381.84	0.17	14.01	483.45	2352.20	10,865.90	596.38	10.40	0.72	0.49	1.75	1.51	1.82	1.14	0.70	7.86
29	8.31	51.55	41.34	162	22,533.01	15.12	395.61	0.21	16.40	430.51	2252.60	9934.20	564.46	9.37	0.71	0.46	1.72	1.89	1.88	1.12	0.76	8.17
30	8.32	51.43	41.28	160	23,152.82	14.09	390.24	0.19	15.86	425.14	2249.70	9681.40	581.45	9.21	0.72	0.50	1.91	1.91	1.94	1.15	0.76	8.28
Mean	8.26	52.17	41.95	146	24,326.40	13.85	385.77	0.21	14.98	452.13	2298.01	10,801.44	594.08	9.86	0.51	0.38	1.44	1.24	1.29	0.88	0.38	6.05
Min	8.00	51.42	41.22	128	22,172.88	12.22	316.99	0.16	13.08	423.32	2246.20	9541.50	512.84	9.17	0.20	0.28	0.96	0.69	0.75	0.68	0.17	3.63
Max	8.49	52.87	42.74	162	25,991.94	15.50	407.45	0.29	16.91	486.99	2355.90	12,646.90	712.91	10.42	0.76	0.51	1.91	1.91	1.94	1.15	0.79	8.28
STD	0.11	0.42	0.44	8.06	1211.19	0.96	18.22	0.04	1.034	20.03	38.05	891.41	44.03	0.38	0.15	0.07	0.27	0.34	0.37	0.14	0.20	1.49
Median	8.26	52.28	42.05	146.50	24,592.53	13.78	390.38	0.21	14.99	444.39	2295.55	10,674.80	584.83	9.98	0.47	0.38	1.38	1.30	1.25	0.84	0.28	5.61
Co. Var	0.01	0.01	0.01	0.05	0.05	0.068	0.05	0.27	0.07	0.04	0.02	0.08	0.07	9.86	0.51	0.38	1.44	1.24	1.29	0.88	0.38	6.05



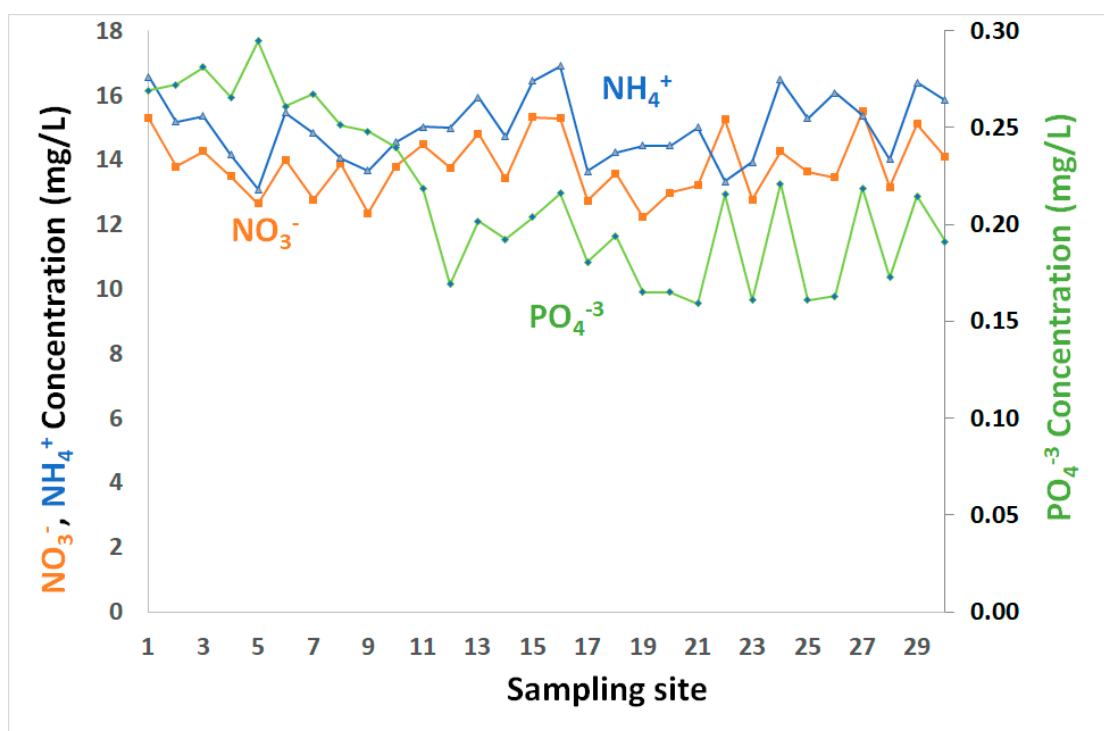


Figure 3.  $\text{NO}_3^-$ ,  $\text{NH}_4^+$  and  $\text{PO}_4^{3-}$  in coastal water of the GoA, Jordan.

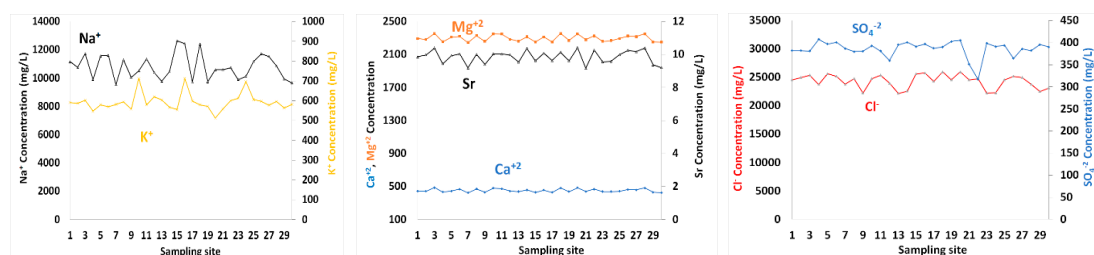


Figure 4. The spatial distributions of  $\text{Na}^+$ ,  $\text{K}^+$ ,  $\text{Mg}^{2+}$ ,  $\text{Sr}$ ,  $\text{Ca}^{2+}$ ,  $\text{SO}_4^{2-}$  and  $\text{Cl}^-$  concentrations in surface seawater of GoA, Jordan.

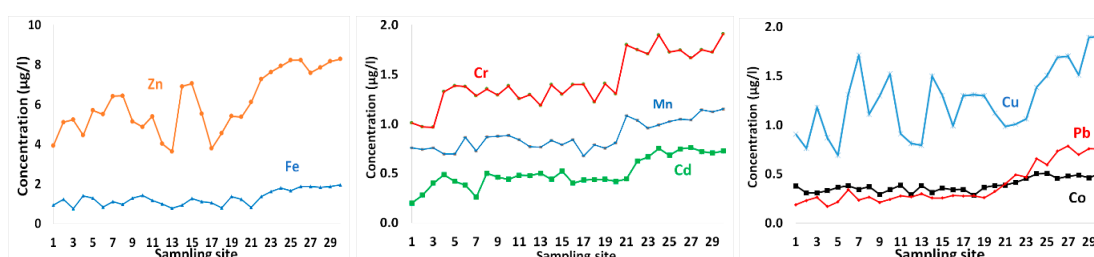


Figure 5. Metal contents in surface seawater in GoA, Jordan.

In addition, these pH values are probably attributable to low growth levels, and production of algal biomass that would contribute organic acids to coastal water when decomposed. While the weather conditions of high temperatures and abundant sunlight allow phytoplankton to grow in abundance, our sampling campaign coincided with a period of nutrient-depleted and stratified water, where photosynthetic activity was at its lowest levels. Manasreh et al. [23] recorded the lowest chlorophyll-a levels in summertime and highest during winter. Microbial decomposition of dead phytoplankton, algae and other flora, produces humic substances, organic acids and amino acids that

raise the seawater acidity. Additionally, higher temperature during summertime will reduce dissolved CO<sub>2</sub> levels in coastal water and increases the pH value.

TA (total alkalinity) concentrations ranged between 128 and 162 mg/L, with an average value of 146 mg/L. These high values are consistent with the coastal water's buffering capacity (due to high contents of calcium carbonates) of the GoA's water. The TDS varied from about 41.22–42.74 g/L with an average of 41.95 g/L. EC varied from 51.42 to 52.87 mS/cm with average of 52.17 mS/cm. Similar values of TDS were reported along the Saudi GoA with an average of 41.4 g/L [9].

The spatial patterns of TDS in water showed insignificant variations (Figure 2). The elevated levels of TDS indicate highly saline conditions in the GoA and are primarily attributable to its geographic location in a subtropical desert region, with very high evaporation rates, very low precipitation, and negligible freshwater input. Manasreh et al. [23] reported an evaporation of 2 m/year in the GoA with an increasing salinity toward the north. In addition to these salinity raising factors, the high TDS values are linked to water stratification and poor water circulation during the sampling period. These factors created unique environmental conditions of higher temperature, evaporation, and salinity than average, compared to the average range for oceans. Lack of input of freshwater into the coastal water contributes to high salinity water. A negligible supply of terrigenous sediments into the water results in clear water conditions with high transparency.

Higher water density is often observed in summertime in response to high salinity (due to poor water mixing). From July–August, a stratified water column dominates with thermocline and pycnocline, occurring at about 250 m in 2013 and near 350 m in 2014 and 2015 [23]. The northwards currents drag warm and saline waters to GoA from the Red Sea [25,29]. The flow of surface water from the Red Sea to the GoA is triggered by the high evaporation rate, where the flowing water offsets the evaporation loss [30]. The GoA surface seawater temperature is 2 °C lower than that of the Red Sea, where the flowing water brings heat that increases temperature, evaporation rates and salinity of surface seawater layer.

In addition, the TDS values become higher in August, corresponding to the summer season, a period of high dust storm events. The atmospheric dust input is an significant source of salts (and metals) to GoA water [8,9]. The GoA is located in a desert-belt region with frequent dust storms, where Negev and Sinai Deserts are in the west and the Arabian Desert is in east. It is believed that large quantities of dust aerosols delivered to GoA is originated from adjacent deserts [31]. Dust deposition will significantly influence the composition of GoA seawater, where the deposition rate of dust in GoA is one of the highest on Earth [13] ranging between 28 g/m<sup>2</sup>/year in the northwestern part [13] to about 34.68 g/m<sup>2</sup>/year in Aqaba city at the northeastern corner of GoA [8]. TDS was significantly correlated with Ca<sup>2+</sup>, Mg<sup>2+</sup>, Na<sup>+</sup>, Sr and Cl<sup>−</sup> with  $r = 0.86, 0.88, 0.81$  and  $0.70$ , respectively (Table 2). These ions are major contributors to seawater salinity.

Inorganic nutrients (nitrate, ammonium, phosphate) are minor constituents of seawater, but are essential for marine ecosystem productivity and growth. Relatively low levels of inorganic nutrients (NH<sub>4</sub><sup>+</sup>, PO<sub>4</sub><sup>3−</sup>, NO<sub>3</sub><sup>−</sup>) (Table 1 and Figure 3) were observed in surface water layer, consistent with the findings of others [23,32–35]. The coastal water in GoA is in extremely oligotrophic conditions, with very limited nutrients supplied to Gulf's water through terrestrial runoff.

Nitrates are present in all water samples, where the concentrations increased slightly in some locations, although not all. NO<sub>3</sub><sup>−</sup> concentrations ranged between 12.22 mg/L up to 15.50 mg/L, with overall mean and median levels of about 13.85 and 13.78 mg/L, respectively (Table 1 and Figure 3). Ammonium levels fluctuated from 13.08 and 16.91 mg/L, with mean and median values of about 15 mg/L (Table 1 and Figure 3). While the nutrient levels generally varied, their variations showed no spatial trends. Nitrate and ammonium showed relatively similar ups and downs and were significantly correlated with  $r = 0.65$  (Table 2).



**Table 2.** Correlation matrices for surface water quality parameters.

[illegible]

Nitrate is the major nitrogen species in the oxic zone, while ammonium dominates in the anoxic zone. The nitrification is a kinetic reaction and is dependent on several water conditions such as salinity, pH, and Eh [36,37]. However, the coexistence of nitrate and ammonium can trigger or slow down nitrogen conversion like nitrification or denitrification [38]. The coexistence of  $\text{NH}_4^+$  and  $\text{NO}_3^-$  may result from poor mixing in groundwater, especially in locations where both nitrogen species are released from active pollution sources [39,40]. The GoA waters are well oxygenated with redox indicators of oxidizing conditions [9]. Manasreh et al. [23] observed that the GoA water is well-ventilated due to the annual water mixing with complete saturation (100%).

This suggests that the presence of high  $\text{NH}_4^+$  levels could be associated with leaks from sewer system and/or because of water discharged from fish farm or fertilizer plume. Phosphate varied from 0.16–0.29 mg/L, with an average of 0.21 mg/L (Table 1 and Figure 3). Relatively higher concentrations were observed in the northern GoA, in close proximity to the phosphate terminal, where deposition of dust containing phosphate during loading/unloading activities may contribute phosphates to seawater. While phosphate showed low levels, it is the limiting nutrient for phytoplankton growth. Weak correlations between phosphate and both ammonium and nitrate were observed with  $r = -0.02$  and 0.18, respectively (Table 2).

Aeolian deposits can provide important nutrients which stimulate the primary productivity in marine ecosystems [41–44], especially in oligotrophic water [45], like the GoA [10,20]. However, they can also deliver various contaminants that negatively impact the aquatic biodiversity.

Similar observations were reported by Badran [46], where phosphate and nitrate levels in surface water varied seasonally, with the lowest in summer and the highest in winter. The sampling period was concurrent with water stratification, and the concentrations of nutrients in surface water layer were low. In the winter season, winds drive convective mixing of deep (nutrient-rich) and surface waters, where nitrate and other nutrients are injected into the euphotic zone, resulting in seasonal plankton blooms [47]. The highest productivity (chlorophyll-a) is expected during the winter season, which declines to minimum levels in summertime [23].

Water stratification and high sunlight irradiation during summertime further draw down the inorganic nutrients in the surface water by enhancing primary productivity at the subsurface water layer (50–75 m) [34]. During photosynthesis, phytoplankton assimilate nutrients, and it is the availability of inorganic nitrogen that often limits the rate of primary production in the sea [32]. Nutrients uptake within the euphotic zone in oligotrophic water body results in a considerable depletion of their levels. Phytoplankton communities in oligotrophic waters are likely to survive by utilizing recycled nutrients [48,49].

Nutrient levels in the southern Red Sea are greater than its northern and central regions. The inflow of surface water from the Red Sea to GoA (to compensate for the high evaporation loss) is a contributing factor to lower levels of nutrients observed in GoA seawater (oligotrophic water). In late summer, an increase of 25% in nutrient levels is observed in the southern Red Sea compared to the central region, due to the inflow of nutrient-rich waters from the Gulf of Aden to the southern area of the Red Sea [50]. The highest levels of phosphate in the southern Red Sea are usually observed in October, following upwellings in the Arabian Sea.

In addition to water mixing, it is likely that nitrate and other nutrients are associated with atmospheric deposition, as this area experiences frequent dust storms. Rare flash floods carrying terrigenous sediments can be a minor contributor to nitrate and phosphate in coastal water. Dust from the phosphate terminal in the GoA provides further evidence of contribution of aeolian dust to coastal water. Fish farming and wastewater discharges may be important sources of nitrate and phosphate, as water samples were collected adjacent to the coastlines closer to touristic, industrial and other human facilities.

Sr content in seawater ranged between 9.17 and 10.42 mg/L (Table 1). The concentrations of  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Na}^+$  and  $\text{K}^+$  in surface seawater layer varied from 423.32–486.99, 2246.2–2355.9, 9541.5–12,646.9 and 512.84–712.91 mg/L, respectively (Table 1). High temperature and evaporation rates are main

contributors to high levels of ions, among others. The spatial distributions of  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$  and Sr in surface water exhibited relatively similar patterns with no trends.  $\text{Ca}^{2+}$  was positively correlated with  $\text{Mg}^{2+}$ , Sr, and  $\text{Na}^+$  with correlation coefficients of 0.97, 0.85, and 0.52, respectively. The correlation coefficients between  $\text{Mg}^{2+}$  and Sr was 0.87, and between  $\text{Mg}^{2+}$  and  $\text{Na}^+$  is 0.59. The  $\text{K}^+$  levels in seawater was not significantly correlated with any cation tested.

$\text{Cl}^-$  exhibited spatial changes in surface water, with concentrations ranging from 22,172.88–25,991.94 mg/L. The  $\text{SO}_4^{2-}$  values in seawater samples varied between 316.99 and 407.45 mg/L. Cl was correlated with  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$  and  $\text{Na}^+$  with correlation coefficients of 0.54, 0.60 and 0.74, respectively. Similar to  $\text{K}^+$ , the  $\text{SO}_4^{2-}$  content showed no significant correlations with any other ions of seawater. TDS values were well correlated with  $\text{Cl}^-$ ,  $\text{Ca}^{2+}$ ,  $\text{Mg}^{2+}$ ,  $\text{Na}^+$  and Sr, with  $r = 0.70, 0.86, 0.88, 0.81$  and  $0.81$ , respectively. Whereas salinity was neither correlated with  $\text{SO}_4^{2-}$  nor with  $\text{K}^+$ . Dust deposition to GoA is also an important contributor to TDS and other ions.

Table 3 compares seawater chemistry of the GoA (northernmost Red Sea) analyzed in this study relative to other regions of the Red Sea. Relatively elevated levels of pH, TA,  $\text{Cl}^-$ ,  $\text{NO}_3^-$ ,  $\text{PO}_4^{3-}$ ,  $\text{NH}_4^+$ ,  $\text{Mg}^{2+}$  and K were observed for GoA, compared to the central and northern Red Sea. Whereas  $\text{SO}_4^{2-}$ , TDS, Ca, and Na showed values that are comparable to or lower than those for other parts of the Red Sea water.

**Table 3.** Comparison of water quality parameters in different regions of the Red Sea and GoA.

	pH	EC mS/cm	TDS g/L	TA mg/L	$\text{Cl}^-$	$\text{NO}_3^-$	$\text{SO}_4^{2-}$	$\text{PO}_4^{3-}$	$\text{NH}_4^+$	Ca mg/L	Mg	Na	K
Present study (northernmost Red Sea)	8.26	52.17	41.95	146	24,326.4	13.85	385.77	0.21	14.98	452.13	2298.01	10,801.44	594.08
Northern Red Sea (Egypt) [51]	7.7	60.3	42	98	23,607	12.4	1260		0.39	738	1570	12,339	287
Central Red Sea (Jeddah, Saudi [52])	8.1	72.55	43.55	128	22,336	1	2440	<0.13	-	496	1512	11,920	-

The concentrations of metals in surface seawater layer are shown in Table 1 and presented in Figure 5. Zn showed the highest concentration, with an average concentration of 6.05  $\mu\text{g/L}$ . Other metals that followed were in the order  $\text{Cr} > \text{Fe} > \text{Cu} > \text{Mn} > \text{Cd} > \text{Pb} = \text{Co}$  (Table 1, Figure 5).

Spatial variability of metals contents in seawater samples exhibited increasing trends toward the south (Figure 5), where the industrial complex is located. Potential impacts from heavy metals are commonly confined to areas in the vicinity of urban or industrialized regions on the coastal edge. However, these levels of metals also suggest that they have probably been derived from multiple sources, including a geogenic origin.

Al-Taani et al. [9] reported high levels of dissolved oxygen, with redox values indicating oxidizing conditions in the coastal water GoA, which may favor immobilization of some metals with relatively low levels in seawater samples. Zn in seawater varied from 3.63–8.28  $\mu\text{g/L}$  with an average of 6.05  $\mu\text{g/L}$  (Table 1 and Figure 5). These values are higher than those reported for the Saudi GoA [9], the offshore surface seawater of Red Sea [53], the average oceanic concentration [54], and the Mediterranean surface seawater [55] (Table 4).

In addition, atmospheric dry deposition is the primary external source of trace metals to GoA [14]. Aeolian dry fluxes of certain trace elements (e.g., Cd, Pb, Cu and Zn) to the ocean water may surpass those of riverine sources [56,57]. Aeolian dust of Zn to GoA ranges between 1.68  $\text{mg/m}^2/\text{year}$  (in Eilat city at the northwestern corner; [31]) and 4.02  $\text{mg/m}^2/\text{year}$  (in Aqaba city in northeastern region; [8]). High concentrations of Zn were observed GoA seawater ranging from 5.71–11.55  $\mu\text{g/L}$  [58] in the vicinity of Industrial Complex.

**Table 4.** Comparison of selected metals ( $\mu\text{g/L}$ ) in surface seawater of GoA relative to other regions.

	Cd	Co	Cr	Cu	Fe	Mn	Pb	Zn
Present study	0.51	0.38	1.44	1.24	1.29	0.88	0.38	6.05
Saudi GoA [9]	0.03	0.236	0.957	6.183	15.255	0.259	0.202	3.323
Offshore surface seawater, northern Red Sea [53]	0.53	0.15	-	0.11	1.58	0.11	0.36	0.22
Oceanic concentration [54]	0.07	0.005 <sup>a</sup>	0.33	0.12	0.04	0.01	0.001	0.4
Mediterranean surface seawater [55]	0.42 <sup>b</sup>	1 <sup>c</sup>	3.4 <sup>d</sup>	0.2 <sup>e</sup>	0.056–0.336 <sup>f</sup>	0.11–0.19 <sup>f</sup>	0.05 <sup>g</sup>	0.17

<sup>a</sup>: [59,60]. <sup>b</sup>: [61]. <sup>c</sup>: [62]. <sup>d</sup>: [63,64]. <sup>e</sup>: [65,66]. <sup>f</sup>: [67]. <sup>g</sup>: [62,65].

Fe contents of seawater varied from 0.75–1.94  $\mu\text{g/L}$  with a mean value of 1.29  $\mu\text{g/L}$ . These concentrations of Fe showed spatial variability in surface water layer with generally greater values in the southern GoA (Figure 5). During the stratified summer, surface water becomes enriched in Fe [13], but the winter mixing of surface and deep water layers, decreases these Fe levels. The average concentration of Fe measured in the present study is relatively comparable to that for the Red sea offshore seawater [53], but higher than the averages for oceanic concentration [54] and the Mediterranean seawater [68] (Table 4). Higher concentrations of Fe were observed in the Saudi GoA with about 15.25  $\mu\text{g/L}$ , suggesting that atmospheric dry deposition in this area is more intense. It is believed that Fe is probably derived from crustal sources. The average dry flux of Fe to the GoA waters varied from about 216  $\text{mg/m}^2/\text{year}$  [31] to 440  $\text{mg/m}^2/\text{year}$  [8].

Cr levels varied from 0.96–1.91  $\mu\text{g/L}$  with average value about 1.44  $\mu\text{g/L}$  (Table 1). These values are higher than those observed in the Saudi GoA seawater [9], the mean oceanic level [54], but less than those for the Mediterranean Sea [63,64] (Table 3). The spatial pattern of Cr suggests that these high levels of Cr at the southern end of GoA are most likely related to discharge of brine water from desalination plant [68]. In addition to the industrial wastewater, mineral dust from fertilizer and cement factories remain potential sources of Cr to GoA seawater. atmospheric aerosol deposition to GoA fluctuated between 0.96  $\text{mg/m}^2/\text{year}$  in Eilat city [31], and about 1.42  $\text{mg/m}^2/\text{year}$  in Aqaba city [8].

Mn content in southern water samples exhibited higher values relative to the northern part of GoA (Table 1 and Figure 5), and it is likely to originate from anthropogenic emissions [31]. The Mn contents in surface seawater ranged between 0.68 and 1.15  $\mu\text{g/L}$ , with an average of 0.88  $\mu\text{g/L}$ , which is three times as much as that measured for the Saudi GoA [9]. These values are also greater than those reported for the offshore water of Red Sea [53], the average oceanic concentration, and those for the Mediterranean Sea [67] (Table 4). Windblown dust of Mn to the uppermost eastern GoA (Eilat city) averaged 5.28  $\text{mg/m}^2/\text{year}$  [31], whereas, in the northernmost extension, a mean value of 10.29  $\text{mg/m}^2/\text{year}$  was reported [8]. In addition, Mn is probably related to desalination plants in the neighboring cities (Eilat, Taba and Haql) [3,9], where various heavy metals, including Mn, may be released with the discharged water of thermal desalination plants, depending on the metal alloys used [68,69].

Cd concentrations in seawater ranged between 0.2 and 0.76  $\mu\text{g/L}$ , with an average of 0.51  $\mu\text{g/L}$  (Table 1). Similarly, elevated levels of Cd were found in the southern GoA (Figure 5), where industrial activities are concentrated. Comparable levels of Cd have been reported by Shriadah et al. [53] for the northern Red Sea (offshore seawater), but higher than those for the average oceanic levels of about 0.07  $\mu\text{g/L}$  [54]. Additionally, the average Cd content in surface seawater measured in the present study is greater than those for Saudi GoA and the Mediterranean Sea (Table 4).

In addition to desalination plants, the anthropogenically derived Cd (and other metals such as Pb and Co) is likely related to the discharge of cooling water and sewage in the southern GoA [26,70,71]. The average concentration of Cd from atmospheric dust varied from 0.012  $\text{mg/m}^2/\text{year}$  in Eilat city [31] to 0.04  $\text{mg/m}^2/\text{year}$  in Aqaba city [8].

Co content in surface seawater exhibited little spatial variations, with relatively higher levels were observed in the southern portion of GoA (Figure 5). Co ranged from 0.28–0.51  $\mu\text{g/L}$ , with average Co values of about 0.38  $\mu\text{g/L}$  (Table 1). Lower average Co contents were reported for the northern Red Sea offshore water [53] and for the Saudi GoA [9]. Co occurs in seawater at concentrations below 0.005  $\mu\text{g/L}$  [59,60]. However, higher average value has been detected in the surface water of Mediterranean Sea [62]. Potential sources of Co in GoA seawater are likely similar to those of Co [9]. Dust particles collected from GoA region showed an average of 0.1  $\text{mg/m}^2/\text{year}$  [31].

Relatively elevated levels of Cu were detected in seawater samples, varying from 0.69–1.91  $\mu\text{g/L}$ , with greater values observed for the southern sampling sites (Figure 5). These values are higher than the average ocean level [54] and the offshore water of Red Sea [53] (Table 3). An average Cu value of 0.2  $\mu\text{g/L}$  was reported for the Mediterranean Sea [67,68]. However, elevated levels of Cu were detected in the Saudi GoA (at the Jordan–Saudi border). The atmospheric dry deposition flux of Cu to GoA region ranged from 0.38  $\text{mg/m}^2/\text{year}$  in Eilat [31] to 0.68  $\text{mg/m}^2/\text{year}$  in Aqaba city [8]. Elevated concentrations of Cu have been reported in Jordanian GoA water, in the range of 0.74–2.28  $\mu\text{g/L}$  [58], and are higher than those measured in this study.

Pb levels in seawater samples varied from 0.17–0.79  $\mu\text{g/L}$ , with a spatial pattern of increasing levels in the southern part of GoA. The average Pb value of 0.38  $\mu\text{g/L}$  is comparable to those measured in offshore surface water sites of Red Sea [53], but higher than those reported for the Saudi GoA [9], the average ocean, and the Mediterranean surface seawater (Table 3). Pb is of anthropogenic origin, mainly from fossil fuel burning [72]. Elevated concentrations of Pb were found in the Jordanian GoA water ranging between 0.73 and 1.43  $\mu\text{g/L}$  [56]. GoA receives high dry flux of Pb varying from 0.8  $\text{mg/m}^2/\text{year}$  in Eilat [31] to about 1.42  $\text{mg/m}^2/\text{year}$  in Aqaba city [8]. All metals tested were significantly correlated (Table 2) indicating that they may have been derived from similar sources.

#### 4. Conclusions

GoA is a place for rich and diverse marine ecosystems. It is highly vulnerable to pollution, where human activities in the bordering countries are intense, with high potential for water contamination. This requires cross-border collaboration to protect these naturally diverse but stressed ecosystems. This study intended to ascertain the seawater quality conditions along the eastern coast of GoA, Jordan. The sampling campaign coincided with a period of low levels of inorganic nutrients, low rates of algal growth with reduced microbial decomposition of dead algal cells. In addition to prevailing saline conditions of high temperatures and high evaporation rates, the water stratification and intense dust storms are the major contributing factors to the observed seawater chemistry. The surface distribution of water quality variables showed spatial variations with no specific patterns, except for metal contents, which exhibited southward increasing trends, closed to the industrial complex. The vast majority of these quality parameters showed relatively higher values compared to those of other regions.

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