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Possible Pre-Seismic Indications Prior to Strong Earthquakes That Occurred in Southeastern Mediterranean as Observed Simultaneously by Three VLF/LF Stations Installed in Athens (Greece)

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Abstract: In this work, we present the analysis of VLF/LF sub-ionospheric propagation data to study anomalies possibly related to very recent strong (M > 5.5) earthquakes (EQs) that occurred in the southeastern Mediterranean in September-October 2021 and January 2022. We used the signal of one transmitter located at Negev in Israel (29.7 kHz) as received by three VLF/LF receivers (two of them using identical SW and HW) installed, at a close distance to each other, in Athens (Greece). This study employed multiple methods and techniques to analyze the reception amplitude data to identify any possible EQ-related anomalies. More specifically, first, we used both statistical and criticality analysis methods such as the "nighttime fluctuation method" (NFM), the "terminator time method" (TTM), and the "natural time" (NT) analysis method. These methods have satisfactorily been applied in the past in a series of other studies leading to interesting results. Moreover, we additionally used two more analysis techniques focusing on the signal's amplitude characteristics. The first is the wavelet analysis of the nighttime part of the signal's amplitude. It is based on the Morlet wavelet function, aiming to unveil the possible existence of atmospheric gravity waves (AGWs) before EQ. The second is named "long wavelength propagation capability" (LWPC), which simulates the amplitude of the signal and is based on the reflection parameters of ionosphere and by searching for increases or decreases of the electron density profile of the ionospheric D layer concerning the shifts of the minima of terminator times (TTs) in the diurnal variation of the signal. Finally, in this work, we summarize our findings and discuss possible "pre-", "co-", and "post-" seismic effects as observed from all the work.

Keywords: earthquakes (EQs); 2021 Crete EQs; 2022 Cyprus EQ; VLF/LF sub-ionospheric propagation data; nighttime fluctuation method (NFM); natural time analysis (NTA); terminator time method (TTM); long wavelength propagation capability (LWPC); wavelet analysis

1. Introduction

The very-low- to low-frequency (VLF/LF) monitoring of the lower ionosphere is an effective and simple technique for probing this part of the atmosphere, especially for detecting ionospheric disturbances caused by various terrestrial and extra-terrestrial



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Copyright: © 2023 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). extreme events, e.g., [1–4]. More and more researchers and amateur radio engineers have shown great interest in installing VLF/LF recording systems (receivers) at different places worldwide to observe and understand these extreme events, e.g., [1–8]. More specifically, seismo-ionospheric studies are the most controversial among many scientists [1]. However, a unified and well-established theoretical pre-seismic model that shows the linkage of three specific planetary layers is called "lithosphere–atmosphere–ionosphere coupling" (LAIC) [1]. This model clearly describes the physical processes behind the preparation of an impending earthquake (EQ), but it is still poorly understood at the moment [1]. Considering those physical mechanisms, this model separates the different observable quantities by grouping them into five suggested channels. These are electromagnetic (EM), chemical, thermal, electrostatic, and acoustic [1,9].

On the other hand, other observed EM quantities have also been used to unveil EQrelated signatures prior to EQs. For example, these are kHz–MHz fracto-electromagnetic (fracto-EM) emissions, ultra-low-frequency (ULF) magnetic field data, and seismic electric signals (SES) [10–13]. Specifically, to these observed EM quantities, different methods and techniques from the theory of complex systems have been applied in past studies by examining in this manner other characteristics of the signal (and the system of which these are observables), such as the information content, fractality, and criticality, e.g., [10–13].

For VLF/LF sub-ionospheric propagation data, two well-established criticality analysis methods have been applied. Specifically, these are "the method of critical fluctuations" (MCF) and the "natural time" (NT) analysis, which can unveil the approach of criticality; they have been applied to the case of the 2016 Kumamoto EQs and 2020 Samos EQ [8,14,15]. However, there are differences in the application of these two methods. The first one was applied to the raw amplitude recordings. In contrast, the second one was used on the three statistical quantities of the "nighttime fluctuation method," which is based on the raw amplitude recordings of the VLF/LF signal [8,14,15].

A second statistical method called the "terminator time method" (TTM) is very important for the detection of shifts of the terminator times (TTs), which appear as minima around the planetary sunrise and sunset time, before the occurrence of an EQ [16]. This behavior of shifts of TTs was first reported in Japan in 1995 before the famous Kobe EQ (M7.1) [16]. In two recent studies of the 2020 Samos EQ, significant shifts of TTs have been observed towards night for two studied sub-ionospheric propagation paths. They have been characterized as possible anomalous shifts prior to the EQ [8,17].

A conventional technique for the analysis of VLF/LF nighttime amplitude recordings is the "wavelet analysis" technique, which is based on the Morlet wavelet function, aiming to unveil the possible existence of atmospheric gravity waves (AGW) prior to an EQ [18–21]. These anomalies of AGW appear as periodic structures in the VLF amplitude recordings [19–21]. More specifically, anomalies associated with AGW have been observed, for example, before the 2016 Imphal EQ in India, where signatures of periodic structures were revealed with a periodicity of ~60 min [19].

Another simple and well-established technique for searching ionospheric anomalies before EQs by simulating the VLF amplitude recordings is called the "long wavelength propagation capability" (LWPC) [22]. Specifically, this technique simulates the amplitude and phase of the signal by knowing the location of the transmitter and the receiver at any place on earth. This technique treats the earth and the lower ionosphere as a waveguide by using single or multiple reflections to guide an emitted VLF/LF signal based on the sub-ionospheric propagation characteristics of the propagation path. Besides the signal's amplitude and phase simulation, another essential factor is the calculation of the lower ionosphere's electron density (ED) profile for a given location and time. By knowing that the radio signals use the D and E layers to be transmitted, they are affected by them.

Many extreme events can disturb the lower ionosphere, such as EQs, solar flares, and geomagnetic storms, by changing locally or globally the ED profile [17,23–25]. However, the calculation of the ED profile is achieved by two significant parameters in LWPC, which are called the "steepness parameter" (β) and "ionospheric reflection height" (h'), taken

from the theory of Wait's exponential ED profile [26,27]. For the 2020 Samos EQ, pre-EQ shifts of TTs have been observed as pre-seismic indications [17]. By numerically simulating around those TTs, by taking into account a small excerpt of the signal's amplitude, the simulated profile was reproduced at the receiver's location with proper corroboration with its observed profile (amplitude recordings) [17]. Using Wait's exponential formula for a particular set, the ED profile was calculated precisely at the time of occurrence of the minima of the shifted TTs. In that work, increases or decreases in the ED profile were generally observed as possible pre-seismic indications before the 2020 Samos EQ [17].

By taking into account all those mentioned above, in this work, we used all these described methods and techniques to search for pre-EQ ionospheric anomalies in the VLF amplitude recordings during strong ($M_L \ge 5.5$) EQs that occurred in the southeastern Mediterranean in September and October of 2021, as well as in January of 2022. The first case concerns three EQs that occurred in September and October of 2021 and hit Crete Island (Greece), causing several damages, or occurred close to Crete, M_L 5.8 on 27 September 2021, $M_L6.3$ on 12 October 2021, and $M_L6.1$ on 19 October 2021, hereafter referred to as "the 2021 Crete EQs" (please see more information in Section 2). The second case concerns an EQ that hit a sea area northwest of Cyprus Island on January of 2022, M_L6.4 on11 January 2022, due to which no damage was reported, hereafter referred to as "the 2022 Cyprus EQ'' (please see more information in Section 2). In our study, we utilized three VLF/LF receivers located in Athens (Greece) to study these EQs [2,8]. These VLF/LF receivers have recently been established searching for ionospheric perturbations caused by various extraterrestrial or terrestrial phenomena, but especially for studying EQs [2,8]. The first one has been operating since March 2020, while the two other identical VLF/LF stations, having different hardware and software from the initial, were installed in June and September of 2021, respectively [2,8].

In the first part of this study, we analyze the raw amplitude recordings of the signal from the three stations using statistical methods such as NFM and TTM. At the same time, we use one criticality analysis method called the "NT analysis method" to analyze the three statistical quantities of NFM. Then, any statistical anomalies and criticality signatures possibly correlated with the EQs under study are searched for a period of 15 days before each of them. In addition, we analyze the raw amplitude data using the "Morlet wavelet analysis searching for AGWs". The results of identified perturbations in the VLF/LF amplitude recordings caused by AGWs are unveiled in 32 to 128 min and in 15 days before each one of the studied EQs. Moreover, in this work, we include the LWPC technique by numerically simulating the amplitude recordings and selecting a small excerpt of the signal where TTs exist. Anomalies associated with increases or decreases in ED according to the shifted TTs are revealed 15 days before each studied EQ (2021 Crete EQs, 2022 Cyprus EQ).

By understanding all those mentioned above, the structure of this manuscript is the following. In the Section 2 of this manuscript, we give details about the sub-ionospheric propagation data, the EQs under study, and other possibly ionosphere-influencing extreme events. In the Section 3, we briefly describe the methods and techniques used for this study. In the Section 4, we present the results; finally, in the Section 5, we give conclusions about the whole study.

2. Sub-Ionospheric Propagation Data, Studied EQs, and Other Non-Examined Extreme Events

This section gives detailed information about the VLF/LF receivers used in our study and the characteristics of VLF/LF sub-ionospheric propagation data. For this study, we choose one transmitter (call sign: ISR), which is located at Negev (37.4094° N, 27.3252° E) in Israel, and its transmitted frequency is 29.7 kHz, while the reception of the signal is conducted by three VLF/LF stations, all located in the prefecture of Attica in Greece. More specifically, two of the three VLF/LF receivers have the same features. The first one is located in a forest area (38.0317° N, 23.6637° E) to the west of Athens, called Aspra Chomata (call sign: ACH), close to the urban complex of Athens, while the second one is located in a suburban area (38.0317° N, 23.8572° E) to the east of Athens, called Gerakas (call sign: GER). At this point, it is noted that these two receivers were recently installed in June and September of 2021, respectively, and they are operating with 30 s sampling. These two receivers can record signals up to 100 kHz, and the detailed description of the setup of their hardware and software and the recorded transmitters are presented in detail in [2].

On the other hand, the third VLF/LF receiver is located at the University of West Attica (call sign: UWA) in the prefecture of Attica in Athens (Greece) and has different features from the other two abovementioned VLF/LF receivers; this receiver has a sampling frequency of 1 s and records frequencies up to 47.5 kHz. The information about the monitored transmitters and key information about this VLF/LF receiver are given in [8]. In addition, a detailed description of the hardware and software is provided on the UltraMSK website (https://www.ultramsk.com (accessed on 29 November 2022)). In Figure 1, we present a satellite image of the prefecture of Attica in Greece. The three VLF/LF stations are shown, together with the distances between each pair of receivers.



Figure 1. Satellite image of the prefecture of Attica in Greece. The three VLF/LF stations (ACH, GER, and UWA) are depicted with light green squares, while the distance between each pair of them is also shown.

In Table 1, we give information about the EQs under study, taken from the seismic catalog of the National Observatory of Athens (NOA) (https://bbnet.gein.noa.gr/HL/databases/database) (accessed on 29 November 2022). All four of them took place in the southeastern Mediterranean from September to October 2021 and January 2022. Specifically, the EQs mentioned above are the 2021 Crete EQs and 2022 Cyprus EQ, which are indicated in this table in the column "EQ group". Table 1 includes the EQ's occurrence time, magnitude, depth, and coordinates of each epicenter. We should also mention that these EQs have a magnitude of 5.5 above, so they are expected to significantly perturb the lower ionosphere [1]. Moreover, we note that for this study, we search for EQ-related signatures within 15 days before each target EQ. In Figure 2, we show the map of the eastern Mediterranean where a representative fifth Fresnel zone of the sub-ionospheric propagation path is shown between each receiver and the transmitter ISR. The EQs chosen for this study are relatively close or within the fifth Fresnel zone of each monitored propagation path between the transmitter ISR and each receiver (UWA, ACH, and GER).

As already mentioned in Section 1, the ionosphere is sensitive to various extreme events occurring in the terrestrial, atmospheric, or space environment (e.g., earthquakes, volcanoes, typhoons, geomagnetic storms, solar flares, lighting activity, and weather storms), while it is not easy to indisputably attribute an ionospheric anomaly to a specific event if more than one of them are happening in close time distance. For this reason, we check

"EQ Group" Based on the Time Occurrence and the Location of the Epicenter of the Studied EQ	Date and Time of Occurrence (UT)	Magnitude (M _L)	Depth (km)	Latitude	Longitude
2021 Crete EQs	27 September 2021 06:17:21	5.8	9.6	35.1521° N	25.2736° E
2021 Crete EQs	12 October 2021 09:24:02	6.3	10.4	34.8944° N	26.4716° E
2021 Crete EQs	19 October 2021 05:32:35	6.1	58.5	34.7131° N	28.2532° E
2022 Cyprus EQ	11 January 2022 01:07:49	6.4	34.7	35.1398° N	31.9537° E

our results against all recorded possibly ionosphere-influencing phenomena that occurred during the time periods of interest.

Table 1. List of examined EQs in the southeastern Mediterranean.

In this direction, we initially checked the geomagnetic conditions for each studied case (2021 Crete EQs and 2022 Cyprus EQ). For the 2021 Crete EQs, the corresponding geomagnetic indices Dst, Kp, ap, and Ap have been reported in a previous work [2]. Specifically, in Fig. 16 of ref. [2], two minor geomagnetic storms, of minimum Dst value ~ -50 nT, happened on 17 September 2021 and 12 October 2021. At the same time, a simultaneous increase by exceeding an appropriate threshold (different for each index) is observed for the rest of the indices (Kp, ap, Ap).

Additionally, for the 2022 Cyprus EQ, the variation of geomagnetic indices (Dst, Kp, ap, Ap), as well as the peak flux (W/m^2) of solar flares, are all shown in five panels for the period from 1 December 2021 until 15 February 2022. We have retrieved the data of Dst, Kp, ap, and Ap from the World Data Center for Geomagnetism of Kyoto (https://wdc.kugi.kyoto-u.ac.jp/wdc/Sec3.html (accessed on 29 November 2022)), while for solar flares we checked the data on solar X-ray flux from the National Oceanic and Atmospheric Administration (https://www.swpc.noaa.gov/products/goes-x-ray-flux (accessed on 29 November 2022)). As is evident from Figure 3, one moderate geomagnetic storm of -91 nT (min Dst) occurred on 14 January 2022, while a simultaneous enhancement (shown with red color) in the two geomagnetic indices (Kp, ap) appears for a very short time (one bin) period. Finally, in the bottom panel, we present all solar flares of the M and X classes that occurred in the abovementioned period. It can be observed that there were several solar flares of the M class in the entire period, but no X class solar flare.

However, other possibly ionosphere-influencing extreme events during the 2021 Crete EQs and 2022 Cyprus EQ have also been checked. We first checked the parameter of convective available potential energy (commonly abbreviated as CAPE) as an indication of atmospheric instability by observing its distribution on the map from the Ventusky search engine (https://www.ventusky.com (accessed on 29 November 2022)), the data of which are provided by the National Oceanic and Atmospheric Administration (NOAA) and the Deutscher Wetterdienst (DWD), while lightning activity was checked using the "Blitzortung.org" lightning-detection network (https://www.blitzortung.org/en/historical_maps.php (accessed on 29 November 2022)) and "lightningmaps.org" (https://www.lightningmaps.org/ (accessed on 29 November 2022)). However, no thunderstorms or typhoons were observed. Finally, the volcanic eruption database of the Global Volcanism Program of the Smithsonian Institution (https://volcano.si.edu/search_eruption.cfm (accessed on 29 November 2022)) was checked for possibly erupting volcanoes during the period of the studied EQs (2021 Crete EQs and 2022 Cyprus EQ), but no volcanic eruptions were recorded.

At this point, we should clarify that there were two strong events in the area of Crete Island that happened during the time period examined for the case of the 2022 Cyprus EQ (2 December 2021 29 January 2022). Both were of $M_L \ge 5.5$ according to the seismic catalog of NOA. Therefore, they may have affected the lower ionosphere. Specifically, these

EQs occurred on 26 December 2021 18:59:02 UT, at $(35.2029^{\circ} \text{ N}, 26.8497^{\circ} \text{ E})$, within the fifth Fresnel zone of the examined paths, with a magnitude of 5.5 M_L and a focal depth of 6.1 km, and on 29 December 2021 05:08:09 UT, at $(34.791^{\circ} \text{ N}, 25.1303^{\circ} \text{ E})$, out of the fifth Fresnel zone of the examined paths, with a magnitude of 5.7 M_L and a focal depth of 67.1 km. Although for the specific studied period (2 December 2021 29 January 2022) we focus on the 2022 Cyprus EQ, we also check for perturbations possibly related to them.



Figure 2. Map of the eastern Mediterranean. The 5th Fresnel zone of the propagation path between the three receivers and the transmitter ISR is indicated, whereas the epicenters of the studied EQs are shown as pentagrams and each occurrence date is shown alongside the corresponding epicenter.



Figure 3. Geomagnetic indices Dst, Kp, ap, and Ap, as well as solar flares during the time period 1 December 2021–15 February 2022, which are indicated as pentagrams. Red-colored parts indicate values exceeding the thresholds marked by the horizontal black dashed lines. In the case of the bottom panel, the threshold separates M class from X class solar flares.

3. Methods

In this section, we describe all the methods and techniques used in this study. Specifically, we present each method in a separate subsection as follows. In Section 3.1, we present the nighttime fluctuation method (NFM), while in Section 3.2, the terminator time method (TTM) is described. In Section 3.3, we present the natural time (NT) method, which can identify criticality before the occurrence of EQs. In Section 3.4, we show the wavelet analysis method for the VLF nighttime amplitude recordings. Finally, in Section 3.5, we discuss the "long wavelength propagation capability" (LWPC), which is a well-defined technique to simulate the electron density (ED) profile versus the ionospheric height at any time and anywhere on Earth based on the location of the transmitter and the receiver.

3.1. Nighttime Fluctuation Method (NFM)

The nighttime fluctuation method (NFM) was proposed for the detection of statistical anomalies caused by ionospheric disturbances concerning a variety of extreme events such as EQs [1,2]. First of all, the raw nighttime amplitude data (in dB) are extracted from the diurnal variation of the amplitude recordings by taking into account a selected nighttime interval. However, the terminator times, represented as minima in the signal's amplitude, are avoided from being included in the "nighttime window". Thus, after determining the nighttime interval, the next step is to initially calculate the mean value (denoted as $\langle A(t) \rangle$) of ±15 days around the day of interest plus the day of interest. Subsequently, the residual variation of the amplitude of the signal (dA(t)), defined as $dA(t) = A(t) - \langle A(t) \rangle$, is calculated, where A(t) is the signal's amplitude at time t. The usage of a ±15-day sliding window reduces the "long-term" variations, unveiling the "short-term" variations. Finally, the daily values for the three statistical parameters "TR" (trend), "DP" (dispersion), and "NF" (nighttime fluctuation) are calculated as:

$$TR = \frac{\sum_{N_s}^{N_e} dA(t)}{N_e - N_s} \tag{1}$$

where *TR* represents the mean value of dA(t), and N_e and N_s are the start and endpoint of the chosen nighttime interval (starting and ending time points);

$$DP = \sqrt{\frac{1}{N_e - N_s} \sum_{N_s}^{N_e} (dA(t) - TR)^2}$$
(2)

where the *DP* is actually the standard deviation of dA(t), and:

$$NF = \sum_{N_s}^{N_e} (dA(t))^2 \tag{3}$$

After calculating of the daily valued time series of these three statistical quantities, the normalized values TR^* , DP^* , and NF^* are computed as $X^* = (X - X_{\pm 15})/\sigma_{\pm 15}$, where $X_{\pm 15}$ and $\sigma_{\pm 15}$ are the mean value and the standard deviation of ± 15 days around the day of interest, respectively. Any statistical anomaly in these daily valued time series that exceeds $\pm 2\sigma$ could possibly be related to an EQ preparation process, a geomagnetic storm [1,2,8,28], or any other phenomenon that can influence the lower ionosphere. The adopted criterion of $\pm 2\sigma$ has been determined from statistical analysis using VLF data of a long time period during which many EQs (with magnitude > 5.5) occurred [29–31]. In fact, this method has recently been applied extensively to identify ionospheric anomalies prior to EQs as an increase in *TR* and decrease in *DP* and *NF*, e.g., [1,8]. It should be mentioned that generally, the usage of a ± 15 -day window around the day of interest includes information from the "future", so this is appropriate only for a posteriori analysis [8]. To analyze in "real time" one should use a single side window.

3.2. Terminator Time Method (TTM)

The terminator time method (TTM) was proposed for the statistical analysis of the occurrence time of the minima in the amplitude and the phase of the VLF signal [1,16], which are close in time to the local (planetary) sunrise time and sunset time. These minima are referred to as sunrise terminators (SRTs) and sunset terminators (SSTs), respectively, or generally as terminator times (TTs), and are created by the interference of different propagation waves (modes of propagation) of the VLF signal—that is, the ground wave and the sky wave [16,32].

Compared with the neighboring days, a significant shift in the SRTs or SSTs is considered an anomaly before an EQ when the lower ionospheric height usually is decreased [32]. In other words, an early appearance of an SRT or a late appearance of an SST, which means an anomalous increase in the duration of the "VLF day" ("VLF daylength", D_{VLF}), as compared with the previous days, is considered to be an EQ precursor [16].

The TTM was initially applied to the strong Kobe EQ (M7.1) that occurred in Japan on 17 January 1995, for which significant shifts in the TTs appeared before the EQ's occurrence [16,33]. By this concept, several other studies have also reported shifts in TTs and consequent increases in D_{VLF} before an impending EQ [34–37]. Furthermore, many other statistical studies have also reported correlations between EQs and TT anomalies, with maximal shifts occurring 0–4 days prior to the main EQ event [34,38–45]. Furthermore, some studies based on numerical simulation of the diurnal variation of the amplitude of the VLF signal, considering the characteristics of the VLF propagation path, the transmitter, and the receiver, are applied to determine the TTs [17,25,37].

In applying the TTM, we initially find the time of appearance of two minima in the diurnal variation of the signal (amplitude or phase), which are close in time with the planetary sunrise and sunset time of each day, respectively. Using these time locations, which are the morning and evening TTs, we form two TT time series—one for the morning minima, denoted t_m , and one for the evening minima, denoted t_e . Subsequently, we use 2 days around the day of interest window (5 days total width) to calculate the running mean for each of the aforementioned time series, forming 2 new time series designated as $\langle t_m \rangle$ and $\langle t_e \rangle$ for the morning and evening TTs, respectively. Finally, the running mean time series are subtracted from the respective TT time series to form the residual TT time series $dt_m = t_m - \langle t_m \rangle$ and $dt_e = t_e - \langle t_e \rangle$, respectively [16,33]. Moreover, we calculate the "VLF daylength" as $D_{VLF} = t_e - t_m$, and similarly to the TT time series, we consecutively calculate the running mean VLF daylength time series, $\langle D_{VLF} \rangle$, and the residual "VLF daylength" time series $dDVLF = D_{VLF} - \langle D_{VLF} \rangle$. Any statistical anomaly in the residual TTs or the residual "VLF daylength" exceeding $\pm 2\sigma$ of the whole considered period is investigated as possibly being EQ-related. We must note that the specific procedure uses the running mean values to reveal the shift of the TTs or prior to an impending EQ by removing their seasonal variability.

3.3. Natural Time (NT) Method

The NT time series analysis method has initially been applied to the ultra-lowfrequency (≤ 1 Hz) seismic electric signals (SES) [46–48] and has been shown to be optimal for enhancing the signals in the time-frequency space [49]. Furthermore, the application of NT analysis to various seismo-EM signals, including VLF sub-ionospheric propagation data, has been presented in detail in [50]. However, in very recent studies, the NT analysis method has been applied to the other channels of LAIC by taking data from satellites and ground-based stations, showing clear findings of the existence of critical dynamics before EQs [51–53]. In the following, we will briefly present the key notions of this method.

Initially, for a number of N events, we determine the NT of the occurrence of the k-th event as $\chi_k = k/N$. Next, we determine the "energy" of each event in NT, which is denoted as Q_k for the k-th event. At this point we have to mention that Q_k corresponds to different kinds of quantities, depending on the time series under analysis. For example, in the case of seismic events Q_k is the seismic energy released (seismic moment), while for

the dichotomous SES signals Q_k corresponds to the SES pulse duration [47]. However, in the case of the fracto-EM emission signals in the MHz band, which are non-dichotomous signals, Q_k is the energy of each event using consecutive amplitude values above a noise threshold as described in [54].

Then, we study the evolution of the pair of (χ_k, Q_k) , where $p_k = Q_k / \sum_{n=1}^{N} Q_n$ is the normalized energy released during the k-th event. The approach of a dynamical system to criticality is identified by means of the variance $\kappa_1 = \langle \chi^2 \rangle - \langle \chi \rangle^2$ of NT weighted with p_k , where $\langle f(\chi) \rangle = \sum_{n=1}^{N} p_k f(\chi_k)$. Hence, the quantity κ_1 can be written as $\kappa_1 = \sum_{k=1}^{N+1} p_k \chi_k^2 - \left(\sum_{k=1}^{N+1} p_k \chi_k\right)^2$. Moreover, the entropy (S_{nt}) in NT is defined as $S_{nt} = \sum_{k=1}^{N} p_k \chi_k \ln \chi_k - \left(\sum_{k=1}^{N} p_k \chi_k\right) \ln \left(\sum_{k=1}^{N} p_k \chi_k\right)$ [11,55]. The entropy in NT is a dynamic entropy, depending on the order of the events [55]. In addition, S_{nt-} , the entropy under time reversal $(Tp_m = p_{N-m+1})$, is also studied [55].

In many studied dynamical systems, it has been found that the value of κ_1 is a measure to quantify the extent of the organization of the system at the onset of the critical stage [11]. The criticality is reached when (a) κ_1 takes the value $\kappa_1 = 0.07$ and (b) at the same time both the entropy in NT and the entropy under time reversal satisfy the condition S_{nt} , $S_{nt-} < S_u = (\ln^2/2) - 1/4$ [11,56], where S_u is the entropy of the uniform distribution in NT [11,55].

In the special case of NT analysis of foreshock seismicity [47–49,55,57], we study the evolution of the quantities $\kappa_{1,}S_{nt,}S_{nt-}$, and $\langle D \rangle$ with time, where $\langle D \rangle$ is the "average" distance between the normalized power spectra $\Pi(\tilde{\omega}) = \left|\sum_{k=1}^{N} p_k exp(j\tilde{\omega}\chi_k)\right|^2$ ($\tilde{\omega}$ stands for the angular frequency in NT) of the evolving seismicity and the theoretical estimation of $\Pi(\tilde{\omega})$ for $\kappa_1 = 0.07$, $\Pi_{\text{critical}}(\tilde{\omega}) \approx 1 - \kappa_1 \omega^2$. Moreover, an "event" for the NT analysis of seismicity is considered to be any data point (EQ) of the original seismicity time series that surpasses a magnitude threshold, M_{Thres} .

The analysis starts with an appropriate low threshold and taking into account only an adequate number of first in the order of occurrence events. Next, the subsequent events, in their original order, are taken into account one-by-one. For each additional event that is taken into account, the quantity χ_k is rescaled within the interval (0,1] and all $\kappa_{1,}S_{nt,}S_{nt-}$, and $\langle D \rangle$ are re-calculated. This way, a temporal evolution of these quantities is attained. The described procedure is repeated for several increasing, values of M_{Thres} for each studied geographic area, and everything is repeated for different overlapping areas.

The seismicity is considered to be in a true critical state, a "true coincidence" is achieved, as soon as (a) κ_1 takes the value $\kappa_1 = 0.07$, (b) at the same time both the entropy in NT and the entropy under time reversal satisfy the condition S_{nt} , $S_{nt-} < S_u$, and three additional conditions are satisfied: (c) the "average" distance $\langle D \rangle$ should be smaller than 10^{-2} , i.e., $\langle D \rangle = \langle \left| \Pi \left(\widetilde{\omega} \right) - \Pi_{\text{critical}} \left(\widetilde{\omega} \right) \right| \rangle < 10^{-2}$ (this is a practical criterion for signaling the achievement of spectral coincidence) [11]; (d) the parameter κ_1 should approach the value $\kappa_1 = 0.070$ "by descending from above", i.e., before the main event the parameter κ_1 should gradually decrease until it reaches the critical value 0.070 (this rule was found empirically) [11,48]; (e) the above-mentioned conditions (a)–(d) should continue to be satisfied even if the considered M_{Thres} or the area within which the seismicity is studied are changed (within reasonable limits).

The use of the magnitude threshold excepts some of the weaker EQ events (those events whose magnitude is $\langle M_{Thres} \rangle$) from the NT analysis. However, the usage of the magnitude threshold is valid because some recorded magnitudes are not considered reliable due to the seismographic network. On the other hand, the application of various M_{Thres} values is useful in determining the time range within which criticality is reached. This is because, in some cases, it is found that more than one time point may satisfy the rest of the NT critical state conditions (a)–(d), and criterion (e) is the one that finally reveals the true time of criticality.

For the application of NT analysis to VLF data, we follow the paradigm of the NT analysis of seismicity by using the non-normalized VLF propagation quantities (defined in Section 3.1) to define the "energy" Q_k and the necessary threshold values as in [8,15].

3.4. Wavelet Analysis

Wavelet analysis is a well-recognized technique that can represent the time-dependent signal into a diffuse two-dimensional time-frequency image, referred to as a scalogram, by applying the continuous wavelet transform, calculating in this manner the wavelet power spectrum (WPS) [18]. Generally, this technique divides the signal into different scale components by unveiling the periodicities of "wave-like" structures of the signal [18]. For example, in VLF amplitude recordings of nighttime amplitude data, periodicities have been revealed in a period from 1 to 128 min before the Imphal EQ, which occurred in India on 4 January 2016 (M = 6.7) [19]. These "wave-like structures" are associated with the existence of the pre-seismic atmospheric gravity waves (AGWs) emitted around the epicenter of the EQ, which directly affect the lower ionosphere and indirectly the nighttime amplitude recordings of the signal [19–21].

In our case, we compute the WPS based on the Morlet wavelet. First, we rearrange the data in a one-minute time sample and then subtract each amplitude value from its tenminute running mean by taking the residual amplitude of the signal [19,20]. Subsequently, we draw the cone of influence (CoI) in the WPS, which represents a boundary, showing that beyond it, the WPS values are not nominal due to the addition of zeros (zero padding), which are required to convert the total number of data points to a power of two to compute the WPS [19,20].

3.5. Long Wavelength Propagation Capability

The long wavelength propagation capability (LWPC) code is a well-known technique for simulating VLF signals and has been extensively used in past studies [6,17,25,37,58–61]. It was developed by the US Navy to numerically simulate VLF amplitude and phase in various directions and ionospheric conditions [22]. For the simulation of different ionospheric conditions, the LWPC uses a set of other models, such as HOMOGENEOUS, CHI, RANGE, and GRID, along with the default model LWPM. In addition, precise inputs must be supplied to the models using some inbuilt substrings of LWPC, namely the TABLE and the EXPONENTIAL. Specifically, the code treats the earth's surface as an ideal conductor, and the ionosphere follows the well-known Wait's two-component model of exponential electron density (ED) profile [26,27]. This code follows the waveguide mode theory of electromagnetic signal propagation and is very efficient for probing the lower ionosphere. This work uses the RANGE model and the EXPONENTIAL substring, where the parameters effective signal reflection height (h' in km) and (β in km⁻¹) are taken as inputs. The altitude profile of lower-ionospheric electron density $(N_e e)$ and the electronneutral collision frequency are taken from [27]. First, we normalize the arbitrary unit of the recorded VLF signal to the logarithmic decibel scale in a non-seismic condition using LWPC. This information is essential to calculate the ideal or normalized signal amplitude at the reception (at the receiver) of the propagation path at the local mid-noon time. The signal amplitude profile, h', and β are treated as the background values for different propagation paths for non-seismic conditions.

For computing the seismogenic perturbed signal profile and the ED variation, we use different sets of h' and β to obtain the observed signal amplitude value at the SRT and SST times for all the propagation paths for the entire period of observations. For computing the ED, we feed those values at a particular ionospheric height (h in km) to the well-known Wait's two-component formula as:

$$N_e(h) = 1.43 \times 10^{13} \exp(-0.15h') \exp[(\beta - 0.15)(h - h')]$$
(4)

To find the anomalies in ED associated with EQs, we use a different set of h' and β . In this study, we focus on the importance of the SRT and SST timings, which are the

prime source of observed perturbation. We follow the trend of these SRTs and SSTs for the entire study period and try to simulate their normalized values using the REXP subprogram in the LWPC. The combinations of h' and β are used to obtain the simulated VLF signal during both nighttime and daytime. For the numerical simulation, we focus on the period of the time of appearance of the terminator time (TT) minima, which existed in the amplitude of the signal around the planetary sunrise and sunset time. We compute the simulated VLF signal exactly and around the morning and evening TTs by giving different combinations of h' and β until we find the specific combinations that lead to a match between the simulated and observed amplitude. For the Crete EQs, for the UWA receiving station, we observed one SRT (SRT1) and two SSTs (SST1 and SST2). For SRT1, we found the range for β to be 0.26–0.45 km⁻¹ and for h' to be 68.1–79.3 km. This means that for the Crete EQs and for SRT1, as observed at the UWA receiving station, all the specific combinations of β and h' values that lead to matching are such that β lies within the range 0.26–0.45 km⁻¹ and h' within the range 68.1–79.3 km. Of course, for each date only a specific combination of h' and β values lead to matching between the observed and the simulated amplitude at SRT1. However, mentioning here all these specific combinations is omitted. For the SST1 and SST2, we determined the β ranges as 0.25–0.48 km⁻¹ and 0.26-0.47 km⁻¹, respectively, and h' ranges as 66.2-79.7 km and 65.7-79.8 km, respectively. For the GER receiving station, we found for SRT1, SST1, and SST2 the ranges for β to be $0.27-0.59 \text{ km}^{-1}$, $0.26-0.55 \text{ km}^{-1}$, and $0.25-0.57 \text{ km}^{-1}$, respectively, while for h' the ranges were found to be 67.8–79.9 km, 65.3–79.5 km, and 48.7–79.5 km, respectively. For the Cyprus EQ, from the UWA station, we choose only SRT1 and SST and the β ranges were found to be 0.23–0.47 km⁻¹ and 0.23–0.42 km⁻¹, respectively, whereas the h' ranges were found to be 71.9–79.3 km and 67.0–78.7 km, respectively. The simulated VLF signal is reproduced with proper corroboration with its observed profile. This simulation of the observed profile for a few points is essential in order to prove that the combinations of h' and β are right. Subsequently, we compute the ED profile from Equation (4), taking only the combinations of h' and β precisely at the time of appearance of morning and evening TTs. Then, we compute an average profile of ED using only the non-shifted (non-perturbated) TT minima, namely, the "quiet days" (QD), which appeared around the studied period of the EQ. Finally, we plot in diagrams of "ED vs. height" the QD (quiet profile) alongside the computed ED profile of each shifted (perturbated) TT. Increases or decreases in the computed ED profiles concerning the height from QD (quiet profile) are considered pre-seismic indications before an EQ.

4. Analysis of the Lower Ionosphere Prior to "2021 Crete EQs" and "2022 Cyprus EQ"

In this section we present the analysis results for the three VLF/LF receivers (UWA, GER, and ACH) using the different methods and techniques presented in Section 3. Specifically, in Section 4.1 we present the results obtained with NFM, while in Section 4.2 we give the results of sequential plots of the variation of filtered signals of amplitude data and we also present the TTM results. In Section 4.3 we show the results of the NT analysis method. Finally, the results of the two techniques of wavelet analysis and LWPC are shown in Sections 4.4 and 4.5, respectively.

4.1. NFM Analysis Results

This section analytically presents the results of the nighttime fluctuation method (NFM) by analyzing the data from the three VLF/LF stations installed in the prefecture of Attica in Athens (Greece). In the two cases of EQs under study (see Section 2), we choose the nighttime interval 21:00–02:00 UT for each baseline (between the transmitter ISR and each receiver). It is important to note that the specific nighttime interval comprises two parts, namely, one part belonging to the previous day (date) and one to the next day. In this particular application of NFM, we attribute each daily value of each statistical parameter to the next day because its (local) nighttime part is significantly larger than that of the previous day.

For the 2021 Crete EQs, we choose the period from 2 September 2021 to 29 October 2021, including at least a range of 15 days before each studied EQ's occurrence date (see Table 1), which is appropriate for searching for EQ-related anomalies. For the 2022 Cyprus EQ, we selected the period from 2 December 2021 to 29 January 2022 for UWA and GER, while for the ACH station, due to a lack of data in November (because of some experimental setup optimizations and later on due to damage to part of the equipment by lightning), the analyzed period is from 18 December 2021 to 29 January 2022. It should be mentioned at this point that in our analysis, we have excluded any kind of artificial excerpts from the daily amplitude recordings by keeping only the natural fluctuations.

In Figure 4, we present an example of the NFM analysis results, using the GER station data, for the study of the 2022 Cyprus EQ. As we can see from Figure 4, a possible EQ-related anomaly appeared in TR^* on 5 January 2022, while a simultaneous increase in NF^* is also presented on the same date. In DP^* , one anomaly appeared on 4 January 2022, while the other (on 11 December 2021) appeared outside of 15 days and seems unrelated to the EQ. It may be associated with two other EQ events ($M_L \ge 5.5$) on 26 December 2021 and 29 December 2021 close to Crete Island (see Section 2). However, as we can see from Figure 4, an anomaly on 2 December 2021 not related to the EQ is simultaneously observed in the three statistical parameters. This anomaly is not likely to be related to the EQ under study or the EQs mentioned above due to the considerable time distance, which exceeds the range of 15 days. Therefore, it cannot be attributed to any possibly ionosphere-influencing terrestrial or extraterrestrial extreme event.



Figure 4. Time evolution of three statistical quantities, TR^* , DP^* , and NF^* , of the NFM analysis method for the sub–ionospheric path GER–ISR is shown in the top, middle, and bottom panels, respectively. The time period represented in the figure is from 2 December 2021 to 29 January 2022. Red solid horizontal lines indicate the corresponding $+2\sigma / -2\sigma$ limits; σ is calculated for the whole studied period of each panel. The date of EQ is shown on top of each panel by the purple vertical line segment, marked on the top of each panel.

In Table 2, we present the results of the NFM analysis for the 2022 Cyprus EQ. Specifically, we present the anomalous data alongside any existing anomaly in the three VLF propagation quantities (TR^* , DP^* , and NF^*) by indicating at which of the three stations (UWA, GER, and ACH) this was observed. The last column shows the attribution of each anomaly to the examined EQ, as well as to any other possibly ionosphere-influencing extreme event. The anomalies that appeared on 2 December 2021 and 27 January 2022 (see Table 2) are not attributed to any extreme event. As already mentioned, the first one appears more than 15 days before the 2022 Cyprus EQ, while the second was not followed by any extreme event.

Date of Appearance of the Anomaly (UT)	TR*	DP*	NF*	Possibly Associated Extreme Event(s)
2 December 2021	GER	GER	UWA, GER	-
11 December 2021		GER		b
21 December 2021		UWA		b, c
4 January 2022	ACH	ACH, GER	ACH	a
5 January 2022	UWA, ACH, GER	ACH	UWA, ACH, GER	a
27 January 2022		ACH		-

Table 2. Results of NFM for the 2022 Cyprus EQ. UWA, GER, and ACH denote the station where the corresponding anomaly in TR^* , DP^* , or NF^* has been identified.

The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow 2022$ Cyprus EQ (6.4 M_L , 11 January 2022); $b \rightarrow a 5.6 M_L$ EQ that occurred on 26 December 2021 18:59:02 (UT) at (35.2029° N, 26.8497° E); $c \rightarrow a 5.7 M_L$ EQ that occurred on 29 December 2021 05:08:09 (UT) at (34.791° N, 25.139° E).

Similar to Table 2, in Table 3 we present the UWA and GER stations' results for the 2021 Crete EQs. All information on anomalies observed in the ISR–GER sub-ionospheric propagation path were taken from [2]. The general picture drawn by Table 3 is that anomalies associated with EQs are observed at different dates within 15 days before each EQ under study.

Table 3. Results of NFM for the 2021 Crete EQs. UWA and GER denote the station where the corresponding anomaly in TR^* , DP^* , or NF^* has been identified.

Date of Appearance of the Anomaly (UT)	TR*	DP*	NF*	Possibly Associated Extreme Event(s)
5 September 2021	UWA	UWA, GER	UWA, GER	-
9 September 2021	GER	GER	GER	a
19 September 2021	UWA, GER			a
28 September 2021		GER		b
4 October 2021		UWA		b, c
7 October 2021	UWA, GER		UWA, GER	b, c
9 October 2021		GER		b, c
11 October 2021	UWA		UWA, GER	с
12 October 2021		UWA		С
27 October 2021	UWA		GER	-

The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow 1st 2021$ Crete EQ (5.8 M_L , 27 September 2021); $b \rightarrow 2nd 2021$ Crete EQ (6.3 M_L , 12 October 2021); $c \rightarrow 3rd 2021$ Crete EQ (6.1 M_L , 19 October 2021).

4.2. Diurnal Variation and TTM Analysis Results

This section presents the sequential plots of the diurnal variation of the amplitude of the filtered (by a Gaussian filter) signal, with TTs noted, as well as the results of the TTM analysis. In this study, we used the three stations to find evidence of TT shifts before each EQ under study (see Section 2). Before TT detection, we resampled the data of the UWA station to 30 s (before the filtration), so as to obtain the same sampling rate as the two other stations (see Section 2). This down-sampling of the UWA data helped us to detect the position of TT minima in parallel for all three stations. After detecting TT shifts from the sequential plots, TTM was applied, as described in Section 3.2.

In Figure 5, we give an example of the diurnal variation in sequential plots for the 2022 Cyprus EQ. In this figure, the diurnal variation of the amplitude is presented in the form of stacked 24 h signals, shifted by 10 dB. The date corresponding to each stacked signal appears in the middle (above the signal), while the selected period of this figure is from 27 December 2021 to 14 January 2022, including 15 days before the 2022 Cyprus EQ. Moreover, in Figure 5, it can be observed that two sets of TTs appear on the morning side (purple and brown circles), while on the evening side (green circles), there is only one set of

TTs. In this study, we show TTs shifted towards night (see Section 2). Specifically, in the morning TTs, the brown minima appeared largely shifted (see black ellipses) towards the night on 1, 2, 8 and 9 January 2022, preceding the date of the EQ (red color). On the other hand, pre-seismic shifts are also indicated in the evening TT, which are marked as black ellipses. More specifically, significant shifts in the evening TT minima have been observed on 28 December 2021 and 9 January 2022, while a gradual shift of TTs appeared to have existed from 2 January 2022 to 6 January 2022.



Figure 5. Diurnal variation of the amplitude of the VLF signal of the propagation path GER–ISR for the time period 27 December 2021–14 January 2022. Each signal is vertically shifted by +10 dB in regards to the signal of the previous day. Brown and purple circles indicate the minima identified as morning TTs before and after the sunrise, respectively. Green circles are the evening TTs, which appear before the sunset. Finally, the date of the EQ is marked with red color. The black circles indicate the existence of shifts of TTs.

The overall results of the 2022 Cyprus EQ by observing shifts of TTs for each station (UWA, ACH, and GER) are presented in Table 4. Specifically, this table gives the date of the appearance of the anomaly in UT, as well as the extreme events possibly associated with the observed anomalies. We mentioned that the chosen time for all the stations is the same as for the GER–ISR sub-ionospheric propagation path. In Table 4, we indicate each shift of TTs with the call name(s) of the stations where this is observed. TTs are referred to on the morning side as t_{m1} and t_{m2} , and for the evening side, as t_e . The t_{m1} and t_e are the first selected sets of TTs towards night, while t_{m2} is the second set chosen. It can be observed that most of the shifts of TTs appeared simultaneously for all stations on various dates before the EQ by considering a searching range of 15 days to find a related anomaly. Furthermore, we mention that the gradual shift of TTs, shown in Figure 5, has also very clearly been observed in the ACH station from 2 January 2022 to 6 January 2022, while in the UWA station some indications of shifts have been found on the specific dates. Yet, from Table 4, it is also noted that only one anomaly on 28 December 2021, found in t_e , is possibly attributed to an EQ that occurred close to Crete Island on 29 December 2021 (see Section 2).

Similarly, by studying the 2021 Crete EQs, in Table 5, we present the summary results of the shifts of TTs detected from diurnal variation sequential plots, taking data from the UWA and GER stations. In this table there are four TT sets, two on the morning side and two on the evening side, denoted as t_{m1} , t_{m2} , t_{e1} , and t_{e2} , respectively. Again, we searched for shifts of TTs 15 days before each EQ under study. As we can see, anomalies recorded by two stations (UWA and GER) have appeared on various dates within a time range of 15 days before each EQ under study. However, we do not have data from the ACH station

for September and October of 2021, as has been mentioned in Section 2, so there are no results for this station included in Table 5.

Table 4. Results from the diurnal sequential plots for the 2022 Cyprus EQ. UWA, GER, and ACH denote the station where the corresponding anomaly in t_{m1} and/or t_e has been identified.

Date of Appearance of the Anomaly (UT)	Shift in t_{m1}	Shift in t_e	Possibly Associated Extreme Event(s)
28 December 2021		UWA, GER, ACH	a, b
1 January 2022	UWA, GER, ACH		b
2 January 2022	GER, ACH	GER, ACH	b
3 January 2022		GER, ACH	b
4 January 2022		GER, ACH	b
5 January 2022		UWA, GER, ACH	b
6 January 2022		UWA, GER, ACH	b
7 January 2022		UWA	b
8 January 2022	UWA, GER, ACH		b
9 January 2022	UWA, GER, ACH	UWA, GER, ACH	b

The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow a 5.7 M_L$ EQ that occurred on 29 December 2021 05:08:09 (UT) at (34.791° N, 25.139° E); $b \rightarrow 2022$ Cyprus EQ (6.4 M_L , 11 January 2022).

Table 5. Results from diurnal sequential plots for 2021 Crete EQs. UWA and GER denote the station where the corresponding anomaly in TTs has been identified.

Date of Appearance of the Anomaly (UT)	Shift in t_{m1}	Shift in t_{m2}	Shift in t_{e1}	Shift in t_{e2}	Possibly Associated Extreme Event(s)
13 September 2021				UWA, GER	a
15 September 2021			UWA	GER	a
17 September 2021		GER			a
18 September 2021		UWA, GER			а
19 September 2021		UWA, GER	UWA		а
20 Septemebr 2021		UWA, GER			а
21 September 2021	UWA				a
22 September 2021	UWA				a
24 September 2021	UWA, GER		UWA		a
25 September 2021	UWA		UWA		a
26 September 2021			UWA		a
27 September 2021			UWA		a, b
28 September 2021	GER		UWA	UWA	b
2 October 2021		GER	GER		b
3 October 2021			GER		b
4 October 2021			GER		b, c
6 October 2021					b, c
7 October 2021	UWA, GER		UWA		b, c
8 October 2021			UWA		b, c
9 October 2021			UWA		b, c
10 October 2021	UWA				b, c
11 October 2021	GER		UWA, GER		b, c
14 October 2021			GER		С
15 October 2021			GER		С
16 October 2021	UWA				С
17 October 2021	UWA		UWA		С

The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow 1st 2021$ Crete EQ (5.8 M_L , 27 September 2021); $b \rightarrow 2nd 2021$ Crete EQ (6.3 M_L , 12 October 2021); $c \rightarrow 3rd 2021$ Crete EQ (6.1 M_L , 19 October 2021).

In this work, we have also applied the TTM analysis to show which pre-seismic anomalies have adequately been shifted (morning or evening TT) for each day by exceeding the threshold of $\pm 2\sigma$. More specifically, we have found the shifted TTs towards nighttime

by finding the amount of shift. The procedure that has been followed is described in Section 3.2. In Figure 6, we give an example of TTM analysis for the 2022 Cyprus EQ. We have found significant shifts of TTs, which exceed the 2σ criterion, searching for anomalies in the same studied period (2 December 2021–29 January 2022), as in the case of the NFM analysis. For the first set of TTs on the morning side towards the day, we calculated the difference of the set of TTs from the running mean window by constructing the dt_{m1} time series (see Section 3.2). The same procedure (see Section 3.2) was applied for the other TT sets by calculating dt_{m2} and dt_e , as well as the dD_{VLF1-1} time series (shift of "VLF daylength"). We note that the symbolization of dD_{VLF1-1} (the same holds for each kind of shifted "VLF daylength", e.g., dD_{VLF2-1}) shows first (reading towards right) the number of sets of TTs that are on the evening side, while the second number indicates the set of TTs that are on the morning side. As we can see, one decrement on 8 January 2022 is observed in the first panel of Figure 6, which is symbolized as dt_{m1} by exceeding (shifted by -40.6 min) the threshold of 2σ of the whole studied period (2 December 2021–29 January 2021). In the second panel for dt_{m2} a shift not related to the EQ anomaly (shifted by -12.2 min) appeared on 8 December 2021. Moreover, in the third panel for dt_e , only one anomalous (shifted by 26.9 min) day on 28 December 2021 existed, while in the fourth panel, presenting D_{VLF1-1} , one increment (shifted by 53.9 min) on 9 January 2022 has been observed.



Figure 6. Temporal evolution of the shifts in the morning and evening TTs, as well as of the "VLF daylength" for the propagation path GERSR during the studied period (2 December 2021 to 29 January 2022). Red solid horizontal lines indicate the corresponding $+2\sigma/-2\sigma$ limits; σ is calculated for the whole studied period for each panel. EQ date is marked at the top of each panel.

In Table 6, we present the overall results of the TTM analysis for the 2022 Cyprus EQ for each of the three stations (UWA, GER, and ACH). This table explicitly shows the date of the anomaly in UT and attributes any observed anomaly (shift) to the 2022 Cyprus EQ or any other possibly ionosphere-influencing extreme event. The anomalies at each shown date are denoted alongside the receiver's call name. It is noted that the selected period for TTM analysis has been chosen similarly to the study of NFM for each station (see Section 4.1). It is evident from Table 6 that pre-seismic indications have been observed from three stations on different dates before the EQ by searching over 15 days. We note that most of these anomalies (anomalous shifts) were identified for the UWA station. It is important to note that one anomalous shift by exceeding the 2σ in D_{VLF1-1} is observed on 21 December 2021, prior to two EQs that occurred close to Crete Island (see Section 2).

Moreover, anomalous dates (shifts), unlikely to be related to the 2022 Cyprus EQ, were found before (7 December 2021), as well as after (24 January 2021), its occurrence (see Table 6). The first anomalous day, before the occurrence of the EQ, is too far from the examined EQ occurrence, while the second anomalous day, after the occurrence of the 2022 Cyprus EQ, cannot be attributed to any other possibly ionosphere-influencing event.

Table 6. Results of TTM analysis for the 2022 Cyprus EQ. UWA, GER, and ACH denote the station where the corresponding anomaly has been identified.

Date of Appearance of the Anomaly (UT)	Excess of -2σ for dt_{m1}	Excess of -2σ for dt_{m2}	Excess of 2σ for dt_e	Excess of 2σ for dD_{VLF1-1}	Possibly Associated Extreme Event(s)
7 December 2021		ACH (shifted -12.1 min)			-
8 December 2021		UWA (shifted -19.6 min), GER (shifted -12.2 min)			-
21 December 2021				UWA (shifted -50.1 min)	b, c
28 December 2021			GER (shifted 26.9 min)		a, c
1 January 2021	UWA (shifted -34.7 min)				a
6 January 2021			ACH (shifted 29.8 min)		а
8 January 2021	UWA (shifted -33.2 min), ACH (shifted -34.5 min), GER (shifted -40.6 min)				a
9 January 2021				UWA (shifted 43.4 min), ACH (shifted 51.3 min), GER (shifted 53.9 min)	a
24 January 2021			UWA (shifted 34 min)		-

The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow 2022$ Cyprus EQ (6.4 M_L , 11 January 2022); $b \rightarrow a 5.5 M_L$ EQ that occurred on 26 December 2021 18:59:02 (UTC) at (35.2029° N, 26.8497° E); $c \rightarrow a 5.7 M_L$ EQ that occurred on 29 December 2021 05:08:09 (UTC) at (34.791° N, 25.139° E).

Similarly, for the 2021 Crete EQs we present the summarized results obtained from the TTM analysis in Table 7. We used the time interval from 2 September 2021 to 29 October 2021, as in the corresponding NFM analysis (see Section 4.1). In this table, we present the shifts found on different dates before each EQ of the 2021 Crete EQs group, searching for possible pre-seismic indications 15 days before each one of them. However, we include all the detected anomalies (shifts) appearing in the studied time interval from 2 September 2021 to 29 October 2021. Thus, on only one date on 4 September 2021, we have observed anomalies (shifts) (see Table 7) from UWA station that exceed the 15 days. Finally, it is profound from Table 7 that most anomalous days come from the UWA station.

Date of Appearance of the Anomaly (UT)	Excess of -2σ for dt_{m1}	Excess of -2σ for dt_{m2}	Excess of 2σ for dt_{e1}	Excess of 2σ for dt_{e2}	Excess of 2σ for dD_{VLF1-1}	Excess of 2σ for dD_{VLF1-2}	Excess of 2σ for dD_{VLF2-1}	Excess of 2σ for dD_{VLF2-2}	Possibly Associated Extreme Event(s)
4 September 2021				UWA (shifted 34.5 min)			UWA (shifted 35 min)	UWA (shifted 35.83 min)	-
6 September 2021				GER (shifted 52.5 min)			GER (shifted 57 min)	GER (shifted 47.75 min)	а
19 September 2021						UWA (shifted 28.4 min)			а
20 September 2021		UWA (shifted -13.7 min), GER (shifted -20.1 min)							a
28 September 2021		_0.11 1111.)	UWA (shifted 26 min)			UWA (shifted 26.3 min)			b
1 October 2021			GER (shifted -18.75 min)	UWA (shifted 36.9 min)			UWA (shifted 32.5 min)	UWA (shifted 39.1 min)	b
2 October 2021		GER (shifted -35.66 min)		,		GER (shifted 44.16 min)		GER (shifted 38 min)	b
7 October 2021	UWA (shifted -45 min), GER (shifted -43.8 min)				UWA (shifted 57.6 min)				b, c
11 October 2021	GER (shifted -32 min)		UWA (shifted 26 min)		GER (shifted 48.7 min)				b, c
12 October 2021	·	UWA (shifted -10.87 min)							b, c

Table 7. Results of TTM analysis for the 2021 Crete EQs. UWA and GER denote the station where the corresponding anomaly has been identified.

The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow 1$ st 2021 Crete EQ (5.8 M_L , 27 September 2021); $b \rightarrow 2$ nd 2021 Crete EQ (6.3 M_L , 12 October 2021); $c \rightarrow 3$ rd 2021 Crete EQ (6.1 M_L , 19 October 2021).

4.3. NT Analysis Results

In this section, we present the results obtained by applying the NT analysis method (see Section 3.3) to the non-normalized VLF/LF propagation quantities (*TR*, *DP*, and *NF*) time series. We applied the NT analysis method to these non-normalized statistical quantities, as first presented for the 2020 Samos EQ and 2016 Kumamoto EQs [8,15]. The procedure to apply the NT analysis in VLF sub-ionospheric data is the same as in the case of the seismicity time series (see Section 3.3). Specifically, we consider all of the daily values of the VLF/LF propagation quantity under analysis that are higher than a certain threshold to be the "events" considered during the NT analysis, while for the k-th event Q_k is considered to be equal to its corresponding daily value.

In this research, we are looking for criticality signatures that are associated with the EQ in the three statistical quantities (*TR*, *DP*, and *NF*) by presenting the results of the NT analysis method for the three stations (UWA, GER, and ACH), taking only data received by the transmitter ISR. Figure 7 presents an example of NT analysis (applied on the TR VLF propagation quantity) for the ISR-ACH sub-ionospheric propagation path prior to the 2022 Cyprus EQ. The analyzed period is from 18 December 2021 to 29 January 2022, as in the corresponding NFM analysis (see Section 4.1). The criticality conditions of NT analysis are satisfied for four thresholds, as presented in the (four) corresponding panels of Figure 7. As we can observe from this figure, criticality has been reached on 8 January 2022, which is very close in time to the examined EQ. In general, the specific criticality indication could have been considered as contaminated by a moderate geomagnetic storm (min Dst -91 nT) that happened on 14 January 2022 (see Figure 3). Yet, if one accounts for the magnitude of the EQ under study, the detected criticality is most probably reflecting the preparation of the 2022 Cyprus EQ than that of the aforementioned moderate geomagnetic storm, let alone all other less-important geomagnetic storms that happened up to ~15 days after the criticality date (see Figure 3). Finally, since only nighttime data have been used, the detected criticality has nothing to do with the preparation of any solar flare.

At this point, we should mention that criticality indications were also found after the date of the 2022 Cyprus EQ in the data acquired from all three receivers. The source of these criticality indications is not so clear, since these could be attributed either to post-EQ power laws due to "locally surviving" critical dynamics, despite the unstable critical point no longer existing, as explained in detail in [8], or to some geomagnetic events that continued up to the first days of February 2022, even though these were not that intense (see Figure 3).

In Table 8, we summarize all our findings obtained with the NT analysis in studying the 2022 Cyprus EQ. In this table, we have included all criticality signatures found by analysis of the three above-mentioned VLF/LF propagation quantities. Specifically, Table 8 indicates for each analyzed VLF/LF propagation quantity, at which receiver(s) and for which date a criticality signature has been identified, as well as the ionosphere-influencing extreme event(s) possibly related to each critical signature. As we can see from Table 8, critical signatures have been found for all three stations. We have marked each critical signature using the call name of the receiver (UWA, GER, or ACH). We should note at this point that the time range used for the NT analysis of the stations UWA and GER has been changed compared to the corresponding period used for NFM analysis (see Section 4.1). Specifically, we chose the shorter time period from 18 December 2021 to 29 January 2022 that was used during the NFM analysis of ACH. This was chosen because if one uses the wider time period (2 December 2021–29 December 2021), the NT analysis does not indicate clear criticality signatures, probably due to a masking phenomenon owing to the two EQ events that happened near Crete Island on 26 and 29 December 2021 (see Section 2). Generally, from Table 8 it can be seen that criticality signatures appear on different dates before the 2022 Cyprus EQ within 15 days prior to the EQ occurrence.



Figure 7. NT analysis of the *TR* VLF propagation quantity time series of the propagation path ISR–ACH for the examined period (18 December 2021–29 January 2022). The presented temporal variations of the NT parameters correspond to the different thresholds of TR_{Th} : (a) 2, (b) 2.5, (c) 3, (d) 3.5, respectively. The limit value of the entropy appears as a horizontal solid light green line, while the value 0.07, along with the region around it, are denoted by a horizontal solid grey and two horizontal dashed grey lines, respectively. The $\langle D \rangle$ limit is shown as a horizontal brown line. The light green patch in each panel indicates the presence of criticality according to the conditions set by NT analysis. The events in each panel depend on the corresponding threshold. Moreover, although the conventional time (date) of occurrence of each corresponding event is noted in the x-axis tick values, the x-axis scale actually follows the NT representation; for this reason, the x-axis is not linear in conventional time. Date format is day/month (for dates that belong to December the year is 2021, while for January the year is 2022).

Date of Appearance of the Criticality Signature (UT)	NT Analysis on TR	NT Analysis on DP	NT Analysis on NF	Possibly Associated Extreme Event(s)
29 December 2021		UWA		а
30 December 2021		GER	ACH	a
4 January 2022			ACH	a
7 January 2022	UWA, GER			a
8 January 2022	ACH		UWA	а
10 January 2022		ACH		а
15 January 2022	UWA, GER			b, c, d, e
20 January 2022	GER	GER		f, g
22 January 2022	ACH	GER		f, g
23 January 2022	GER, ACH	GER		f, g
24 January 2022	ACH			f, g
25 January 2022	GER, ACH		UWA	f, g

Table 8. Results of the NT analysis for the 2022 Cyprus EQ. UWA, GER, and ACH denote the station where the corresponding criticality has been identified.

The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow 2022$ Cyprus EQ (6.4 M_{L} , 11 January 2022); $b \rightarrow post-EQ$ power laws due to "locally surviving" critical dynamics, despite the unstable critical point no longer existing; $c \rightarrow a$ minor (Kp~5) geomagnetic storm that occurred on 15 January 2022 21:00:00 (UT); $d \rightarrow a$ minor (Kp = 5) geomagnetic storm that occurred on 18 January 2022 00:00:00 (UT); $e \rightarrow a$ minor (Kp = 5) geomagnetic storm that occurred on 19 January 2022 03:00:00 (UT); $f \rightarrow a$ minor (min Dst = -66 nT) geomagnetic storm that occurred on 3 February 2022 10:00:00 (UT); $g \rightarrow a$ minor (min Dst = -61 nT) geomagnetic storm that occurred on 4 February 2022 20:00:00 (UT).

Similar to the 2022 Cyprus EQ, in Table 9, we present the results of the NT analysis for the 2021 Crete EQs. Specifically, we analyzed the VLF propagation quantities (see Section 3.1) from the UWA and GER stations by keeping the same time period that we used in the analysis of NFM. Criticality indications that appeared on different dates have been found only before the first 2021 Crete EQ within 15 days before its occurrence. A minor (min Dst = -64 nT) geomagnetic storm that occurred on 17 September 2022 21:00 (UT) is not likely to have influenced the results if one accounts for the magnitude of the first 2021 Crete EQ.

Table 9. Results of the NT analysis for the 2021 Crete EQs. UWA and GER denote the station where the corresponding criticality has been identified.

Date of Appearance of the Criticality Signature (UT)	NT Analysis on TR	NT Analysis on DP	NT Analysis on NF	Possibly Associated Extreme Event(s)
9 September 2021	GER			а
10 September 2021		UWA, GER		а
11 September 2021	UWA			a
12 September 2021	UWA		UWA	а
19 September 2021			GER	а

The following letter appearing in the column "Possibly associated extreme event(s)" denotes: $a \rightarrow 1$ st 2021 Crete EQ (5.8 M_L , 27 September 2021).

4.4. Wavelet Analysis Results

Following the procedure of applying the wavelet analysis in sub-ionospheric propagation data as described in Section 3.4, here we present the analysis results for the 2021 Crete EQs and the 2022 Cyprus EQ (see Section 3.1). In this analysis, we compute the WPS of the nighttime amplitude recordings by taking recorded data by the three stations (UWA, GER, and ACH). Specifically, we chose the same nighttime interval (21:00–02:00 UT) as we have already used in the NFM analysis (see Section 4.1). For determining the date of each WPS with the selected nighttime interval, we follow the same reasoning as in NFM (see Section 4.1).

In Figure 8, we give an example of analysis for the 2022 Cyprus EQ using the UWA receiver's recordings. Each panel of this figure clearly shows the WPS of VLF nighttime amplitude recordings for each day, covering a period of 16 days (16 panels). Considering this figure, one significant anomaly correlated with pre-seismic AGW activity has been found 11 days (on 31 December 2021) before the EQ (the date marked at the top of the panel is colored as red). This observed enhancement in the power of WPS was found to have a periodicity of 31 to 54 min and lies within the CoI. However, more wave-like structures of intermediate intensity (and within the CoI) were observed at different dates (27 December 2021, 2, 10 and 11 January 2022), which are also possibly related to the EQ.

All the wavelet analysis findings for 2022 Cyprus EQ are summarized in Table 10, including all the results from the three stations (UWA, GER, and ACH). This table, specifically, includes the date and time of the occurrence of the AGW anomaly, its range of periodic structures, and its intensity. Furthermore, this table shows the attribution of any detected AGW anomaly to possibly ionosphere-influencing extreme event(s). However, from Table 10 we can observe only the AGW anomalies that appeared before the 2022 Cyprus EQ at dates where no other possibly ionosphere-influencing extreme event occurred, except from one case for which the anomaly could be due to another EQ of 5.7 M_L that occurred on 29 December 2021 close to Crete Island.

Finally, in Table 11, we show the results of the wavelet analysis for the 2021 Crete EQs, in a form similar to the form of Table 10. Again, it can be observed that many AGW anomalies have been detected on different dates before each one of the EQs under study. However, none of these AGW anomalies are related to any other possibly ionosphere-influencing extreme event.



Figure 8. Wavelet power spectrum (WPS) of UWA station. Horizontal axis refers to time of the night after 21:00 UT. The white parabolic line is the cone of influence (CoI). Vertical axis refers to periodicity in minutes and color corresponds to the normalized power (magnitude of periodic structure). The date (11 January 2022) of the EQ is marked with red color at the top of the last panel.

Table 10	. Results	of wavelet	analysis f	for the	2022 C	yprus EQ.
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Receiver Call Name	Date of Appearance of the Anomaly (UT)	Time of Appearance of the Anomaly (UT)	Range of Periodic Structures (min.)	Intensity of the Wave-Like Structures	Possibly Associated Extreme Event(s)
	27 December 2021	21:45:57-22:24:00	31–36	moderate	a, b
	31 December 2021	22:10:48-23:21:00	31–54	high	а
UWA	2 January 2021	00:07:48-00:54:00	31–38	moderate	а
	10 January 2021	21:43:48-23:07:48	31–50	moderate	а
	11 January 2021	21:45:36-22:27:00	31–38	moderate	а
GER	5 January 2022	21:37:48-22:12:36	31–38	high	a
	9 January 2022	23:48:00-00:25:48	31–36	moderate	а
ACH	10 January 2022	00:24:36-01:12:00	31–41	high	а

The terms "high" and "moderate" in the column "Characterization of the intensity of power of the wave-like structures" verbally characterize the intensity of the wave-like structures with reference to the normalized linear scale of power (33–66% for "moderate" and 67–100% for "high"). The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow 2022$ Cyprus EQ; $b \rightarrow a 5.7 M_L$ EQ that occurred on 29 December 2021 05:08:09 (UTC) at (34.791° N, 25.139° E).

Table 11. Results of wavelet analysis for the 2021 Crete EQs.

Receiver Call Name	Date of Appearance of the Anomaly	Time of Appearance of the Anomaly (UT)	Range of Periodic Structures (min.)	Intensity of the Wave-Like Structures	Possibly Associated Extreme Event(s)
	13 September 2021	22:48:36-00:22:48	31–47	moderate	а
	17 September 2021	22:52:48-00:43:48	31–36	high	а
	17 September 2021	23:21:00-00:36:00	47-71	moderate	а
	4 October 2021	22:07:48-00:31:48	31–38	high	b, c
	8 October 2021	23:21:36-00:27:00	31–36	high	b, c
ΙIWA	8 October 2021	22:54:36-00:01:48	71–94	high	b, c
01111	9 October 2021	23:45:36-00:57:00	31-50	moderate	b, c
	10 October 2021	21:58:48-00:55:48	41–47	moderate	b, c
	11 October 2021	23:46:48-01:09:00	36-54	high	b, c
	12 October 2021	23:04:48-00:21:00	41-62	moderate	b, c
	13 October 2021	23:46:48-01:49:48	54-66	moderate	с
	16 October 2021	22:09:36-23:18:36	31–33	moderate	с

Receiver Call Name	Date of Appearance of the Anomaly	Time of Appearance of the Anomaly (UT)	Range of Periodic Structures (min.)	Intensity of the Wave-Like Structures	Possibly Associated Extreme Event(s)
GER	12 September 2021	00:00:00-00:45:36	31–36	high	а
	28 September 2021	22:43:48-00:09:36	31-50	high	b
	28 September 2021	22:58:48-00:22:48	57-87	moderate	b
	4 October 2021	23:03:36-00:24:36	35–43	high	b, c
	9 October 2021	23:36:36-01:00:00	43-66	high	b, c
	10 October 2021	00:03:00-01:09:36	33–46	moderate	b, c
	12 October 2021	23:33:36-00:18:00	33–44	moderate	b, c
	13 October 2021	22:36:00-23:54:36	31–41	moderate	с
	16 October 2021	21:42:36-00:09:00	31–35	high	с
	19 October 2021	21:43:48-00:33:36	31–38	high	с

Table 11. Cont.

The terms "high" and "moderate" in the column "Characterization of the intensity of power of the wave-like structures" verbally characterize the intensity of the wave-like structures with reference to the normalized linear scale of power (33–66% for "moderate" and 67–100% for "high"). The following letters appearing in the column "Possibly associated extreme event(s)" denote, respectively: $a \rightarrow 1st 2021$ Crete EQ (5.8 M_L , 27 September 2021); $b \rightarrow 2nd 2021$ Crete EQ (6.3 M_L , 12 October 2021); $c \rightarrow 3rd 2021$ Crete EQ (6.1 M_L , 19 October 2021).

4.5. Simulation Results form Long Wavelength Propagation Capability Code

We present the simulation results in two parts. As mentioned above, we have focused on the morning and evening terminator times for the simulation. Therefore, we first apply the different sets around the SRT and SST to obtain the simulated amplitude value at the time of the terminators. In previous reports, e.g., [17], the simulation has been conducted to check for the entire period of sunrise and sunset times and a comparison has been presented. After that, the primary computation is associated with the perturbed ED profile during the TTs.

We used the different sets of h' and β , as mentioned in Section 3.5, for the entire periods of observation in Equation (4) to obtain the altitude profile of ED. Figures 9–11 show the altitude profile of ionospheric ED at SRT1, SST1, and SST2, respectively, for the ISR–UWA path in respect to the 2021 Crete EQs. The "QD" is a quiet day without EQ or space weather events. It is evident from Figure 9 that the ED decreases due to the presence of an EQ. The amount of changes in the ED profiles is not fixed. For the first 2021 Crete EQ on 27 September 2021, a moderate shift is observed on 24, 25 and 27 September 2021. It is noted that the ED profile tends to decrease before the EQ, and after the EQ, it goes in the opposite direction, leading to the quiet condition of QD. These findings corroborate the previous results when a perturbed ED usually tries to acquire its normal state (see the Introduction section). For the third 2021 Crete EQ on 19 October 2021, the changes in the ED are more significant, and the maximum decrease in ED is observed on 16 October 2021, 17 October 2021, and also on the EQ day on 19 October 2021. Therefore, both pre- and co-seismic effects are present in this case.

Figures 10 and 11 present similar ED profiles for SST1 and SST2 for the ISR–UWA path. For the sunset times, we examine both the sunset terminators. In contrast to the SRT, for SST1, we found significantly perturbed ED profiles only for the second 2021 Crete EQ on 12 October 2021. The maximum shifts in ED take place on 11 October 2021 and 12 October 2021. It is evident that the ED profile for SST1 returns to its original values after the EQ on 12 October 2021, and no further significant changes have been observed. For the SST2, the changes in the ED are random with no such prominent decreasing nature.

The variation of ED for the ISR–GER path with respect to the 2021 Crete EQs is shown in Figures 12–14. For the ISR–GER path, we examine one SRT (SRT1) and two SSTs (SST1 and SST2), as for the UWA signal. For the 27 September 2021 EQ, no changes in ED have been observed. However, for the EQ on 12 October 2021, significant changes in the ED profile are observed on 11 October 2021 from its typical values, while for the 19 October 2021 EQ, a greater decrement in ED is observed on the day of the EQ.



Figure 9. Altitude profile of ionospheric ED around the time of SRT1 obtained from ISR–UWA path in respect to the 2021 Crete EQs. The y–axis and x–axis denote the ionospheric height in km and the ED in logarithmic scale, respectively. "QD" and "EQ" lines represent the variation of normal days and the EQ day, respectively.



Figure 10. Same as Figure 9 for SST1 of ISR–UWA path.



Figure 11. Same as Figure 9 for SST2 of ISR–UWA path.



Figure 12. Altitude profile of ionospheric ED around the time of SRT1 obtained from ISR–GER path with respect to the 2021 Crete EQs. The *y*-axis and *x*-axis denote the ionospheric height in km and the ED in logarithmic scale, respectively. "QD" and "EQ" lines represent the variation of normal days and the EQ day, respectively.



Figure 13. Same as Figure 12 for SST1 of ISR–GER path.



Figure 14. Same as Figure 12 for SST2 of ISR–GER path.

In the case of SSTs, it is evident from Figure 13 that the ED profiles are a bit random for this terminator. For the EQ on 27 September 2021, the values of ED at SST1 are found to be unperturbed both on the EQ day and prior to the same. This is consistent with the ISR path ED profile during SST, where we failed to observe any significant changes in ED. However, for the EQs on 12 October 2021 and 19 October 2021, substantial changes in the ED profile are observed. For the 12 October 2021 EQ, the maximum changes in ED are observed on the day of the EQ, followed by intermediate changes on 11 October 2021. Similarly, for the 19 October 2021 EQ, significant changes in ED are found on the day of the EQ, followed by intermediate changes on 14 October 2021.

For the SST2 variation, we observed only significant changes for the EQ on 27 September 2021. However, we observed the change long back on 13 September 2021 and 15 September 2021. This is a bit unusual compared to the other ED profiles for which we found the changes in ED within a shorter period of time before the EQ. It is also evident from Figure 13 that there are examples of enhancement of ED in the intermediate periods of those EQs. This effect can be attributed to three consecutive EQs during the observation period. We also observed the shift in TT in opposite directions. Therefore, the modal conversion of the waves possibly behaves oppositely. It is well-established that during the SRTs and SSTs, a conventional waveguide mode conversion occurs due to variation in the ED profiles due to solar irradiance. As the seismogenic perturbation is not controlled but random both ways, TT shifts have been observed in many previous cases, but it is to be noted that the perturbation in ED primarily decreases consistently during or before the EQ.

The similar variation of the altitude profiles of ED at SRT and SST for the ISR–UWA path during the 2022 Cyprus EQ on 11 January 2022 is shown in Figures 15 and 16, respectively. It is evident from Figure 15 that a considerable amount of decrement of ED is observed on the day of the EQ (11 January 2022), and intermediate decrease in EDs from the quiet day ED profiles (QD) are observed on 1, 8 and 9 January 2022. For the SST, a similar decrement in the ED profiles has been observed before the EQ day (on 5, 6, 7 and 9 January 2022), while for the day of the EQ, the ED profiles slightly increased. This is quite an exciting result where the co-seismic effect is somewhat opposite to the pre-seismic variation.



Figure 15. Altitude profile of ionospheric ED around the time of SRT obtained from ISR–UWA path with respect to the 2022 Cyprus EQ. The *y*-axis and *x*-axis denote the ionospheric height in km and the ED in logarithmic scale, respectively. "QD" and "EQ" lines represent the variation of normal days and the EQ day, respectively.



Figure 16. Same as Figure 15 for SST of ISR–UWA path.

4.6. Summary of Analysis Results

By taking into account all the results presented in the previous ubsections of Section 4, in this Subsection we summarize the main findings of the whole research:

Within 15 days before each studied EQ (2022 Cyprus EQ and 2021 Crete EQs), anomalies and criticality signatures have been revealed for different days from all the analysis methods and techniques (see Tables 2–11). It is noted that any possible influence on the presented results from ionosphere-influencing extreme events (e.g., solar flares, geomagnetic storms, EQs, volcanoes), other than the EQs under study have been discussed in detail, whereas any possible attribution of the revealed anomalies and criticality signatures to them has been included in Tables 2–11.

By analyzing with NFM, taking data from the three sub-ionospheric propagation paths (ISR-UWA, ISR-GER, ISR-ACH) with respect to the 2022 Cyprus EQ, we can observe several VLF anomalies exceeding the $+2\sigma/-2\sigma$ limits (see Table 2). These are found in all the sub-ionospheric propagation paths in the three VLF propagation quantities. However, two other EQ events (occurred on 26 December 2021 and 29 December 2021, respectively) that also occurred in the southeastern Mediterranean, not the focus of this study, are found to be correlated with only two anomalous days on 11 December 2021 and 21 December 2021 (see Table 2). Moreover, it is also found that two anomalous days (2 December 2021 and 27 December 2021) are not associated with any possibly ionosphere-influencing extreme event. Similar to the 2022 Cyprus EQ case, there are VLF anomalies attributed to the three 2021 Crete EQs in the three VLF propagation quantities (exceeding the $+2\sigma/-2\sigma$ limits) for the two sub-ionospheric propagation paths for which adequate data were available (ISR-UWA and ISR-GER) (Table 3). In this case, no other extreme event occurred around the studied time period. However, two detected anomalous days (5 September 2021 and 27 October 2021) in Table 3 are not attributed to any known possibly ionosphere-influencing extreme event.

From the sequential plots of the diurnal variation of the received amplitude in regard to the 2022 Cyprus EQ, we have observed pre-EQ shifts of TT minima towards the nighttime in all studied sub-ionospheric propagation paths (ISR–UWA, ISR–GER, ISR–ACH) for two out of the three TTs (t_{m1} and t_e) (Table 4). For the third TT (t_{m2}) there is no observed shift. Moreover, one of the two EQs not the focus of this study (mentioned above) is found to be possibly related to one shifted day (28 December 2021). In a more detailed analysis, by applying the TTM (see Table 6), it is evident that shifts (also noting the amount of shift) of TTs exceeding $\pm 2\sigma$ are found in all three paths for the three TT sets (t_{m1} , t_{m2} , and t_e). However, three observed anomalous shifts, detected earlier and later than the EQ on 7 December 2021, 8 December 2021, and 24 January 2021, are not linked with any extreme event. Yet, the two aforementioned EQs that are not the focus of this study are likely to be associated with two anomalous days with 21 December 2021, while one of them is associated with 28 December 2021, by pointing out that their occurrence is closer in time to these observed anomalies, in contrast to the examined EQ. From the sequential plots of the diurnal variation of the received amplitude in regard to the 2021 Crete EQs (see Table 5), we found pre-EQ shifts of TT minima towards nighttime in four TT sets (t_{m1} , t_{m2} , t_{e1} , and t_{e2}) by observing shifts of TT minima for the paths ISR–UWA and ISR–GER. On the other hand, using the TTM, shifts were also detected for both paths (Table 7), while only one shifted day on 4 September 2021 is not associated with an EQ.

By applying the NT analysis on non-normalized VLF propagation quantities in the case of 2022 Cyprus EQ, it is found that the criticality signatures are detected at different dates before and after the EQ for all three propagation paths (ISR–UWA, ISR–GER, ISR–ACH) (see Table 8). More specifically, the criticality signatures that appeared before the EQ can be attributed to the specific seismic event under study, while the source of the criticality indications found after the 2022 Cyprus EQ (during the time period 15 January 2022–25 January 2022) is not so clear. These could be attributed either to post-EQ power laws [8] due to "locally surviving" critical dynamics, despite the unstable critical point no longer existing, or to some geomagnetic events, even though these were not very intense. In the case of applying the NT analysis method to the non-normalized VLF propagation quantities for the 2021 Crete EQs, criticality has been detected in both paths for which adequate data were available (ISR–UWA and ISR–GER), but only for the first EQ that occurred on 27 September 2021 (see Table 9).

Regarding the wavelet analysis (see Table 10) of VLF data for the 2022 Cyprus EQ, it is found that AGW-related anomalies existed in all three propagation paths on different days before the studied EQ. High- and moderate-intensity patches of wave-like structures are generally unveiled in a range of ~30–60 min, with varying time intervals. At this point, it needs to be mentioned that only one AGW-related anomaly is associated with the EQ that occurred in the southeastern Mediterranean on 27 December 2021 that was not the focus of this study as presented in Table 10. Similarly, by analyzing the VLF data for the 2021 Crete EQs, AGW activity appeared at different dates before each studied EQ in the propagation paths ISR–UWA and ISR–GER. The detected high- and moderate-intensity patches of wave-like structures appeared, most of them from ~30–60 min, while very few exceed the periodicity of ~60 min. It is finally mentioned that no other possibly ionosphere-influencing extreme event is associated with these observed AGW anomalies.

The LWPC simulation of perturbed ED finds significant variation during the SRT and SST times for the four EQs that are the focus of this study. In Figures 9–16, it is evident that, for all the EQs, the electron density becomes mostly depleted. Moreover, it needs to be mentioned that we have observed both pre- and co-seismic perturbation of ED profiles for these EQs. The changes vary from roughly 0.5×10^9 to 3×10^9 /m³. Some exciting features are found in this analysis. The ED profiles during SRT seem more prominently perturbed compared to the SST. For the Crete EQs, not all the paths show identical ED profiles for SRT and SSTs. The usual ED perturbation for SRT lies within 0 to 7 days before the EQs (Figures 9, 10, 12 and 13). On the other hand, during the SST2, the UWA signal does not have any effect, while for the GER signal, SST2 shows a variation quite a long time (14 days) before the EQ (Figures 11 and 14). However, for the 2022 Cyprus EQ, the ED variation shows an anomalous decrement for both the SRT and SST period. We anticipate this inhomogeneity due to different modal interference processes at different stations.

5. Conclusions

This paper elaborately presented the analysis of VLF/LF sub-ionospheric propagation data from one VLF transmitter (with call name ISR and operation at 29.7 kHz). At the same time, their acquisition was performed by three VLF/LF receivers located in Athens (Greece). In this attempt, we used several analysis methods in order to detect pre-seismic indications in a time range of 15 days before the 2021 Crete EQs and the 2022 Cyprus EQ. This aim was indeed achieved, confirming previous studies

On the other hand, except the EQs in the focus of this study, several other possibly ionosphere-influencing extreme events have been found close within a relatively short time distance (within 15 days) from the revealed anomalies and criticality indications. For this reason, we attributed any possible candidate extreme event to any of each possible aforementioned anomaly/criticality indication.

Moreover, it was found that, in some cases, the identified anomalies or criticality signatures were revealed after the occurrence of an EQ. This could be due to other possible ionosphere-influencing extreme events that occurred within a relatively short time distance (~15 days) after the detected anomalies or criticality signatures, or, for the case of criticality indications, to possible post-EQ power laws due to "locally surviving" critical dynamics, despite the unstable critical point no longer existing [8].

Moreover, by applying the LWPC simulation for TT minima, using data from the three VLF/LF stations, we found perturbations (increases or decreases) of ED for all the studied EQs. Past studies, e.g., [17,25,37], which similarly showed ED perturbations on shifted TTs, are in agreement with the findings of the present work.

To sum up, this paper clearly shows that the lower ionosphere appears to carry a pre-seismic imprint before the occurrence of each studied EQ, using various analysis methods. However, more research efforts are needed in the future in order to find more evidence to attribute an anomaly to a specific extreme event by taking a further step towards understanding the nature of the phenomena preceding an impending EQ.

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