

## Article

# Reduction of the VLF Signal Phase Noise Before Earthquakes

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**Abstract:** In this paper we analyse temporal variations of the phase of a very low frequency (VLF) signal, used for the lower ionosphere monitoring, in periods around four earthquakes (EQs) with magnitude greater than 4. We provide two analyses in time and frequency domains. First, we analyse time evolution of the phase noise. And second, we examine variations of the frequency spectrum using Fast Fourier Transform (FFT) in order to detect hydrodynamic wave excitations and attenuations. This study follows a previous investigation which indicated the noise amplitude reduction, and excitations and attenuations of the hydrodynamic waves less than one hour before the considered EQ events as a new potential ionospheric precursors of earthquakes. We analyse the phase of the ICV VLF transmitter signal emitted in Italy recorded in Serbia in time periods around four earthquakes occurred on 3, 4 and 9 November 2010 which are the most intensive earthquakes analysed in the previous study. The obtained results indicate very similar changes in the noise of phase and amplitude, and show an agreement in recorded acoustic wave excitations. However, properties in the obtained wave attenuation characteristics are different for these two signal parameters.

**Keywords:** ionosphere; earthquakes; observations; VLF signal; signal processing; acoustic and gravity waves



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## 1. Introduction

In addition to periodical ionospheric changes, which can be predicted and estimated by different models (see, for example, [1,2] and references therein), sudden events can induce significant ionospheric disturbances and affect many contemporary technologies based on satellite and ground-based electromagnetic (EM) signal propagation [3]. Consequently, variations of the recorded EM signal properties can be used for detections and analyses of influences of many phenomena on this atmospheric layer including processes which induce different kinds of natural disasters [4–8].

In the last several decades, studies of the lower ionosphere disturbances are mostly based on observations by very low/low frequency (VLF/LF) radio signals [9–13] and processing of the corresponding recorded data in both the time and frequency domains. Increases or decreases of the signal amplitude and/or phase are recorded in many studies focused on research of ionospheric disturbances induced by earthquakes [9], solar activity [14–18], tropical cyclones [19,20], solar eclipse [21,22] etc.

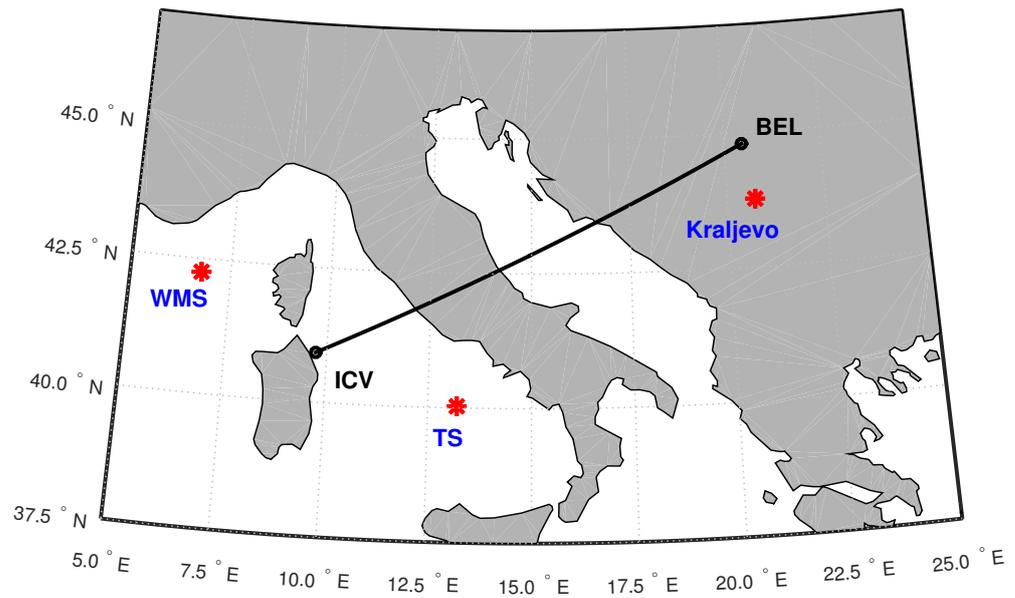
The acoustic gravity waves (AGW) are recently mentioned in association with earthquakes mainly as a result of the strong oscillations caused by the seismic shock (Jin et al., 2015 [23]) or tsunami (Manta et al., 2020 [24]). But for a long period of time they were considered as a main agent of the seismo-ionospheric coupling as the earthquake precursor (Korepanov et al., 2009 [25]). But the main problem was the lack of convincing experimental evidences of the AGW generation before earthquakes and corresponding physical mechanism of their generation. As the possible sources the mosaic distribution of gas emission before earthquakes was proposed (Mareev et al., 2002 [26]), or the surface thermal anomalies (Molchanov et al., 2004 [27]). Research of AGW in the lower ionosphere is also presented in several studies. They relate to disturbances induced by the solar terminator [28], geomagnetic storms [29,30], tropical cyclones [19,31,32] earthquakes [33], solar eclipse [34], and they are based on analyses of the VLF/LF signals.

Studies of the ionospheric changes as precursors of earthquakes report changes usually a few days before events [9,13,35]. These lower ionosphere disturbances are detected as the solar terminator shift [13,36,37] and deviations of the time evolutions of signal characteristics from values recorded during days with unperturbed conditions [5,38] (in time domain), and as variations in the wavelet power spectrum [5,39] (in frequency domain), and all of these changes are shown for both signal amplitude and phase. The recent analysis presented in [33] shows reduction of the amplitude noise of the VLF signal less than one hour before the earthquake occurred near Kraljevo, Serbia, on 3 November 2010 (the seismotectonic model of this event is presented in [40]), as well as excitation and attenuation of the acoustic waves. In addition, the similar changes in the amplitude noise are also recorded for 12 other earthquakes with different magnitudes during three whole days. It was concluded that all considered EQs with the magnitude larger than 4 were connected with the recorded noise amplitude reduction. However, contrary to the previous cases this pioneer study which indicates a possible new ionospheric precursor of earthquake is provided only for the amplitude i.e., variations of the phase, is not considered. For this reason, in this study we extend the research presented in [33] and investigate if the recorded changes in amplitude noise are also visible in analysis of the VLF signal phase, and if excitations and attenuations of the acoustic and gravity waves can be visualized from the recorded phase. We show analysis of the phase of the ICV signal emitted in Italy and recorded in Serbia in periods around four EQs with magnitude greater than 4 which are connected with the noise amplitude reductions in study shown in [33].

The paper is organized as follows. Descriptions of observations and data processing are given in Section 2. Results of this study are divided in two parts: those related to reduction of the phase noise is shown in Section 3.1 and those related to analysis of the acoustic and gravity waves are presented in Section 3.2. Finally, conclusions of this study are summarized in Section 4.

## 2. Observations and Data Processing

In this study we analyse phase of the ICV signal emitted by a transmitter located in Isola di Tavolara, Italy (40.92 N, 9.73 E) and received in Belgrade, Serbia (44.8 N, 20.4 E) in time periods around four EQs, considered in [33] and connected to the short-term noise amplitude reduction. Two of these events occurred near Kraljevo, Serbia, one in the Tyrrhenian Sea (TS) and one in the Western Mediterranean Sea (WMS) (their epicentres are shown in map given in Figure 1). As it can be seen in Table 1, their magnitudes were greater than 5 for two events (the first one near Kraljevo, and in the Tyrrhenian Sea) while other two had magnitude larger than 4. These magnitudes are larger than those related to other EQs considered in [33].



**Figure 1.** Propagation paths of the VLF signals recorded by the Belgrade receiver station (BEL) in Serbia and emitted by the transmitters ICV in Italy. Locations of the main considered EQs are shown as stars. One EQ occurred in the Tyrrhenian Sea (TS) and Western Mediterranean Sea (WMS), while two EQs were near Kraljevo (their epicentres are shown by the same star).

**Table 1.** List of the main earthquakes considered in this study. EQ date, time *t*, epicentre locations (latitude (LAT) and longitude (LON)) and magnitudes (M) are given in <http://www.emsc-csem.org/Earthquake/> (accessed on 25 February 2021). The variable *d* denotes the distance between the EQ epicentres and signal propagation path.

No.	Date	t (UTC)	LAT (°)	LON (°)	<i>d</i> (km)	M	Location
<b>Kraljevo—03/11/2010</b>							
1	2010/11/03	00:56:54	43.74	20.69	126.0	5.4	Serbia (near Kraljevo)
31 EQs until 8 UT—31 in Serbia, 3 in Italy and 1 in Bosnia and Herzegovina							
<b>Tyrrhenian Sea (TS)—03/11/2010</b>							
2	2010-11-03	17:12:30	42.4	13.35	11.4	2	Central Italy
	2010-11-03	17:48:04	43.75	20.7	120.7	2.5	Serbia (near Kraljevo)
	2010/11/03	18:13:10	40.03	13.2	219.1	5.1	TS
	2010-11-03	18:47:23	43.73	20.67	121.7	2.1	Serbia (near Kraljevo)
<b>Kraljevo—04/11/2010</b>							
1	2010-11-04	20:33:01	43.75	20.7	120.7	1.9	Serbia (near Kraljevo)
	2010/11/04	21:09:05	43.78	20.62	114.9	4.4	Serbia (near Kraljevo)
	2010-11-04	21:55:40	45.81	7.55	562.9	1.2	Northern Italy
	2010-11-04	23:43:05	43.78	20.62	114.9	3.3	Serbia (near Kraljevo)
	2010-11-05	00:16:14	43.74	20.64	119.6	2.8	Serbia (near Kraljevo)
	2010-11-05	01:38:48	43.76	20.69	119.4	2.5	Serbia (near Kraljevo)
<b>Western Mediterranean Sea (WMS)—03/11/2010</b>							
4	2010-11-09	16:45:13	43.59	12.36	165.9	2.3	Central Italy
	2010-11-09	18:23:36	42.25	6.77	287.7	4.3	WMS

During the considered four time intervals, additional weaker EQs also occurred near the considered signal propagation path. After the most intensive EQ (Kraljevo, 3/11/2010) 31 additional EQs occurred before 8 UT. 27 of these events were in Serbia, 3 in Italy and 1 in Bosnia and Herzegovina. Their magnitudes were lower than 3 except in one case when it was 3.3. Because of this large number we only give common information of their occurrences in Table 1. Processes in the lithosphere below the monitored ionospheric area were not so intensive during the other three time intervals. For this reason all additional EQ events are indicated in Table 1.

In this analysis, we process the 0.1-s resolution datasets. This procedure consists of three parts: (1) phase unwrapping, (2) determination of the unwrapped phase noise, and

(3) application of Fast Fourier Transform (FFT) to the unwrapped phase in order to examine excitations and attenuations of the acoustic and gravity waves.

- **Phase unwrapping.** The recorded signal phase  $P_{\text{data}}$  represents the deviation of the signal phase with respect to the phase generated at the receiver. Because of that it has a component of constant slope. However, this component does not affect the presented analysis, and for this reason we did not remove it. On the other side, all recorded values are given within a principal phase interval and for further analysis it is necessary to unwrap it. The obtained time evolutions of the unwrapped phase  $P$  are shown in Figure 2 where the vertical lines indicate times of EQ occurrences. Red lines represent the main EQ considered in corresponding time periods while the additional events listed in Table 1 are coloured in black. To visualize the magnitudes of these additional events, we divided them in three categories: 1. magnitude below 2.5, 2. magnitudes from 2.5 to 3, and 3. magnitudes from 3 to 4. These categories are represented by thin dotted, thin dashed and tick dotted black lines, respectively.
- **Determination of the phase noise.** To obtain the noise  $P_{\text{noise}}$  of the unwrapped phase  $P$  we calculate its deviation  $dP(t) = P(t) - P_{\text{base}}(t)$  from the basic phase  $P_{\text{base}}$  at time  $t$ . Here,  $P_{\text{base}}$  is obtained in a procedure described in [33] as the mean value of unwrapped phase in the defined time bins around time  $t$ . Finally, noise of  $P$  is determined as the maximum of  $|dP|$  after elimination of the largest  $p$  percent of its values. To find this value, we first sorted the values of  $|dP|$  into an ascending array  $dP_{\text{as}} = \text{sort}(|dP|)$  of  $N$  members, and determined the value of the phase noise as the value of the term that is  $i_{\text{noise}} = N \cdot (100 - p)/100$  in this array:

$$P_{\text{noise}} = dP_{\text{as}}(i_{\text{noise}}). \quad (1)$$

In this study we use  $p = 5\%$  like in [33].

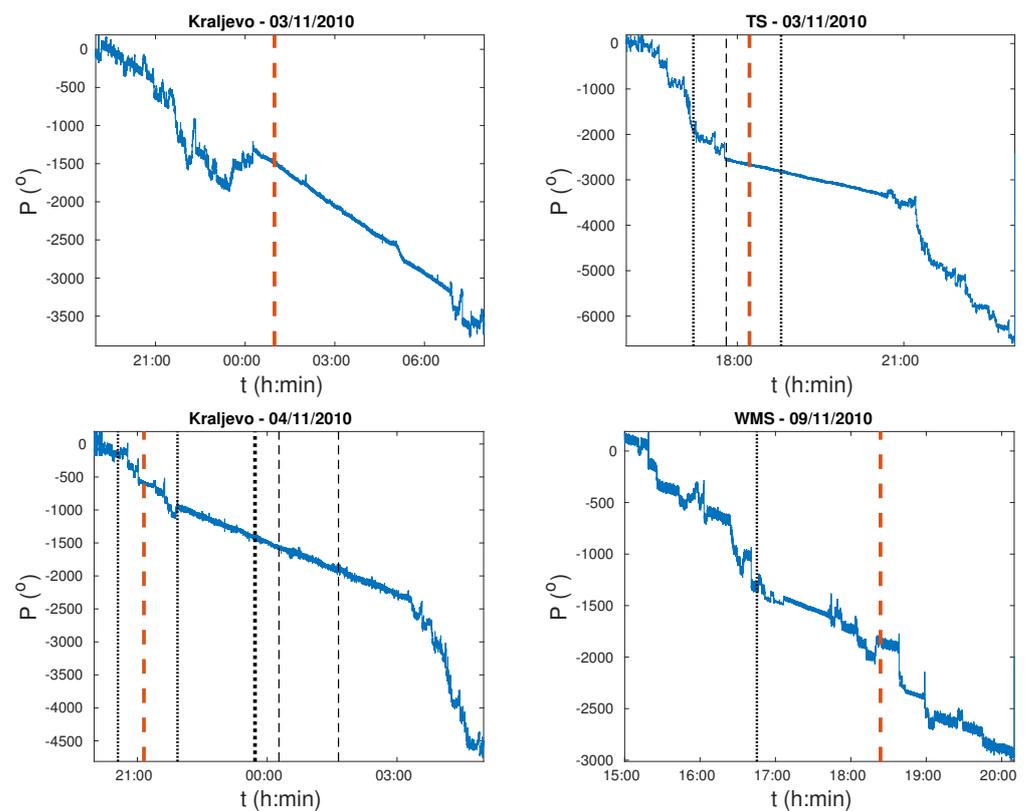
- **Acoustic and gravity waves—excitations and attenuations.** Research of the acoustic and gravity waves in this paper is based on processing of the VLF signal phase. We analyse their excitations and attenuations in periods around the considered EQs using the procedure given in [33]. It is based on the application of the Fast Fourier Transform (FFT) on fixed window time intervals (WTI) within the considered time periods. Keeping in mind that WTI affects the maximum of observable wave period and precision in the analysis of the observed variations we choose three WTIs of 20 min, 1 h and 3 h.

The goal of this procedure is to analyse the recorded phase in frequency domain and connect the wave-periods for which important changes are recorded to the acoustic and gravity waves. The acoustic cut-off  $\tau_0$  and the Brunt-Väisälä  $\tau_{BV}$  wave-periods representing minimal and maximal periods for the acoustic and gravity waves, respectively, are determined from the expressions:

$$\tau_0 = \frac{4\pi v_s}{\gamma g}, \quad \tau_{BV} = \frac{2\pi}{N_{BV}}, \quad (2)$$

where  $\gamma = 5/3$  is the standard ratio of specific heats and  $g = 9.6 \text{ m/s}^2$  is gravitational acceleration. The adiabatic sound speed squared  $v_s^2 = \gamma k_B T_0 / m_a$  is obtained for the gas temperature  $T_0 = 220 \text{ K}$  (estimated from the International Reference Ionosphere (IRI) model [41] and assumed average mass of atoms  $m_a \approx 10^{-25} \text{ kg}$ . The Boltzmann constant  $k_B$  is  $1.3807 \cdot 10^{-23} \text{ J/K}$ . Details of this procedure can be found in, for example, [42,43].

As it is obtained in [33] waves with periods  $T < \tau_0 = 176.7, \text{ s}$  and  $T > \tau_{BV} = 180.4 \text{ s}$ , represents acoustic and gravity modes, respectively.



**Figure 2.** Time evolutions of the unwrapped phase  $P$  for the considered periods. Red lines represent the main EQ considered in corresponding time periods. The additional events listed in Table 1 are coloured in black. Times of the additional considered EQ events with magnitude below 2.5, from 2.5 to 3, and from 3 to 4 are represented by thin dotted, thin dashed and tick dotted black lines, respectively.

Here we point out that during the considered time periods there are not recorded other events which can influence the signal phase. Detailed analysis, described in [33], indicated that influences of receiver, transmitter, meteorological and geomagnetic conditions, which are suggested as the most important non-ionospheric sources of the VLF signal variations [44], can be ignored.

### 3. Results and Discussions

Results of determination of the phase noise and periods of the excited and attenuated acoustic and gravity waves are presented in Sections 3.1 and 3.2, respectively.

#### 3.1. Signal Phase Noise

Time evolutions of deviation of the wrapped phase from its basic values  $dP$  and the phase noise  $P_{\text{noise}}$  obtained by the proposed methodology are shown in Figures 3 and 4, respectively. As one can see, reduction of the phase noise is recorded for all four main EQ events and it is clearly visible in the first two cases whose magnitudes are greater than 5.

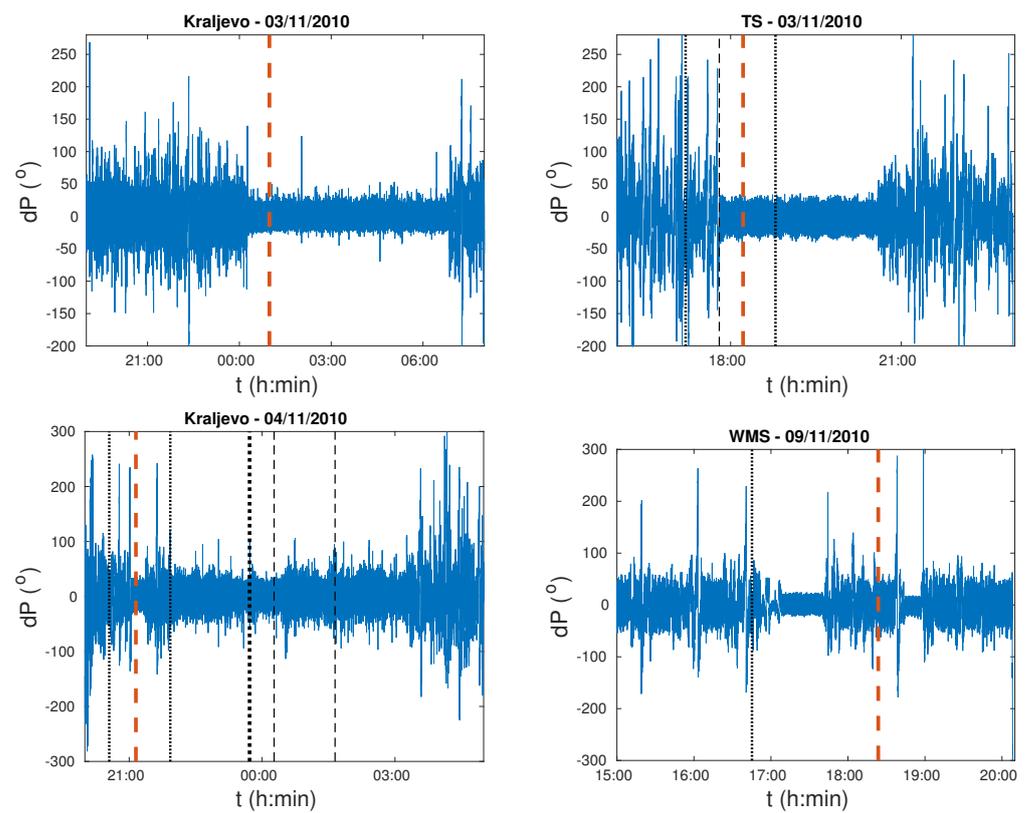


Figure 3. The same as in Figure 2 but for phase deviation  $dP = P_{data} - P_{base}$ .

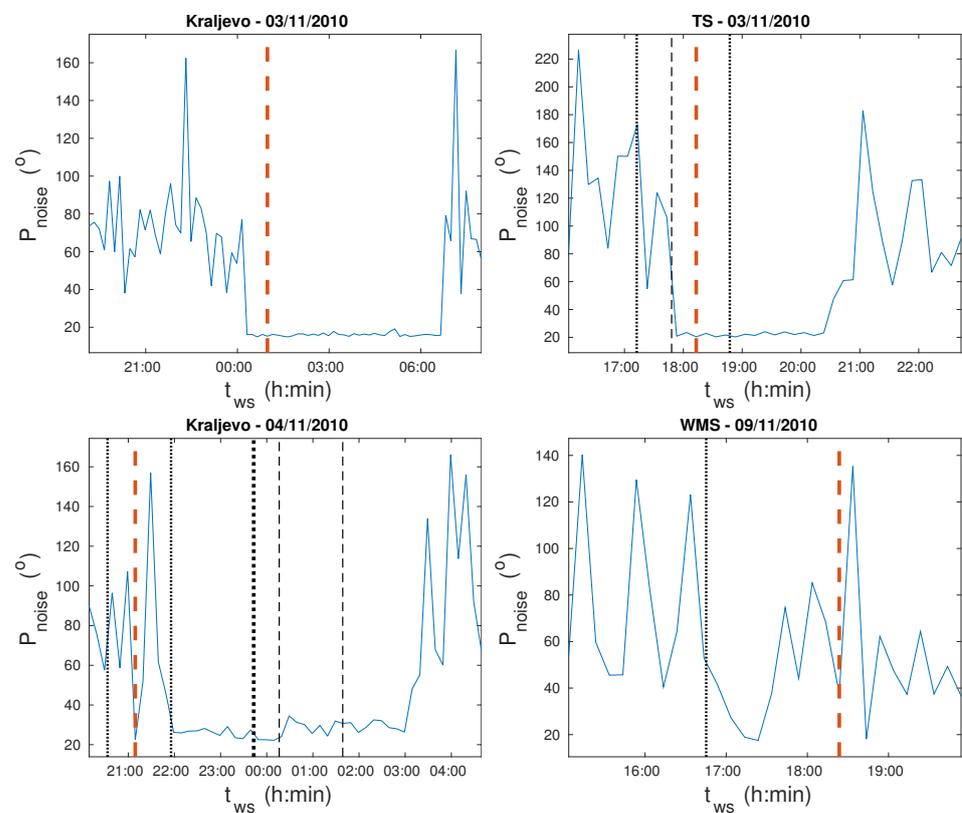


Figure 4. The same as in Figure 2 but for the phase noise  $P_{noise}$ .

It is worth noting that after the first EQ event near Kraljevo which occurred on 3 November 2010, additional 31 EQs occurred in areas near the considered signal propagation path. Twenty nine of them occurred when the phase noise reduction is clearly visible, while two events of weak intensities (magnitudes of 2.1 and 2.2) occurred after increasing of the phase noise. The most intensive additional EQ had magnitude of 3.3. However, despite the large number of accompanying earthquakes, no significant variations were observed either in  $dP$  or in  $P_{\text{noise}}$ . The absence of these variations is also noticeable in the second case (EQ in Tyrrhenian Sea of magnitude 5.1) when three more earthquakes (magnitudes of 2, 2.1 and 2.5) were recorded within about 1.5 h.

In the cases of the other two EQs which magnitudes were between 4 and 5, analysis of  $dP$  and  $P_{\text{noise}}$  time evolutions is not so simple like in the first two considered time intervals. Namely, although the reductions in phase noise are recorded before and after these events, they are not related each other. In the case of the EQ which occurred near Kraljevo on 4 November 2010 significant reduction in  $P_{\text{noise}}$  is recorded several minutes before the EQ and lasts about 20 min. This phase reduction is followed by the short-term increase in  $P_{\text{noise}}$  and additional significant decrease which yield to the second reduction lasting about 5 h. During the second reduction four additional EQs are recorded. The first one occurred in North Italy while the other three near Kraljevo (like the main one) with magnitudes of 3.3, 2.8 and 2.5. As one can see in the bottom left panels of Figures 3 and 4, although the small increase in noise is recorded after the second additional EQ near Kraljevo, significant reduction which can be related with EQ events observed within a time window of about 3 h and includes the time of the last EQ.

Reduction of the phase noise begins about 1 h before the EQ in the Western Mediterranean Sea but it is also possible relate it with the EQ occurred in Central Italy just before the decrease in  $dP$ . This reduction is followed by noise increase which begins more than a half of hour before the EQ and lasts about 1 h before the reduction is recorded again. This event is interesting because position of the EQ epicentre is, contrary to the other events, northern than the signal propagation path. The possible recorded time shift of the reduction time opens a question of influence of position of the EQ epicentre with respect to the signal propagation path. This task requires a specific statistical analysis and it will be in focus of our forthcoming research.

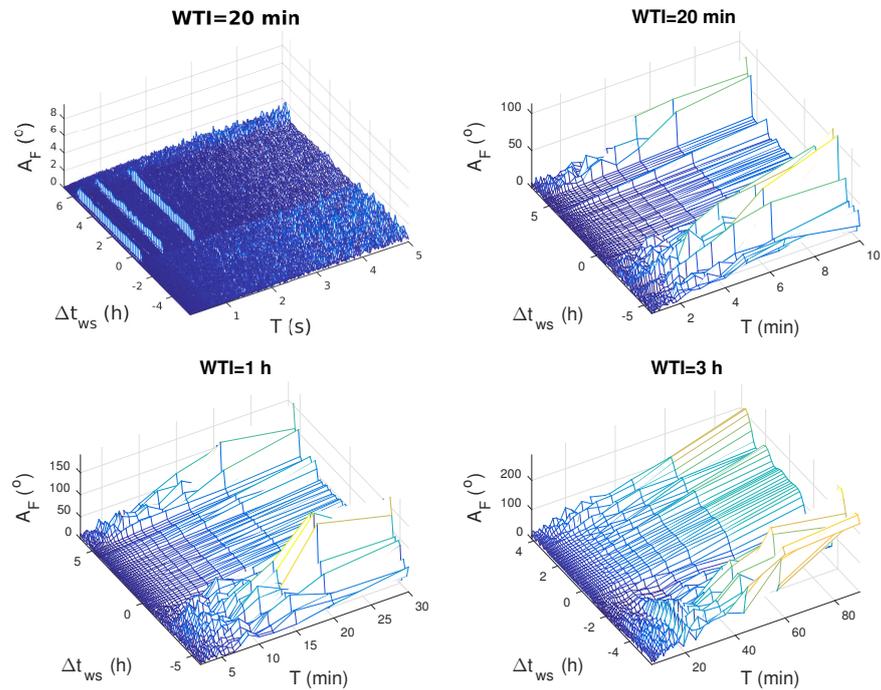
By comparison with noise amplitude reduction analysed in [33] we can conclude that the characteristics of phase reduction are the same for the first two cases: they last for several hours, begin before and end after an EQ event, and there are not observed changes that could be related to other earthquakes of lower intensity. In the third and fourth cases phase noise reductions are also recorded, but they are shortly interrupted by the noise amplifications. Also, it cannot be claimed that the strongest earthquake in the observed period masks the potential relationship between phase noise reduction and weaker EQs.

### 3.2. Acoustic and Gravity Waves

In the second part of this study we analyse signal phase in frequency domain. We apply FFT to the recorded data to research possible excitations and attenuations of the acoustic and gravity waves that can be considered as ionospheric disturbances connected to earthquakes.

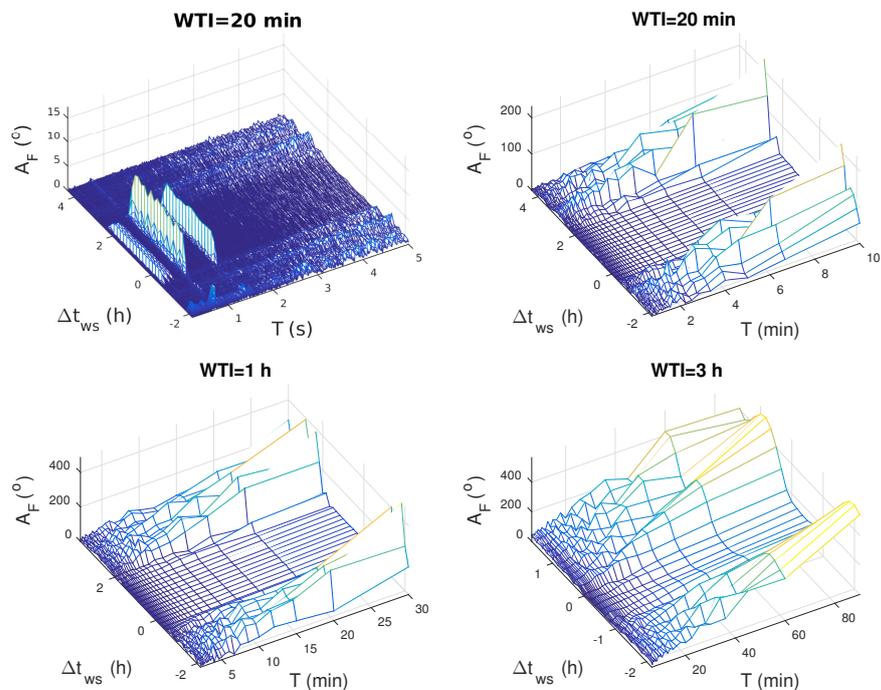
To better visualize periods of the excited/attenuated waves we apply the same procedures like in [33]: (1) we consider three WTI of 20 min, 1 h and 3 h and, (2) in order to better present changes for smaller and greater wave periods  $T$ , the obtained values for all WTIs are considered for smaller and greater wave period domains separately. In this study we showed lower periods  $T$  for the first WTI, and greater periods for all the three WTIs. The results of the analyses for the considered four time periods are shown in Figures 5–8.

**Kraljevo—03/11/2010**



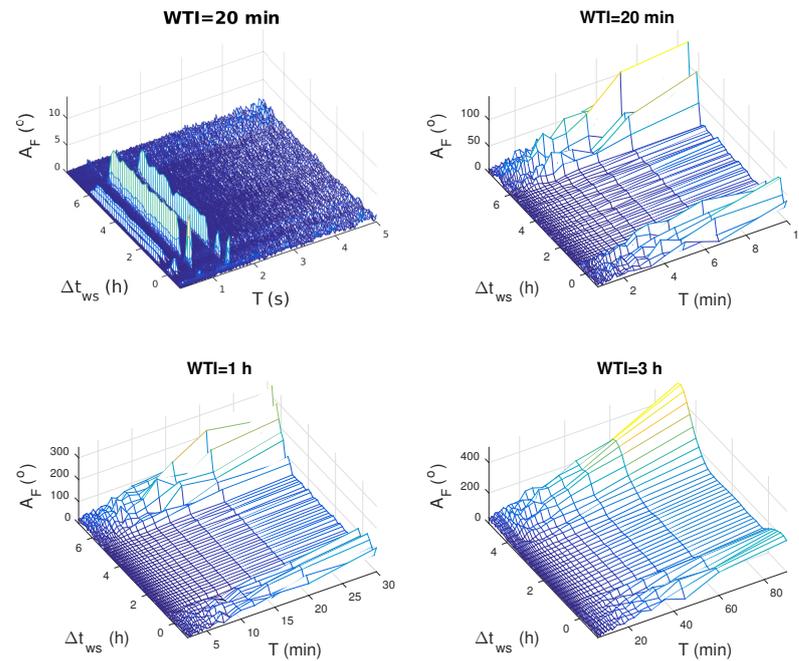
**Figure 5.** Fourier amplitude of waves with period  $T$  obtained by applying FFT to the ICV signal phase recorded in time around EQ occurred near Kraljevo on 3 November 2010 with window time intervals (WTI) of 20 min (upper panels), 1 h (bottom left panel), and 3 h (bottom right panel) which begin with a  $\Delta t_{WS}$  shift with respect to the EQ time.

**Tyrrhenian Sea—03/11/2010**



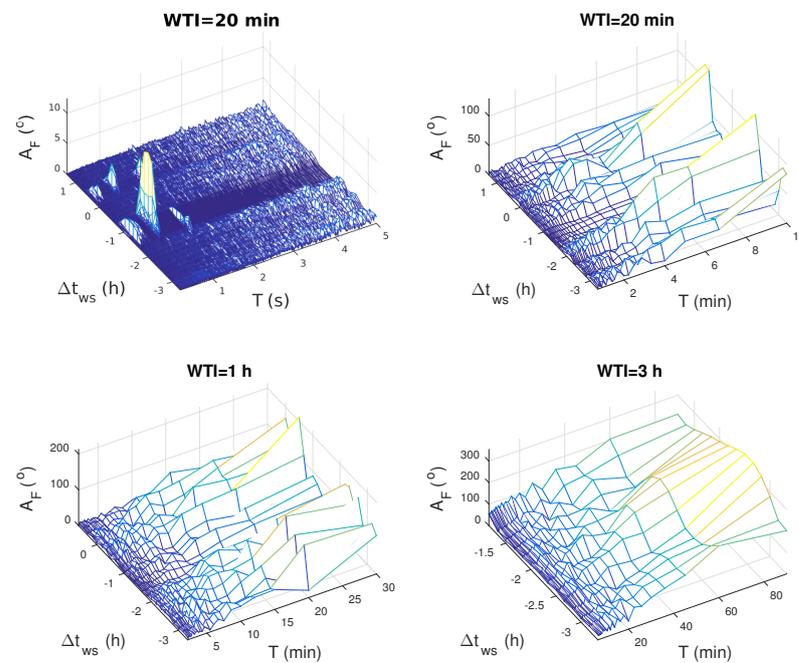
**Figure 6.** Fourier amplitude of waves with period  $T$  obtained by applying FFT to the ICV signal phase recorded in time around EQ occurred in the Tyrrhenian Sea on 3 November 2010 with window time intervals (WTI) of 20 min (upper panels), 1 h (bottom left panel), and 3 h (bottom right panel) which begin with a  $\Delta t_{WS}$  shift with respect to the EQ time.

**Kraljevo—04/11/2010**



**Figure 7.** Fourier amplitude of waves with period  $T$  obtained by applying FFT to the ICV signal phase recorded in time around EQ occurred near Kraljevo on 4 November 2010 with window time intervals (WTI) of 20 min (upper panels), 1 h (bottom left panel), and 3 h (bottom right panel) which begin with a  $\Delta t_{WS}$  shift with respect to the EQ time.

**Western Mediterranean Sea—09/11/2010**



**Figure 8.** Fourier amplitude of waves with period  $T$  obtained by applying FFT to the ICV signal phase recorded in time around EQ occurred in the Western Mediterranean Sea on 9 November 2010 with window time intervals (WTI) of 20 min (upper panels), 1 h (bottom left panel), and 3 h (bottom right panel) which begin with a  $\Delta t_{WS}$  shift with respect to the EQ time.

Similarly to the analysis of the noise amplitude shown in [33], excitations are recorded in periods when phase noise is reduced at several values of  $T$  which are smaller than 1.5 s:

- Kraljevo—03/11/2010: 0.2 s, 0.23 s, 0.47 s (weak increase of the Fourier amplitude), 0.7 s and 1.4 s.
- Tyrrhenian Sea—03/11/2010: 0.23 s, 0.35 s, 0.47 s, 0.7 s and 1.4 s.
- Kraljevo—04/11/2010: 0.23 s, 0.35 s, 0.7 s and 1.4 s;
- Western Mediterranean Sea—09/11/2010: 0.23 s, 0.35 s, 0.47 s (during the first time period when noise reduction is recorded), 0.7 s and 1.4 s.

As one can see, the common periods are 0.23 s, 0.35 s, and 1.4 s while waves at period of 0.47 s are excited in three cases whereby these excitations are weak in the first case, while in the fourth case they are recorded only for one of two periods of phase noise reduction. Excited waves at 0.2 s are recorded only for the first considered EQ event. Because the obtained values are lower than calculated maximum period of the acoustic waves ( $\tau_0 = 176.7$ ) we can conclude that acoustic waves are excited in periods when noise reduction occurs.

Although the previous results related to the noise reductions and excitations of the acoustic waves are very similar as those shown in analysis of the amplitude, there is a difference in wave attenuation. Namely, the Fourier amplitude for several discrete values before the reduction time are reported in [33]. In this study, these peaks are not recorded. In addition, in all four considered periods attenuations are clearly visible at all periods  $T$  (except those values for which excitation is recorded) for phase while these attenuations are much less pronounced in the case of the amplitude for time period around Kraljevo EQ on 3 November 2010. Recent results obtained for the Kumamoto M7.2 earthquake on 15 April 2016 in Japan [45] are in close correlation with our results, so we can suppose that registered oscillations are the acoustic gravity waves of the same origin as is described in this paper. A similar correlation of our results can be also found in their comparison with data presented in analysis of the M7.8 earthquake with the epicenter in Nepal on 25 April 2015 [46].

#### 4. Conclusions

In this paper we analysed the VLF signal phase in time periods around four earthquakes which magnitude were greater than 4. The goal of this study was to extend analyses of reductions of the noise amplitude of VLF signals, excitations and attenuations of the acoustic and gravity waves presented in [33] to the corresponding analyses of the phase of the VLF signal. We analysed data recorded by the receiver located in Belgrade, Serbia for VLF signal emitted by the ICV transmitter in Italy.

The obtained results of this study can be summarised as follows:

- In the cases of EQs with magnitudes greater than 5, a multi-hour noise reductions was observed. As in the case of the amplitude, they begin before the earthquake. In these cases, no changes that could be related to other earthquakes of lower intensity were observed.
- In the cases of EQs with magnitudes between 4 and 5, phase noise reductions are also recorded, but they are shortly interrupted by the noise amplifications. Specific reductions are potentially related to different EQs, i.e., it cannot be claimed that the strongest earthquake in the observed period masks the potential relationship between phase noise reduction and weaker EQs.
- Because the recorded phase reductions are very similar like those in the case of the amplitude the choice of the signal characteristic which can be used in the corresponding studies depends only on the quality of the recorded data and do not affect the results of study.
- Excitations of the acoustic waves are recorded for all four periods. The obtained wave-periods are below 1.5 s which is in agreement with results obtained in analysis of the amplitude.

- Attenuations of the acoustic and gravity waves are recorded continuously with wave-period except for those  $T$  corresponding to wave excitations. This result does not agree with those obtained when analysing amplitude variations where attenuations are primarily recorded for discrete values of wave periods, while similar continuous attenuations are much less pronounced.

**Author Contributions:** Conceptualization, methodology, investigation, resources, formal analysis, writing—original draft, preparation, visualization, A.N.; software, data curation, A.N. and S.T.M.; validation, P.F.B., S.P., L.Č.P. and M.R.; writing—review and editing, all authors. All authors have read and agreed to the published version of the manuscript.

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**Informed Consent Statement:** Not applicable.

**Data Availability Statement:** Publicly available datasets were analysed in this study. This data can be found here: <http://www.emsc-csem.org/Earthquake/> (accessed on 28 February 2021); [https://ccmc.gsfc.nasa.gov/modelweb/models/iri2012\\_vitmo.php](https://ccmc.gsfc.nasa.gov/modelweb/models/iri2012_vitmo.php) (accessed on 29 February 2021). The VLF data used for analysis is available from the corresponding author.

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**Conflicts of Interest:** The authors declare no conflict of interest.

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