



Article Comprehensive Validation and Comparison of Three VIIRS Aerosol Products over the Ocean on a Global Scale

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Abstract: Three parallel Visible/Infrared Imager Radiometer Suite (VIIRS) aerosol products (SOAR, NOAA, and AERDT) provided data since 2012. It is necessary to study the performances and advantages of different products. This study aims to analyze the accuracy and error of these products over the ocean and compare them with each other. The results show that the three VIIRS ocean aerosol retrievals (including total aerosol optical depth (AOD), fine mode fraction, Ångström exponent (AE), and fine AOD (AODF)) correlate well with AErosol RObotic NETwork (AERONET) retrievals (e.g., correlation >0.895 for AOD and >0.825 for AE), which are comparable to the newest moderate-resolution imaging spectro-radiometer (MODIS) retrievals. Overall, the SOAR retrievals with quality filtering have the best validation accuracy of all parameters. Therefore, it is more recommended to use. The differences in the annual AOD spatial patterns of different products are small (bias < 0.016), but their AE spatial patterns are evidently different (bias > 0.315), indicating the large uncertainty of VIIRS AE. Error analysis shows that the scattering angle and wind speed affect aerosol retrieval. Application of the non-spherical dust model may reduce the dependence of retrieval bias on the scattering angle. Overall, this study provides validation support for VIIRS products usage and possible algorithm improvements.

Keywords: VIIRS; fine mode fraction; fine mode aerosol optical depth; AERDT; AERDB

1. Introduction

Aerosols, as an important component of the atmosphere, influence the cloud property, Earth's radiation balance, and global climate change [1–4]. Excessive aerosols affect human health and life as pollutants (such as $PM_{2.5}$) [5]. With the development of remote sensing technology, satellite aerosol observation from space has been widely used in related scientific research [6].

The ocean surface reflectance is stable and low, which is suitable for aerosol measurements using satellites [7,8]. This is because the aerosol signal accounts for a higher proportion of satellite-observed reflectance at a low surface reflectance. For example, Geogdzhayev et al. [9] developed a two-channel aerosol retrieval algorithm over the ocean and applied it to global Advanced Very High-Resolution Radiometer (AVHRR) observations for retrieving aerosol optical depth (AOD) and Ångström exponent (AE). Preliminary validation results of the AOD retrieved by the algorithm show that it correlates well with the AOD measured by the sun photometer [10]. Owing to the good and stable radiation performance of sea-viewing wide field-of-view sensor (SeaWiFS) and moderate-resolution imaging spectro-radiometer (MODIS) sensors, it is reliable to use them to retrieve aerosol



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Copyright: © 2022 by the authors. Licensee MDPI, Basel, Switzerland. This article is an open access article distributed under the terms and conditions of the Creative Commons Attribution (CC BY) license (https:// creativecommons.org/licenses/by/ 4.0/). properties over the ocean. For example, Sayer et al. [11] developed the Satellite Ocean Aerosol Retrieval (SOAR) algorithm for SeaWiFS. Its validation results show that the correlation of AOD is >0.85, and the correlation of AE is 0.78 when AOD is greater than 0.3. The MODIS operational aerosol product over the ocean also has good validation accuracy of AOD and AE [12]. Although these preliminary validations and comparisons are promising, they are limited compared to a large number of aerosol validations over land. And most studies validated AOD and AE for ocean aerosol products, however, fine mode fraction (FMF) and fine mode AOD (AODF) parameters had only limited comparison (e.g., [13]). These parameters are important and can be used in climate change research [14,15].

As the successor to AVHRR, SeaWiFS, and MODIS, the Visible/Infrared Imager Radiometer Suite (VIIRS) is also used for aerosol retrieval over the ocean. There are currently three parallel VIIRS ocean aerosol products, including the EDR product developed by National Oceanic and Atmospheric Administration (NOAA), the AERDT product developed by the National Aeronautics and Space Administration (NASA) Dark Target (DT) team, and the SOAR product developed by the NASA Deep Blue (DB) team [16–18]. Huang et al. [19] validated the AOD and AE of NOAA product using the AErosol RObotic NETwork (AERONET) measurements. Saver et al. [20] evaluated the AOD, AE, and FMF parameters of the SOAR product. Sawyer et al. [17] evaluated the performance of AERDT AOD and compared it with MODIS operational AOD. Note that the above three VIIRS ocean aerosol products directly retrieve AOD and FMF and can derive AE and AODF. Similar to product validations for other sensors (mentioned above), studies evaluating these four parameters from VIIRS simultaneously are lacking. Comparing the performance of different aerosol products (retrieved from the same or different sensors) and providing usage recommendations is important to data users. For example, Sayer et al. [21] compared MODIS DT, DB, and the merged product of DT and DB over land and provided their usage recommendations. Liu et al. [22] comprehensively evaluated and compared the performance of MODIS DT, DB, and Multiangle Implementation of the Atmospheric Correction AOD products over mainland China. He et al. [23] validated and compared the performance of VIIRS EDR and DB AOD products over land in the Chinese mainland area. However, similar studies over the ocean for all available VIIRS products (Table 1) are limited. Therefore, it is necessary to compare the three different VIIRS aerosol products over the ocean to understand their performances and advantages. This study aims to comprehensively evaluate and compare the accuracy, error, and spatial pattern of AOD, AE, FMF, and AODF retrieved by the three VIIRS ocean aerosol algorithms using consistent validation standards and data.

Section 2 summarizes the data sets used, including the VIIRS, MODIS, and AERONET aerosol products. Section 3 presents the validation and comparison results. Section 4 provides some discussions. Section 5 summarizes this study.

2. Materials and Methods

The VIIRS sensor onboard the S-NPP satellite has 22 channels. Its spatial resolution is approximately 375–750 m and it passes through the equator at approximately 1:30 p.m. local time. Since 2012, it has provided extensive observations of land, atmosphere, and ocean from space. Based on the useful VIIRS data, three different aerosol products were developed. This study uses data from the full year of 2018. For convenience, NOAA denotes NOAA-EDR product or algorithm over the ocean, AERDT denotes AERDT product or algorithm over the ocean, AERDT denotes AERDT product or algorithm over the ocean. Hereafter, the wavelength of AOD, FMF, and AODF is 550 nm, and the wavelength of AE is 550–870 nm.

2.1. AERDT Aerosol Algorithm and Product

To produce long-term aerosol records from satellites, the NASA Dark Target (DT) team developed the VIIRS DT algorithm to extend the current satellite DT product (e.g., from MODIS DT to AERDT) [17,24]. In general, the VIIRS ocean DT algorithm inherits

from that of MODIS. However, because of the difference between the spectral response functions of VIIRS and MODIS sensors, parameters such as atmospheric water vapor, ozone correction coefficients, and Rayleigh optical depth of the VIIRS spectral bands need to be re-derived [25]. The AERDT algorithm takes as input the 750 m resolution TOA reflectance of 6-minute granule data processed by NASA. Then unsuitable pixels (such as clouds, ice, ocean glints, and land) are removed. A 6×6 spatial window is used to aggregate pixels to a coarser resolution for a higher signal-to-noise ratio and better retrieval efficiency. In the 6×6 box, the algorithm removes the darkest and brightest 25% of pixels in the 0.87 µm band. Additionally, it takes the average of the remaining pixels as the reflectance of the box, which is an approximately 6-km resolution. The box reflectance will be trace gas corrected to obtain the final TOA reflectance used for the aerosol retrieval. The appropriate TOA spectral reflectance (0.55, 0.67, 0.87, 1.24, 1.61, and 2.25 µm) is compared with the spectral reflectance calculated by the look-up table, and the average of the acceptable retrieval solutions is taken to obtain the AOD and FMF, and then the AE is calculated. The spatial resolution of all parameters of the AERDT product is 6 km (Table 1).

 Table 1. Summary of the three VIIRS aerosol datasets used in this study.

Products	Field	Parameter	Download
AERDT	Effective_Optical_Depth_Average_Ocean Optical_Depth_Ratio_Small_Ocean_0p55micron	AOD FMF	https://earthdata.nasa.gov, last access: 23 May 2022.
	Angstrom_Exponent_1_Ocean Land_Ocean_Quality_Flag	AE Quality flag	
NOAA	AerosolOpticalDepth_at_550 nm SmallModeFraction AngstromExponent QF1_VIIRSAEROEDR	AOD FMF AE Quality flag	https: //www.avl.class.noaa.gov, last access: 23 May 2022.
SOAR	Aerosol_Optical_Thickness_550_Ocean_Best_Estimate Fine_Mode_Fraction_550_Ocean_Best_Estimate Angstrom_Exponent_Ocean_Best_Estimate Solar_Zenith_Angle Viewing_Zenith_Angle Scattering_Angle Wind_Speed	AOD FMF AE Solar zenith angle Satellite zenith angle Scattering angle Wind speed	https://earthdata.nasa.gov, last access: 23 May 2022.

2.2. NOAA Aerosol Algorithm and Product

NOAA's ocean aerosol algorithm, described in Jackson et al. [16], is very similar to that of the AERDT. They share similar retrieval strategies and aerosol model parameters, but there are some differences. The NOAA algorithm does not retrieve aerosols over large inland lakes. It replaces the 0.55 μ m band used by the AERDT algorithm with the 0.746 μ m band. Additionally, it reports the retrieval of the optimal solution instead of the average of the acceptable solutions. The upper limit of AOD retrieval of the NOAA algorithm is 2.0, while that of the AERDT algorithm is 5.0. Their pixel aggregation methods are different. NOAA algorithm first retrieves 750 m resolution aerosol properties and then aggregates them to a 6-km coarse resolution (Table 1).

2.3. SOAR Aerosol Algorithm and Product

The SOAR algorithm used by the AERDB product is described in Sayer et al. [18], which differs from the NOAA and AERDT algorithms. It has two retrieval schemes, a full retrieval, and a backup retrieval, depending on whether the ocean surface is turbid. In the open ocean, SOAR uses the full retrieval based on seven bands (0.49, 0.55, 0.67, 0.87, 1.24, 1.61, and 2.25 μ m), while on turbid water surfaces it uses the backup retrieval based on the four near-infrared bands (0.87, 1.24, 1.61, and 2.25 μ m). SOAR's aerosol model is based on internal limited mixing of the coarse and fine modes, i.e., limiting FMF to a reasonable range (Table 2 in [18]), rather than the free mixing approach used by AERDT and NOAA

algorithms. SOAR uses a non-spherical dust optical model instead of the spherical one used by AERDT and NOAA [26]. The improvement of the dust model has a large impact on the retrieval results, which will be discussed later. In terms of retrieval strategy, SOAR uses the optimal estimation theory, with AOD and FMF as free parameters, and its initial value is determined by the AOD and FMF nodes in the look-up table with the smallest residual error. When the spectral AOD and FMF are determined, then AE is calculated using 550 and 870 nm AOD. SOAR first retrieves aerosol properties at a 750 m resolution and then aggregates them to a 6-km resolution (Table 1).

2.4. MODIS Aerosol Data

MODIS aerosol products (MOD04 and MYD04) over the ocean have been widely validated and used (e.g., [12,27,28]). Therefore, this study used MYD04 data as a comparison to characterize the relative accuracy of VIIRS aerosol products. MYD04 is chosen instead of MOD04 because MYD04 has better validation accuracy and is close to the VIIRS imaging time [29]. MYD04 also records AOD, AE, and FMF properties, but at a 10-km resolution.

2.5. AERONET Data

The uncertainty of direct-solar AOD observations from AERONET is small (~0.01 in visible), so it serves as a ground benchmark to validate satellite retrievals [30–32]. The newly released version 3 data applies automated cloud screening and instrument anomaly quality control technique, which greatly speeds up the production of level 2 data [33]. In this study, level 2 AERONET AOD and AE data of version 3 are used. AERONET Spectral Deconvolution Algorithm (SDA) and standard inversion products provide FMF parameters [34,35]. Because of the better temporal sampling of SDA FMF, it is used to compare satellite FMF and AODF retrieval [36]. Note that the uncertainty of the SDA FMF is approximately 0.1 [34].

2.6. Validation Method

All AERONET sun photometers provide AOD at 440, 675, 870, and 1020 nm, some instruments provide other additional wavelengths (e.g., 340, 380, and 500 nm). However, none of them provide 550 nm AOD, instead, 550 nm AOD is the most commonly used by satellite retrieval. Therefore, the quadratic log interpolation method was used to calculate the AERONET 550 nm AOD from the other wavelengths of AOD using Equation (1) [20,30,36]. Then, the 550 and 870 nm AODs are used to calculate the AE, which is for direct comparison with the satellite-retrieved AE. AERONET SDA product only provides 500 nm FMF. For direct comparison with FMF and AODF retrievals from VIIRS, the SDA total and fine mode AOD is interpolated to 550 nm using corresponding AE, and then the 550 nm FMF is calculated using Equation (2) [20]. Since the spatiotemporal sampling of AERONET retrieval and satellite retrieval is inconsistent, they require spatiotemporal matching. In this study, a temporal window of ± 30 min is used to calculate the median AERONET retrieval, and a spatial window of ± 25 km is used to calculate the median satellite retrieval. Note that only the satellite retrievals with quality filtering are used for comparison. The minimum number of AERONET values allowed within the window is one, and the minimum fraction of valid satellite values is 20% [17,20].

$$\log(\tau_{\lambda}) = a_0 + a_1 \log(\lambda) + a_2 \log(\lambda)^2 \tag{1}$$

where τ_{λ} is the AOD at wavelength λ ; a_0 , a_1 , and a_2 are the fitting coefficients, which are commonly calculated by AOD at 440, 500, 675, and 870 nm using each AERONET record.

$$f_{550} = \tau_{550}^{\rm t} / \tau_{550}^{\rm t} \tag{2}$$

where τ_{550}^t , τ_{550}^t , and f_{550} are fine mode AOD, total AODT, and FMF at 550 nm, respectively.

The metrics used to evaluate VIIRS ocean aerosol products include: (1) the number (N) is the volume of matching pairs. (2) the correlation coefficient (R) is used to characterize

the variation consistency between satellite retrieval and sun photometer retrieval. (3) Mean bias (MB) is used to characterize the overall overestimation or underestimation of satellite retrieval. (4) Mean absolute error (MAE) and root mean square error (RMSE) are used to characterize the overall error of satellite retrieval. Note that MAE and RMSE are related to the level of comparison values and RMSE is easily affected by outliers. (5) The expected error is used to characterize the fraction of satellite retrieval that meets the expected accuracy, and it is generally believed that the fraction should be greater than 68% (i.e., one standard deviation). Equations (3)–(5) represent the expected errors for the AOD (EE), AE (EE_AE), and FMF (EE_FMF), respectively. The Global Climate Observing System (GCOS) provides the expected accuracy of AOD for long-term climate change research as Equation (6).

$$EE = \pm (0.03 + 10\% \times \tau_{AERONET}) \tag{3}$$

$$EE_AE = \alpha_{AERONET} \pm 0.4 \tag{4}$$

$$EE_FMF = f_{AERONET} \pm 0.2$$
(5)

$$GCOS = maximum(0.03, 10\% \times \tau_{AERONET})$$
(6)

where, $\tau_{AERONET}$, $\alpha_{AERONET}$, and $f_{AERONET}$ is AERONET AOD, AE, and FMF, respectively.

3. Results

3.1. Overall Comparison Results with AERONET

3.1.1. Aerosol Optical Depth

The validation results of the total AOD are shown in Figure 1. SOAR AOD metrics (EE:0.681 and GCOS:0.545) are similar to those of Sayer et al. [20] (EE:0.66 and GCOS:0.49–0.52). The NOAA AOD RMSE (0.075) is slightly larger than the result (0.065) of Huang et al. [19]. The EE of MYD04 AOD in the study of Levy et al. [12] is higher than that of our result (Figure 1d), which is owing to their use of a more relaxed EE formula. The ocean AOD of all products has a good correlation with the AERONET AOD with the R of 0.895–0.917. SOAR AOD has the best performance with smaller errors (MAE and RMSE) and higher EE and GCOS fractions. The fraction of SOAR AOD falling on EE is 0.681, which meets the goal of EE (~0.68). The fraction within EE of NOAA AOD is 0.634. Both of them have better accuracy than MYD04 AOD (EE of 0.607). In contrast, AERDT using the same algorithm as MYD04 has lower accuracy than MYD04 AOD. Overall, except for AERDT AOD, SOAR and NOAA AOD are equal to or better than the MYD04 AOD level.

3.1.2. Ångström Exponent

AE utilizes the different responses of two wavelengths to aerosol to characterize the size of aerosol particles. This parameter is useful for distinguishing between fine and coarse dominant aerosol types. Therefore, the AE parameter is also the focus of this study. When the AOD loading is low, there is a non-negligible uncertainty in the calculated AE; therefore, the AE parameters are compared only when the AERONET 550 nm AOD is >0.2.

The validation results of AE are shown in Figure 2. Since only the AE with AOD > 0.2 is compared, validations for SOAR, NOAA, and MYD04 AE are better than previous studies [12,18,19]. Although each product captures the variation in AE, i.e., R > 0.825, they all overestimate low AE and underestimate high AE. The same problem was also found in the previous validation of MODIS AE (e.g., [37]). Zhou et al. [38] proved that using a non-spherical dust model can reduce the overestimation of low AE and FMF. Fortunately, most of their retrievals are within the EE_AE (>0.760), and their RMSEs are all smaller than EE_AE (defined as ± 0.4), indicating that their AE retrievals are quantitatively meaningful. From the two metrics of EE_AE and RMSE, the AEs of SOAR and AERDT are better than that of MYD04, while the AE of NOAA is slightly worse than that of MYD04.



Figure 1. Total AOD at 550 nm validation results using AERONET measurements. (a) SOAR, (b) NOAA, (c) AERDT, and (d) MYD04 products.



Figure 2. Fine mode fraction at 550 nm validation results using AERONET measurements. (**a**) SOAR, (**b**) NOAA, (**c**) AERDT, and (**d**) MYD04 products.

3.1.3. Fine Mode Fraction and AOD

The definitions of VIIRS-retrieved FMF and AERONET-retrieved FMF differ slightly, but they are related and can be compared with each other [13] (Figure 3). The comparative metrics (Figure 3a) of SOAR FMF are similar to previous studies (e.g., [18,20]). Except for NOAA FMF, the performance of SOAR and AERDT FMF is close to or better than that of MYD04 FMF. The SOAR algorithm is better at retrieving low FMF cases, while other algorithms overestimate low FMF. This suggests that NOAA, AERDT, and MYD04 using similar algorithms (spherical dust aerosol model) tend to increase the fraction of fine modes to fit satellite-observed reflectance. Using non-spherical dust models may improve their FMF retrieval [38]. Overall, the FMFs of all algorithms are quantitatively meaningful, i.e., they capture variations in FMF (R > 0.789), and the EE_FMF is >0.713.



Figure 3. VIIRS AE at 550 nm comparison results with that of AERONET. (**a**) SOAR, (**b**) NOAA, (**c**) AERDT, and (**d**) MYD04 products.

A related parameter to FMF is the fine mode AOD (Equation (2)), which can be used to estimate anthropogenic radiative forcing [14,15]. The accuracy of SOAR AODF is very high (e.g., R:0.872, EE:0.803, GCOS:0.727). This indicates that it can be applied to scientific research. Conversely, the scatter plot (Figure 4b–d) and EE metrics show that the NOAA, AERDT, and MYD04 algorithms significantly overestimate AODF. The overestimated total AOD (Figure 1) and FMF (Figure 3) together lead to the overestimation of AODF. Therefore, these algorithms need to be improved in the future.



Figure 4. Fine mode AOD (AODF) at 550 nm validation results using AERONET measurements. (a) SOAR, (b) NOAA, (c) AERDT, and (d) MYD04 products.

3.2. Quantitative Comparison

To reduce the impact of sample differences, the same matchups (same site and same time) of all VIIRS retrievals with AERONET retrievals are used to discuss the relative accuracy of different VIIRS products (Figure 5). The AOD accuracy of SOAR and NOAA is similar and they are close to the 1:1 line. While the AERDT systematically overestimates the AOD, it remains within the expected error envelope. AE and FMF comparison results (Figure 5b,c) show similar product differences, i.e., NOAA and AERDT overestimated low parameter values, which is consistent with the overall validation results (Figures 2 and 3). In contrast, the AE and FMF of SOAR are close to the 1:1 line. For the AODF parameter, SOAR underestimates high AODF. Interestingly, the AODFs of AERDT and NOAA almost overlap, they both overestimate the low AODF and underestimate the high AODF. Overall, the comparison results show that SOAR has the best performance on all parameters.

3.3. Pixel-by-Pixel Comparison

3.3.1. Aerosol Optical Depth

To understand the differences in spatial patterns of different VIIRS total AOD, a pixel-by-pixel comparison is analyzed (Figures 6–8). The three VIIRS products were resampled to a daily mean at 1° spatial resolution for pixel-by-pixel matching. As can be seen in Figure 6a, the aerosols around the coast are heavily influenced by the land aerosols, and the AOD is higher than 0.15, while in the open ocean, the AOD is less than 0.15. In general, the AOD spatial patterns of SOAR and NOAA have little difference (MB = 0.001), while the AOD spatial patterns of AERDT and the other two products have larger differences. Their biases (-0.014 for SOAR and AERDT, 0.016 for AERDT and NOAA) are approximately half of the SOAR annual mean (0.028). Figure 7 shows that correlations between different AOD retrievals are high (R > 0.85) in most marine areas, especially in high AOD loading areas (R > 0.9). In regions under low AOD loading, their

correlations are slightly reduced, which is understandable because the AOD variations in these regions are small. Figure 8 shows that when the latitude is $<\pm40$ degrees, the retrieval frequencies of the three products are similar, and their retrieval frequency ratio is approximately 0.9–1.1. In coastal zones, SOAR had better retrieval frequencies, possibly due to its backup algorithm retrieving a portion of turbid water [18]. With increasing latitude, AERDT has the highest retrieval frequency, followed by SOAR, and NOAA has the lowest retrieval frequency. When all quality AODs instead of high-quality AODs are used to calculate the retrieval frequency (not shown), SOAR outperformed that of AERDT except above 65 degrees north latitude and except above 60 degrees south latitude, NOAA's retrieval frequency exceeds that of the SOAR, indicating that SOAR and NOAA adopts a stricter quality control approach.



Figure 5. Comparison results of VIIRS (**a**) total AOD, (**b**) AE, (**c**) FMF, and (**d**) fine mode AOD parameters with those of AERONET. The dashed black line is 1:1 and the shading is the expected error envelope.

3.3.2. Ångström Exponent

The AE spatial pattern of SOAR is shown in Figure 9a. In the oceans close to the mainland, high AE indicates fine-mode-dominated aerosols in these regions. In most oceans, the MB of SOAR AE and AERDT, and NOAA AE are negative, which is consistent with previous studies [20]. The AE pattern MB (0.062) of AERDT and NOAA is small, which is related to their use of similar algorithms (Section 2.3). This does not mean that they produce more accurate AE retrievals. Because the overall validation result (Figure 5b) shows that compared with AERDT and NOAA algorithms, SOAR AE has a better validation performance, which implies that the spatial pattern of SOAR AE may be more accurate. The correlation coefficients (Figure 10) of SOAR AE and other algorithm AEs are related to the annual mean of AE, that is, in regions with high AE, their correlation coefficients are high, and in regions with low AE, their correlation coefficients are low. This is because the

open ocean is rarely polluted by aerosols from the land, and therefore, AE variations are slight. For the AE parameter, the serious differences between different algorithms indicate that the uncertainty of AE retrieved by VIIRS is large, and more validations are needed to improve the algorithm in the future.



Figure 6. Spatial pattern comparison of three VIIRS products. (**a**) SOAR mean AOD (0.028) matched AERDT, (**b**) mean bias (-0.014) of SOAR and AERDT, (**c**) mean bias (0.001) of SOAR and NOAA, (**d**) mean bias (0.016) of AERDT and NOAA.



Figure 7. AOD correlation coefficient of (a) SOAR and AERDT, (b) SOAR and NOAA.



Figure 8. AOD retrieval frequency ratio of (a) SOAR and AERDT, (b) SOAR and NOAA.



Figure 9. Spatial pattern comparison of three VIIRS products. (**a**) SOAR mean AE (0.352) matched AERDT, (**b**) mean AE bias (-0.315) of SOAR and AERDT, (**c**) mean AE bias (0.241) of SOAR and NOAA, (**d**) mean AE bias (0.062) of AERDT and NOAA.



Figure 10. AE correlation coefficient of (a) SOAR and AERDT, (b) SOAR and NOAA.

3.4. Error Analysis

3.4.1. Aerosol Optical Depth

Figure 11 shows the dependence of AOD bias (AOD_{VIIRS}-AOD_{AERONET}) on imaging angles and sea surface wind speed (Table 1). Almost all AOD biases are positive, indicating a systematic overestimation of AOD. As the solar zenith angle increases, the AOD bias decreases. One potential reason is that low latitudes (usually small solar zenith angles) have high AOD loadings and complex sources (Figures 6a and 9a), while high latitudes (usually large solar zenith angles) have low AOD loadings. This leads to a larger AOD bias for small solar zenith angles. Another potential reason is that the solar zenith angle (3D radiative effect) affects the selection of the non-spherical model and AOD retrieval in the algorithm [39,40]. When the scattering angle is $<135^\circ$, all VIIRS products show similar AOD bias; however, when the scattering angle is >135°, their AOD bias is very different, suggesting that the scattering angle (>135°) may have a large influence on the ocean AOD retrieval. According to previous studies (e.g., [26,38]), this AOD bias can be attributed to the scattering phase function of the dust aerosol model. The dust model of spherical approximation used by AERDT and NOAA produces scattering phase functions that deviate from the actual shape, while the SOAR using the non-spherical dust model maintains a stable AOD bias when the scattering angle is >135 (Figure 11c). The study by Lee et al. [26] illustrates that improving the dust aerosol model can reduce the dependence of the AOD bias on the scattering angle. However, it is difficult to determine the right

non-spherical model at the pixel scale, therefore, most studies are based on several fixed non-spherical models [41]. Figure 11d shows that the sea surface wind speed affects the AOD bias, that is, the greater the wind speed, the greater the bias. Although all VIIRS algorithms account for variations in wind speed, suggesting that none of the algorithms can eliminate the effect of wind speed on retrieval.



Figure 11. AOD error (AOD_{VIIRS}-AOD_{AERONET}) dependence analysis of on (**a**) solar zenith, (**b**) satellite zenith, (**c**) scattering angle, and (**d**) wind speed.

3.4.2. Ångström Exponent

Figure 12 shows the dependence of AE bias. Unlike AOD, the solar zenith angle has little effect on AE bias. However, the AE bias of NOAA and AERDT algorithms shows a dependence on the scattering angle, which is also caused by a scattering phase function that does not match the actual shape of the dust aerosol [38]. The SOAR algorithm uses a non-spherical dust model, so its AE bias does not depend on the scattering angle [26]. Therefore, it is recommended that current algorithms update the non-spherical dust model and aerosol model selection strategy to improve aerosol retrieval. Figure 11d shows that the AE bias of all VIIRS products is positively correlated with sea surface wind speed.



Figure 12. AE error (AE_{VIIRS} - $AE_{AERONET}$) dependence analysis of on (**a**) solar zenith, (**b**) satellite zenith, (**c**) scattering angle, and (**d**) wind speed.

4. Discussion

VIIRS has provided data since 2012 and has three different aerosol algorithms applied to it simultaneously. An issue is whether these products can achieve the accuracy of existing products (such as MODIS) and can be applied scientifically. Another issue is the different performances and advantages of the three VIIRS products. This study uses AERONET ground-based observations to simultaneously validate and compare the total AOD, AE, FMF, and AODF of the three VIIRS aerosol products over the ocean using a consistent approach and data.

The validation results on the global scale show that the VIIRS ocean aerosol retrieval performance is good, reaching and possibly better than that of the MODIS aerosol retrieval in the same period. Liu et al. [42] compared the NOAA ocean AOD and the MODIS Collection 5.1 ocean AOD and also found that their accuracy was comparable. Note that users need to be wary of AERDT's AOD and NOAA's FMF and AE, which may not reach the accuracy of MODIS. Sawyer et al. [17] compared AERDT and MODIS AOD over the ocean and also found that the validation accuracy of the AERDT ocean AOD was slightly worse than that of the MODIS. Overall, SOAR shows the best validation accuracy of the aerosol retrievals (including total AOD, AE, FMF, and AODF) in the three products. Especially, the fraction of SOAR AODF better than the GCOS goal (EE goal) is 0.727 (0.803), which suggests that it meets the requirements of scientific research. Notably, we found that all quality retrievals for SOAR do not filter land pixels fully (not shown), so quality filtering for SOAR is recommended. Conversely, NOAA, AERDT, and MODIS significantly overestimate AODF with above EE > 0.418 (Figure 4). This is unavoidable because they overestimate both total AOD and FMF (Figures 1, 3 and 5). Therefore, the AODF of these algorithms should be used with caution for scientific research. There is a small difference in the spatial pattern of AOD between SOAR and NOAA, which is consistent with a previous study [18]. However, the AOD spatial patterns of AERDT versus NOAA and SOAR show

larger differences. Since NOAA and SOAR AODs show better validation accuracies, their AOD spatial patterns are more credible. The AE of SOAR is systematically smaller than that of NOAA and AERDT. Although the difference in AE between AERDT and NOAA is small (MB = 0.062), it does not mean that they retrieve more accurate AEs because SOAR shows a better validation of AE (Figures 2 and 5). Therefore, VIIRS AEs require further validation and analysis. The FMFs and AEs retrieved by VIIRS are correlated [13]; therefore, they generally exhibit the same validation characteristics. For example, AERDT and NOAA overestimated low AE and low FMF (Figure 5). Conversely, SOAR performs well under low AE and FMF conditions. Using the non-spherical dust model may improve the retrievals of FMF and AE [38]. Proper selection of dust pixels and limited mixing ratios of coarse and fine modes are also possible directions for improvement [18,43].

Error analysis results show that scattering angle and sea surface wind speed have an impact on AOD and AE biases, suggesting potential points for algorithm improvement. For example, the non-spherical dust model is used in the SOAR algorithm; therefore, its retrieval has a smaller dependence on the scattering angle [26]. Previous studies have shown that simply adding non-spherical models to existing algorithms, while tempting, may not solve the problem. This is because the simple modification does not guarantee that the algorithm can stably select a suitable non-spherical model for the retrieval of dust aerosols, or a spherical model for the retrieval of non-dust aerosols [38,41]. Although all VIIRS algorithms consider the sea surface wind speed parameter, retrieval biases are still dependent on it. Previous studies have also found a similar problem [44,45]. The possible reason is that high wind speeds increase the reflectance of glint and foam, further enlarging the reflectance of the ocean surface [44]. Therefore, the aerosol signal accounts for a smaller proportion of satellite-observed reflectance, leading to greater AOD and AE bias. AOD and AE both have a relatively strong relationship with wind speed, which may be another possible reason why the retrieval bias is dependent on wind speed [46].

5. Conclusions

As the successor to sensors such as MODIS, three algorithm teams have developed different aerosol products to continue long-term MODIS aerosol observations. Over the ocean, all three algorithms can retrieve AOD, AE, FMF, and AODF properties. However, the validation of FMF and AODF is ignored by most studies. They are important and can better determine anthropogenic radiative forcing [1]. The main conclusions of the validation and comparison results in 2018 are: (1) SOAR total AOD retrievals (quality filtering) have a better validation accuracy. After quality and AOD loading filtering, the AE and FMF validation accuracies of the three algorithms are similar, but SOAR may be slightly better. (2) SOAR AODF meets the goal of GCOS (its retrieval stability is also necessary to be studied (e.g., [31])), while the AODF of AERDT and NOAA should be used with caution in the scientific study due to their lower validation accuracy. (3) The differences in AOD spatial patterns of the three VIIRS products are small, but the differences in AE spatial patterns are large. (4) Scattering angle and sea surface wind speed affect the retrieval bias. The suitable non-spherical dust model may reduce the dependence of retrieval bias on the scattering angle. Higher wind speeds result in greater ocean surface reflectance, which results in a reduced sensitivity of satellite-observed reflectivity to aerosol signals. This may explain the dependence of retrieval bias on wind speed. Since this study only used AERONET observation data in 2018 and on coasts and islands, the result may not represent the full view of the performance of these three ocean products. More robust studies will be carried out in the future.

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